

Ultracarbonaceous Antarctic Micrometeorites (UCAMMs)

*« Samples from the external regions of
the protoplanetary disk »*

Cécile Engrand¹ and colleagues²...

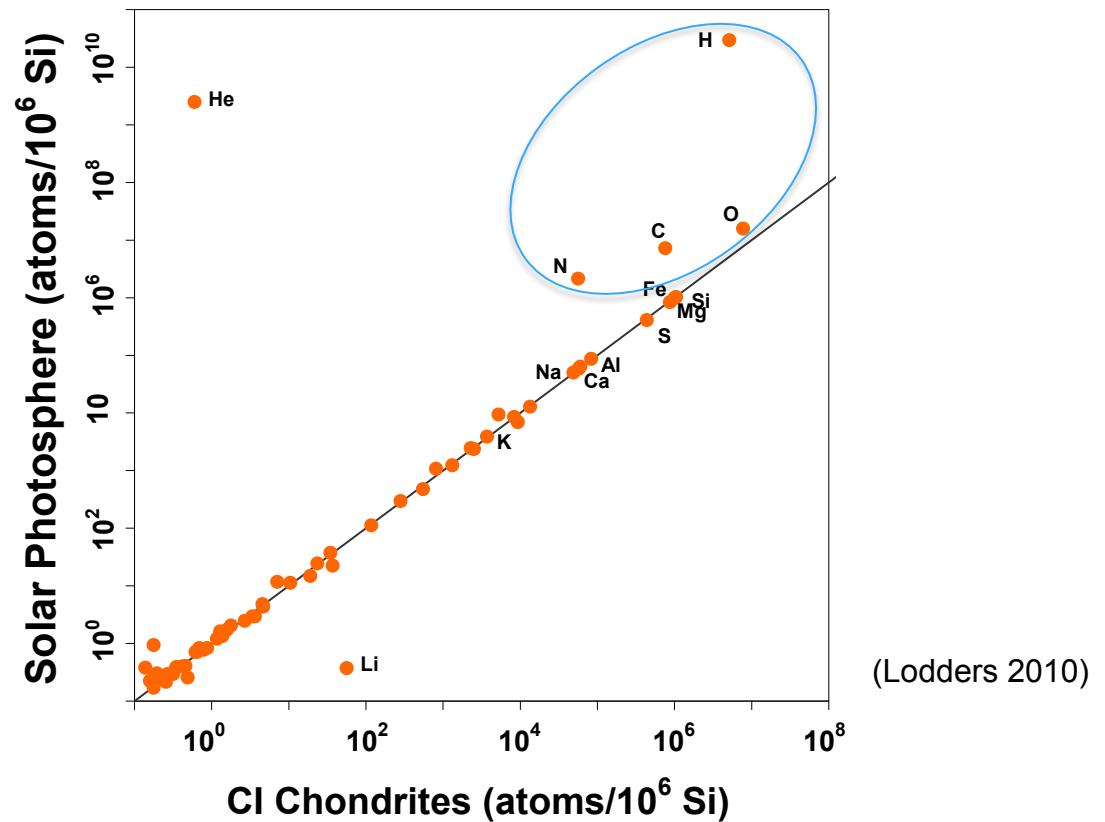
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²next slide...

...colleagues

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- IAS: E. Dartois
- Institut Curie: T.D. Wu, J.-L. Guerquin-Kern
- UMET: H. Leroux
- IMPMC: K. Benzerara, L. Remusat
- IPAG: E. Quirico, L. Bonal, F.R. Orthous-Daunay, V. Vuitton
- GANIL: B. Augé et al.
- Univ. New Mexico: E. Dobrica
- Univ. Hawai'i: J.P. Bradley, H. Ishii

Composition of the solar nebula?



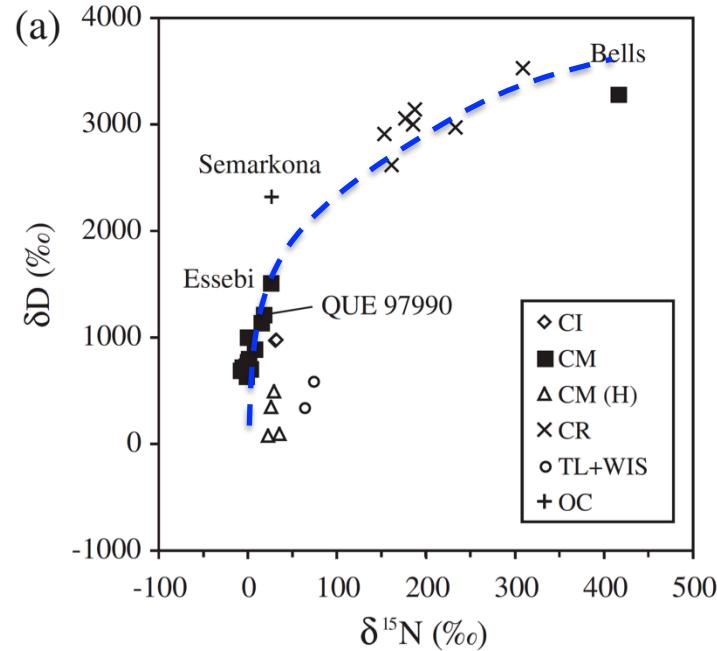
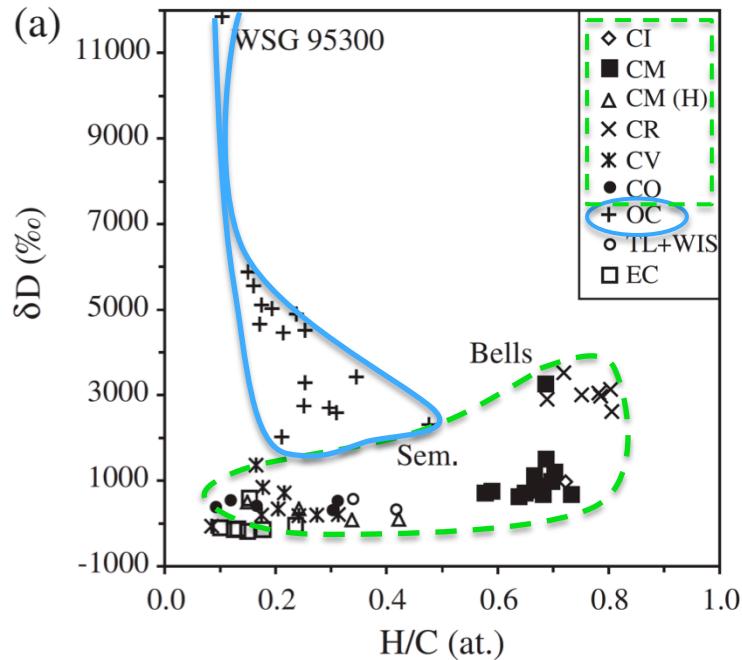
- Carbonaceous chondrites have the solar photosphere's composition
- They only represent ~ 5% of the meteorites
- The light elements are depleted wrt solar values
- Organic matter in extraterrestrial material?
- Comets?

Primitive organic matter in the Solar System

⇒ « insoluble organic matter »

- OM in meterites
 - Origin in Solar nebula vs ISM (presolar molecular cloud)
- OM in comets
 - remote sensing and in situ measurements (Stardust, Rosetta)
- OM in cometary* dust : IDPs and UCAMMs
 - Laboratory measurements of samples collected on Earth

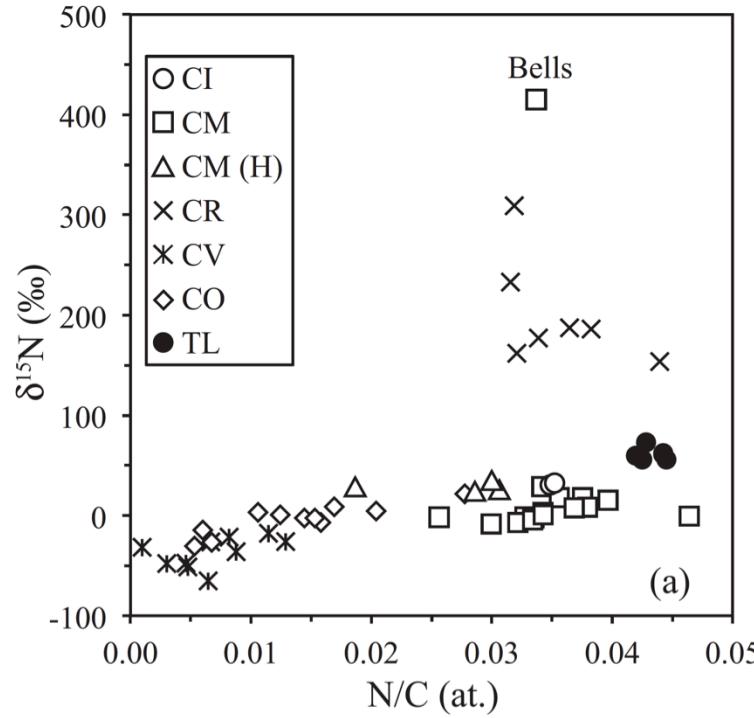
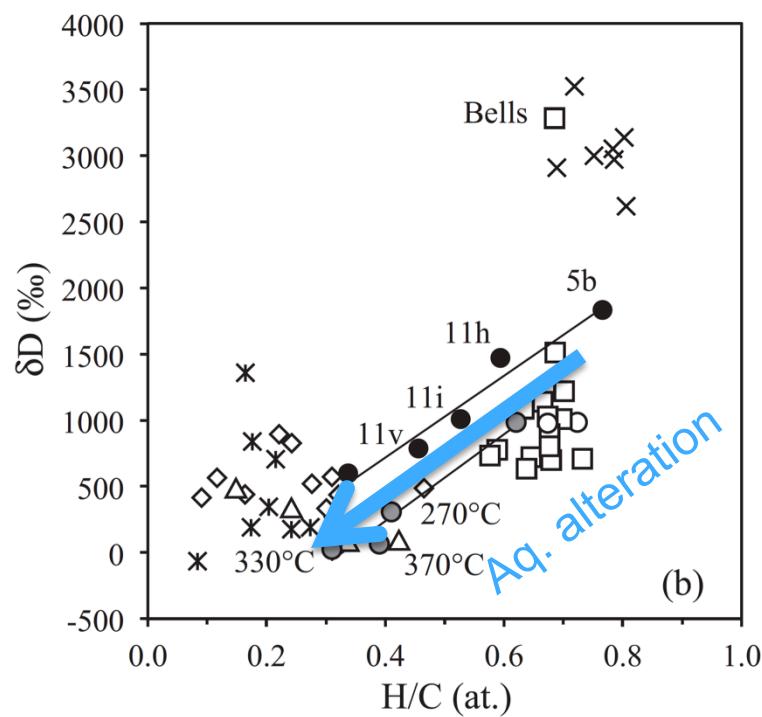
IOM in meteorites



Alexander+2007, 2010, 2014

- CCs : decrease of δD with thermal metamorphism
- UOCs: increase of δD with thermal metamorphism
- Correlation δD - $\delta^{15}\text{N}$?

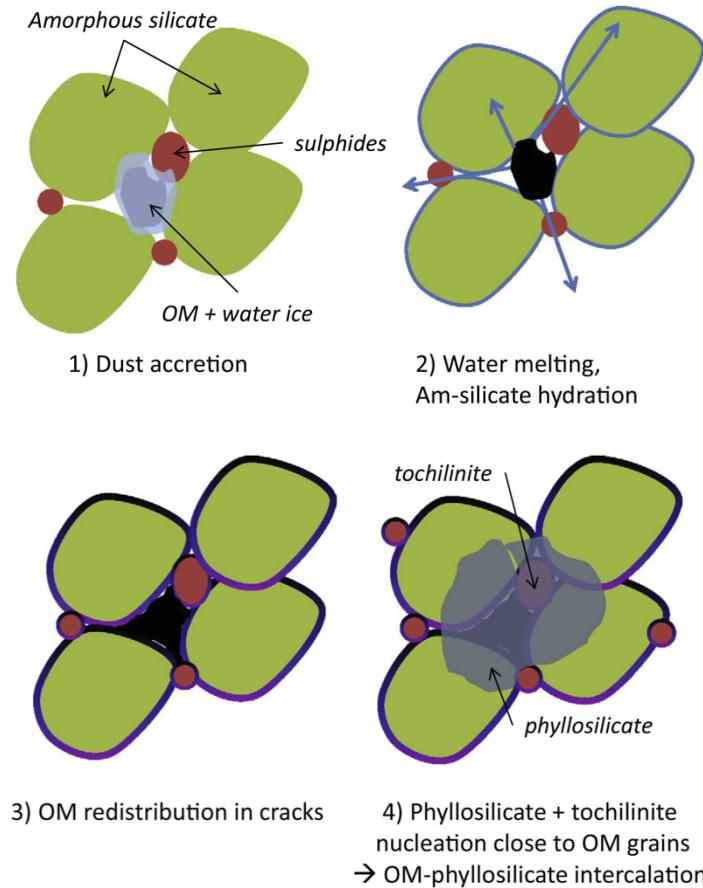
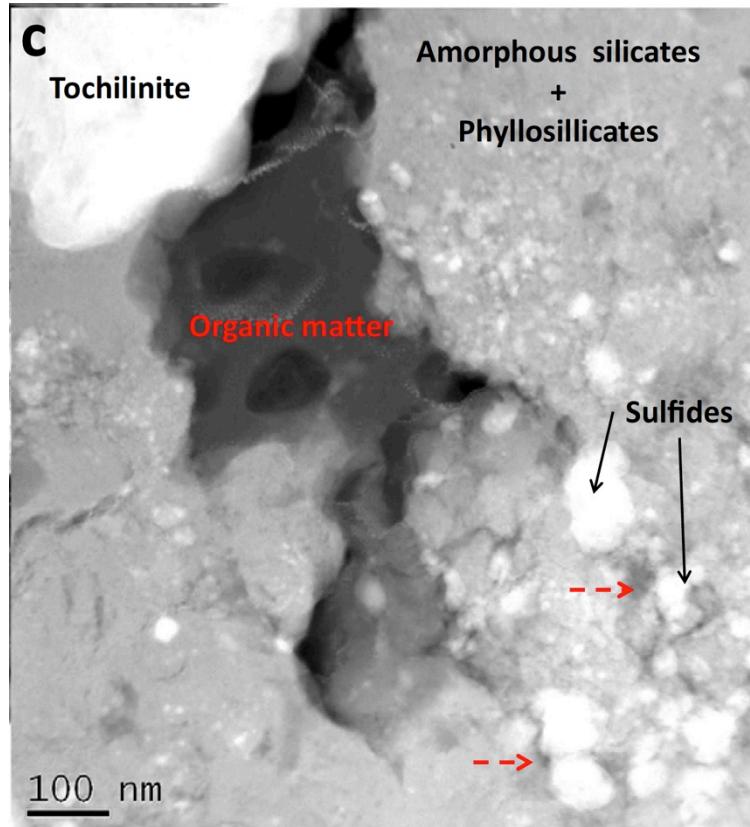
A unique IOM precursor?



Alexander+2014
Oba & Naraoka 2009

- CR-like IOM precursor + aqueous alteration?
(Alexander+2014)
- D fractionation from reduction of H_2O by Fe ?

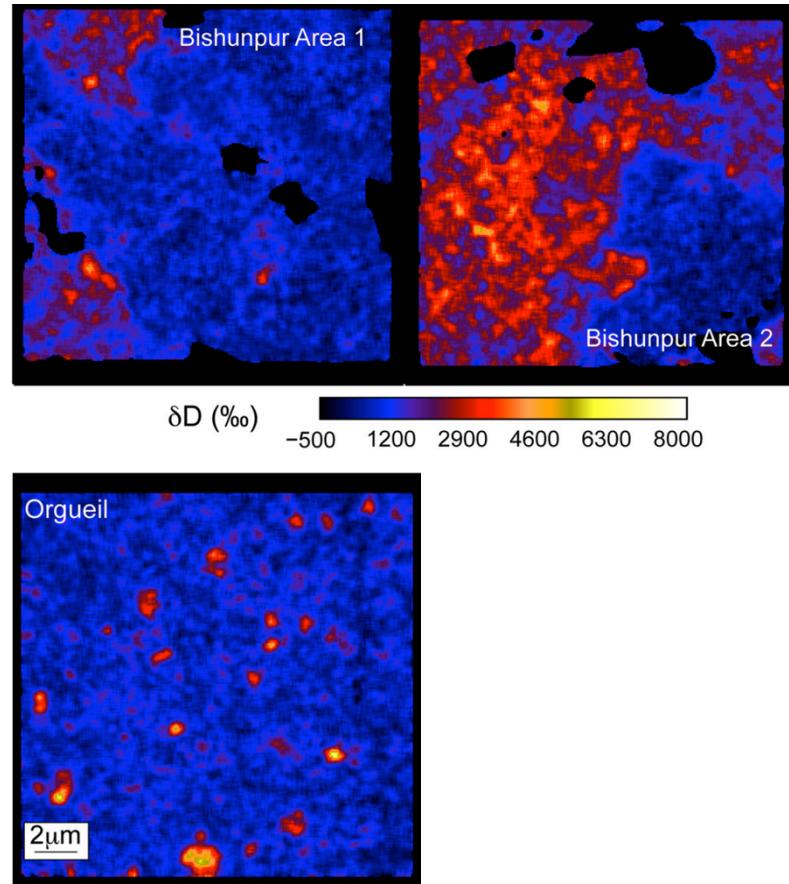
Ice-OM accretion? (CR meteorite)



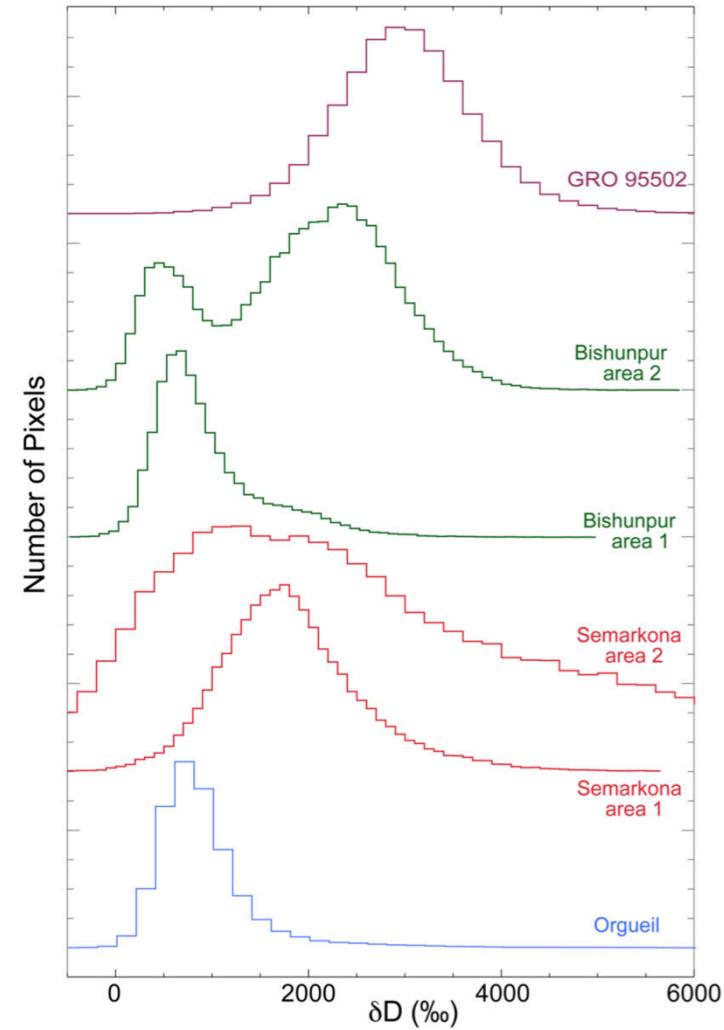
Le Guillou+2014 GCA

- Link OM – water? (OM – hydrated minerals association)

Two IOM phases in UOCs

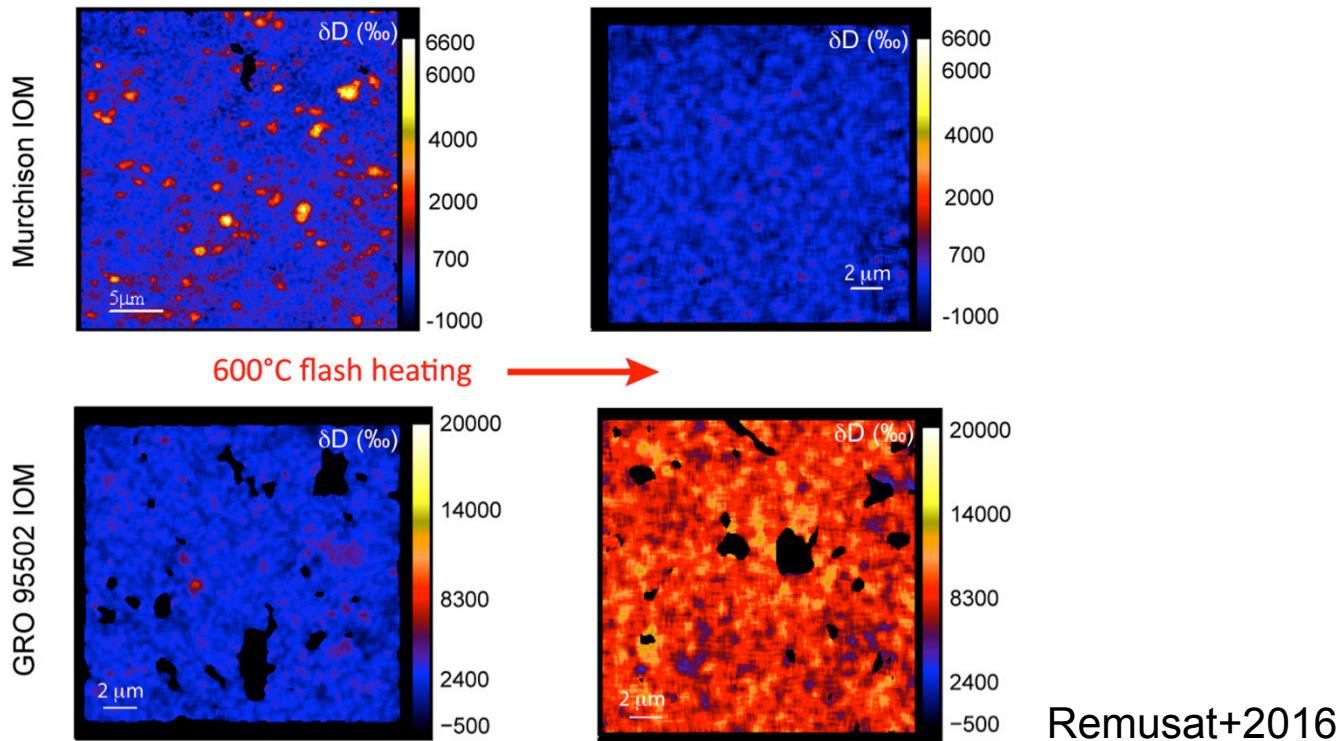


Remusat+2016



- Bimodal distribution of D/H in UOCs

D/H and thermal metamorphism

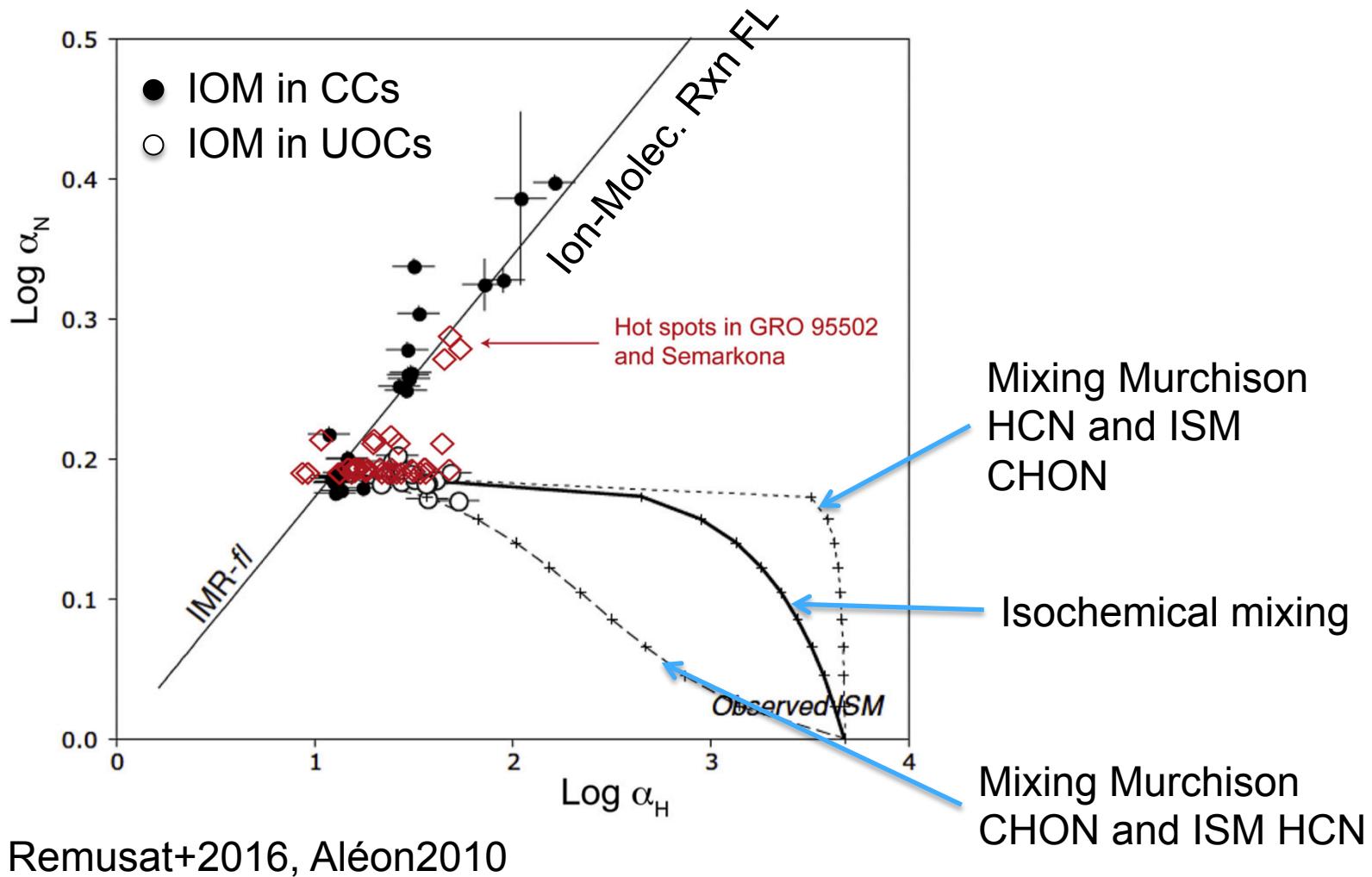


- Heating : Similar structural and molecular evolution for UOC and CC IOMs (structural reorganization and increase of aromatic/aliphatic carbon ratio).
- Heating : increase D/H for UOC; decrease D/H for Murchison (CC)
- **Not the same IOM precursor?**
- **ISM component in UOC IOM? (or preferential H loss?)**

Origin of N isotopic anomalies

- interstellar or cold molecular cloud origin via low-temperature ion–molecule reactions
(Terzieva & Herbst 2000, Rodgers & Charnley 2008, Aléon 2010)
 - ^{15}N enhancements in two dense prestellar cores (Hily-Blant+2013)
 - variations in N isotopic ratio across the Galaxy (Adande & Ziurys 2012)
- solar system origin via photochemical self-shielding (Clayton 2002; Lyons+2009)
 - Experimental reproduction of ^{15}N enrichments by VUV photolysis (Chakraborty+2013)

Mostly solar origin for CC and UOC IOM



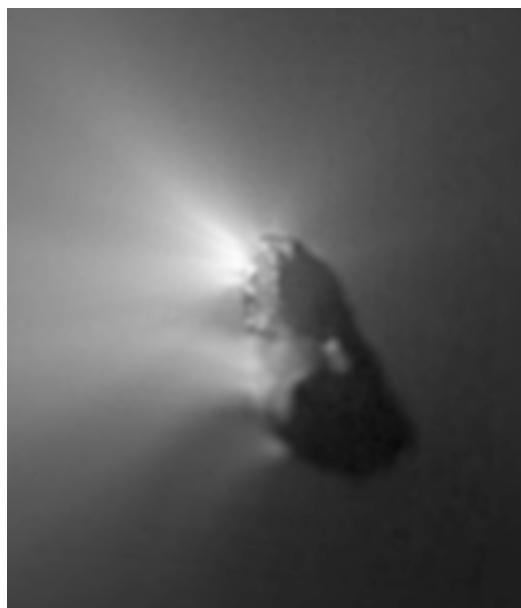
- UOCs: Mixing at the molecular level? (no hotspot with $D/H = 0.1 > 0.1\mu\text{m}$)
- Some hotspots for 2 UOCs on the IMR-fl (presolar cloud or outer regions)

IOM in meteorites

- UOC IOM :
 - only D-rich organics
 - homogeneous $\delta^{15}\text{N} \sim 0$
 - very few ^{15}N anomalies w/ large D-enrichments (IMR)
 - 2 origins? (solar system + ~ 1% presolar cloud?)
- CC IOM:
 - Accretion and mixing of both D- and ^{15}N -rich organics
 - Ion–molecule or grain surface chemistry in cold environments could explain both D- and ^{15}N -enrichments
- Comets?

Spatial exploration of comets

1986



Comet 1P/Halley
1986
Giotto/ESA

1992

Grigg-
Skjellerup
(Giotto)

2001

Borelly
(Deep
Space 1)

2004

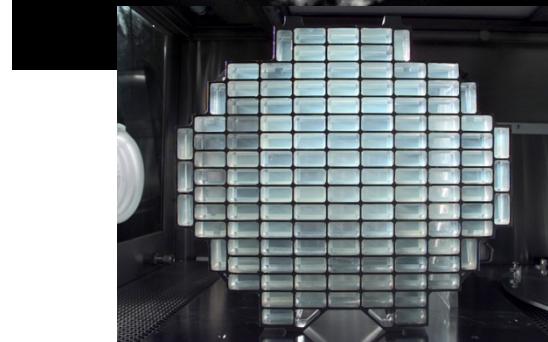
Wild 2
(Stardust)

2005

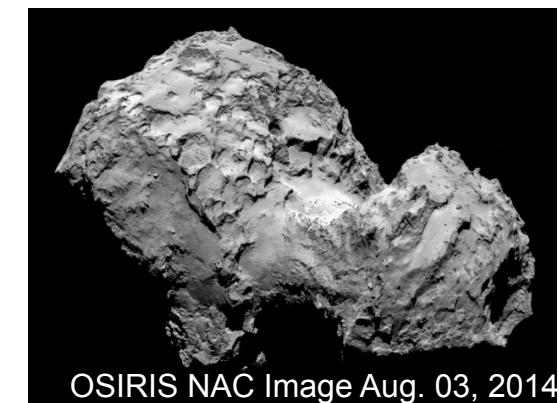
Tempel 1
(Deep
Impact)

2014-16

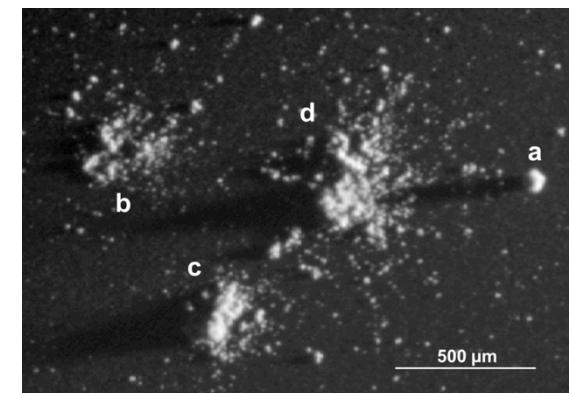
67P/C-G
(Rosetta)



Comet 81P/Wild 2
(6 jan. 2004)
Distance 500 km
Stardust/NASA



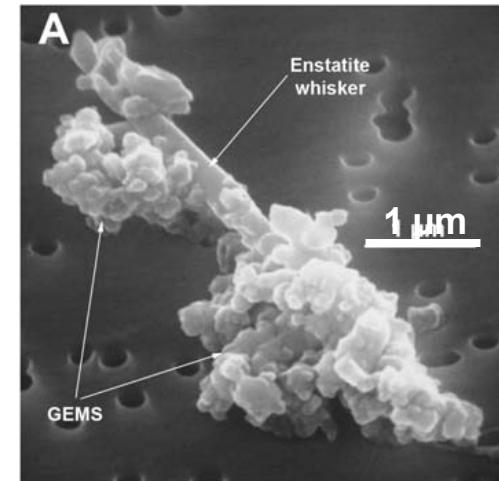
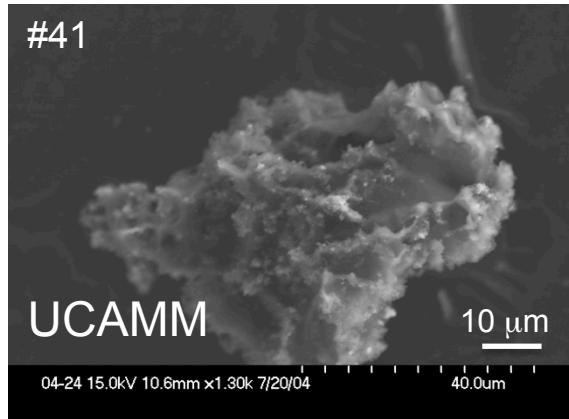
OSIRIS NAC Image Aug. 03, 2014



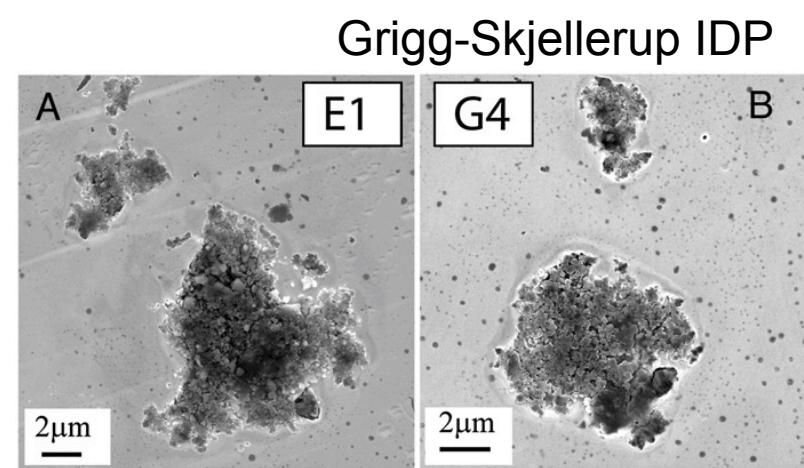
500 μm
Langevin+2016

Cometary dust collected on Earth?

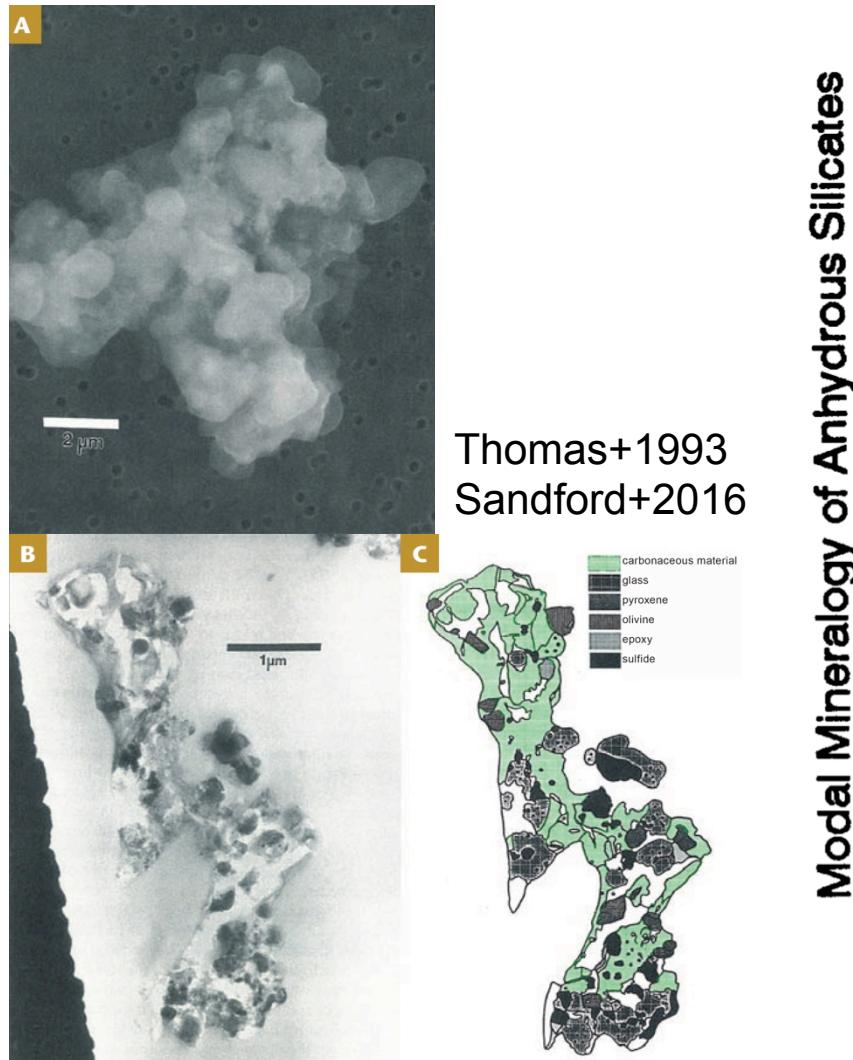
- Cosmic Dust flux ~ 30 000 ton/year
 - (Mass max ~ 200 μm)
- Theoretical dual origin : comets and asteroid
- ... ~ 90% from JF comets (Nesvorný+2010, Poppe2016...)
- Collection in stratosphere (IDPs) and Antarctica (MMs)
- Cometary dust:
 - IDPs : CP-IDPs & Timed collection (26P/G-S, 21P/G-Z)
 - UCAMMs



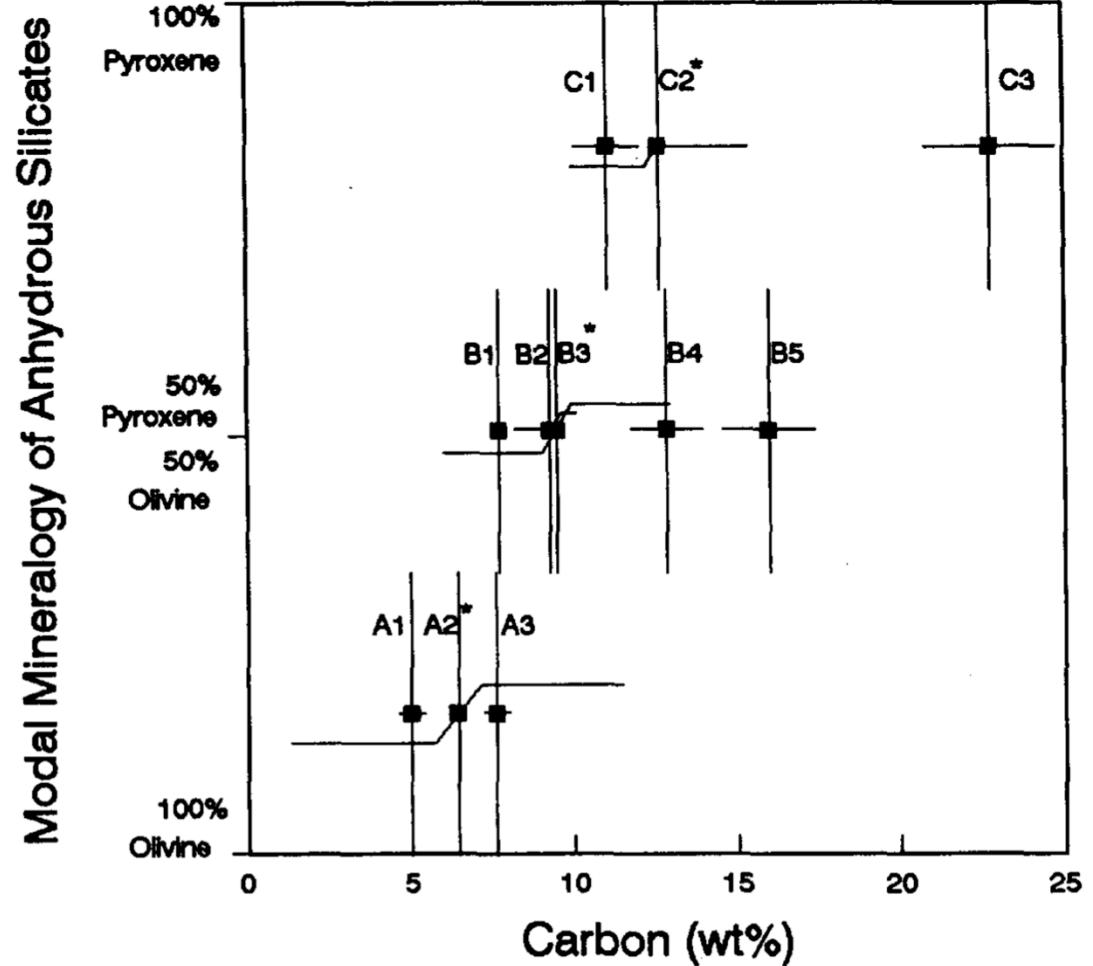
Ishii+2008



CP-IDPs are C-rich & Px-dominated



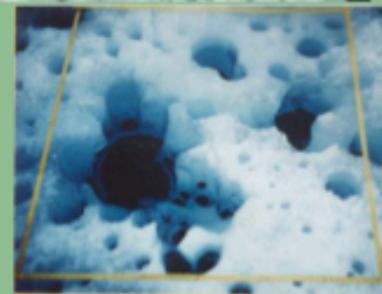
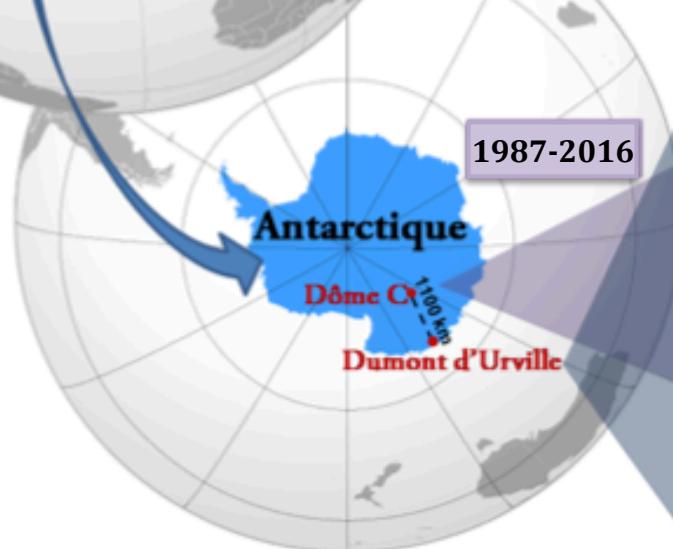
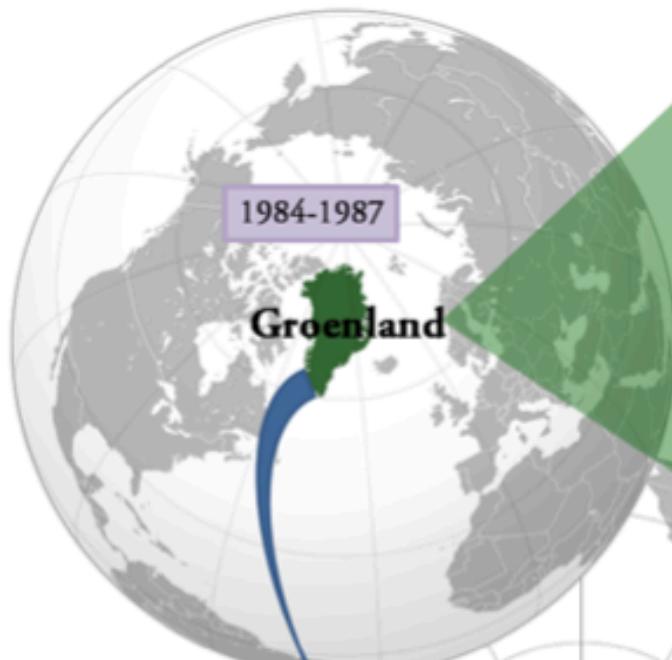
Thomas+1993
Sandford+2016



Thomas+1993

- A cometary origin of CP-IDPs is proposed

French Micrometeorite collections

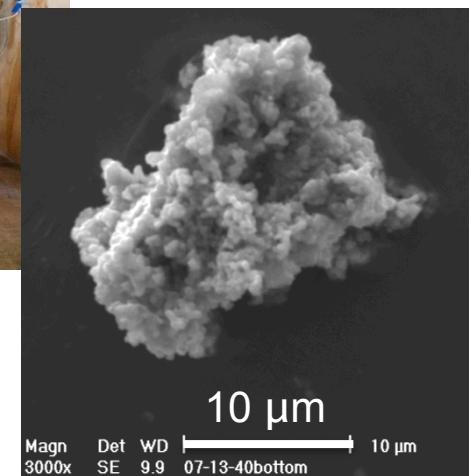


Maurette+ Nature 1987, 1991
Duprat+ ASR 2007

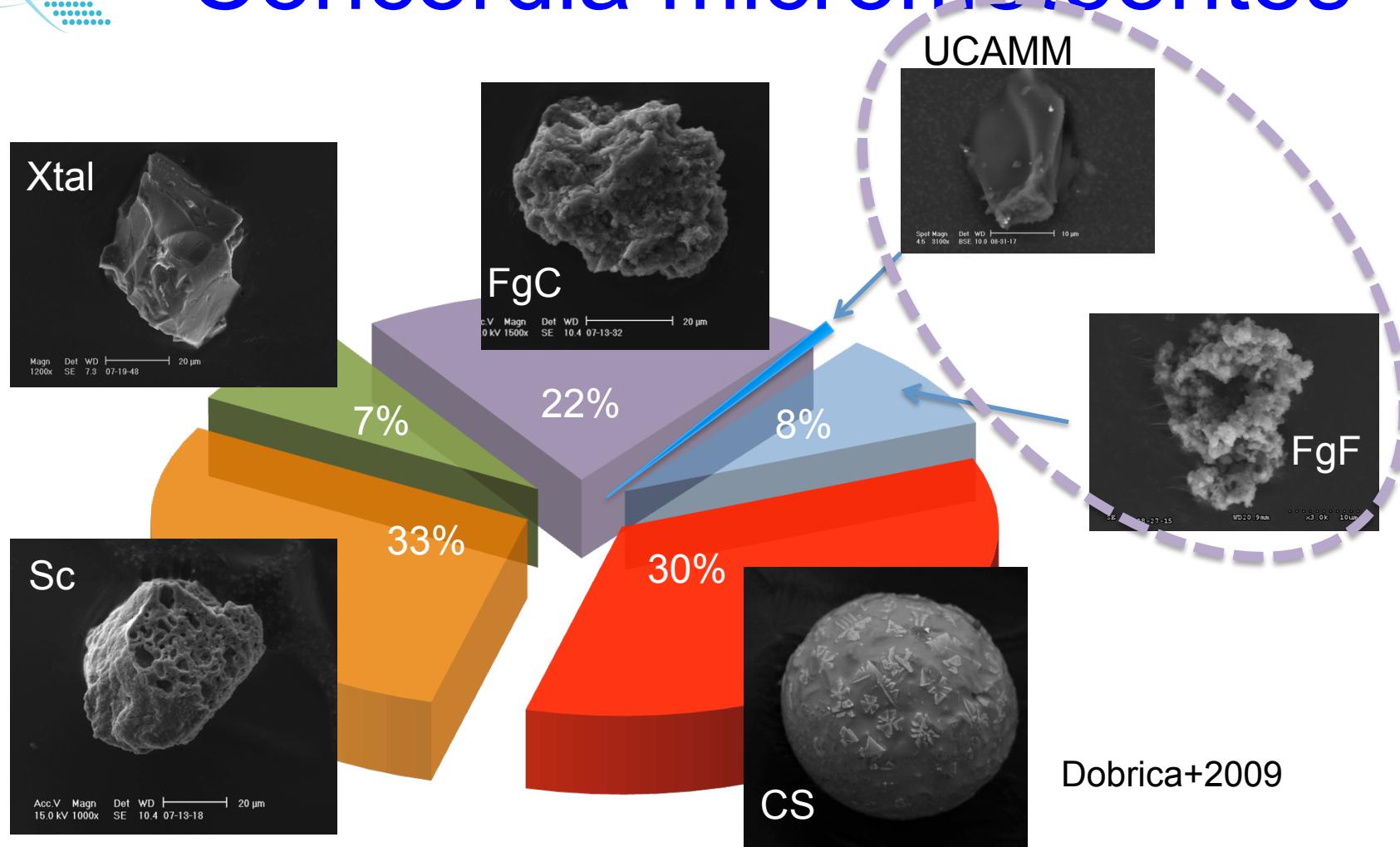
Micrometeorites @ CONCORDIA

Micrometeorites from central Antarctic snow

- **CONCORDIA Collection (CSNSM)**
 - > 5000 interplanetary dust particles CONCORDIA (75°S, 123°E) (IPEV-CNRS)
 - **Very good preservation**
 - **2000, 2002, 2006, 2014 & 2016 field campaigns**



Concordia micrometeorites



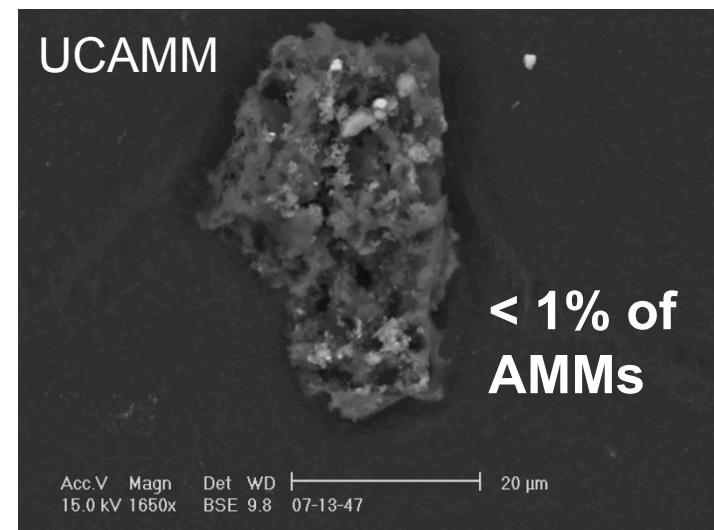
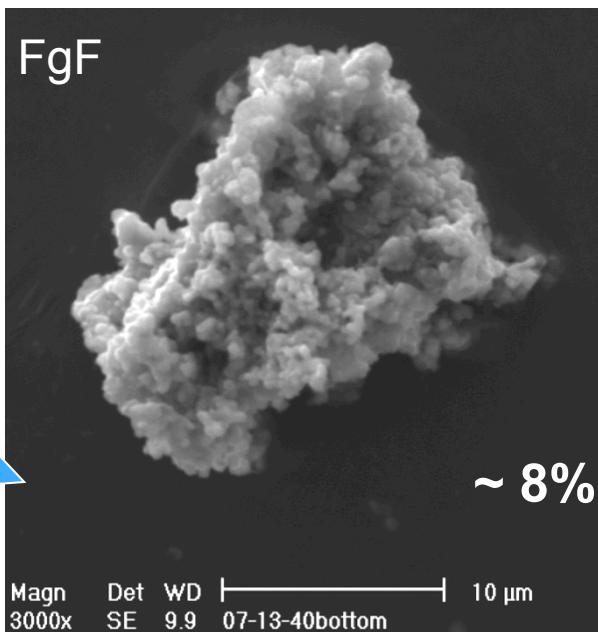
- size range: 20-250 μm
- Average size $\sim 45 \mu\text{m}$

The CONCORDIA MM collection

- Two new families of particles:

Friable MMs (FgF, Fine-Grained Fluffy) = CP-IDPs?

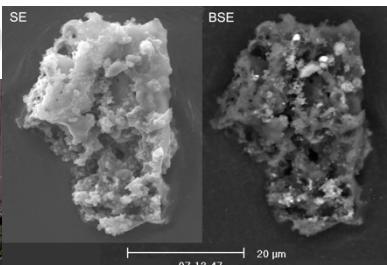
- Ultracarbonaceous Antarctic MMs (UCAMMs)



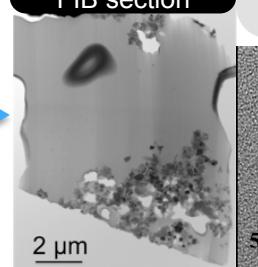
Micromanipulation



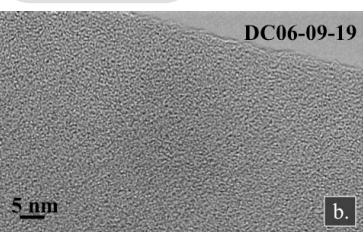
SEM + EDX



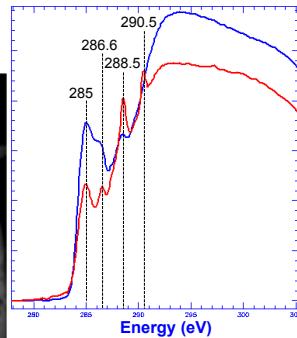
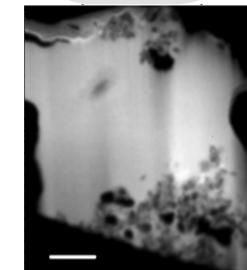
FIB section



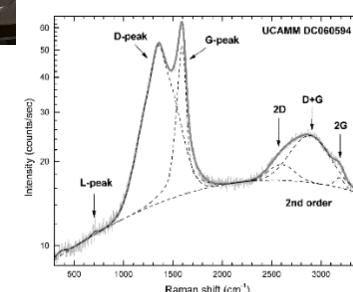
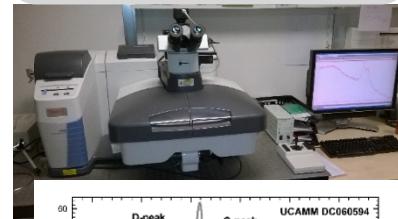
TEM



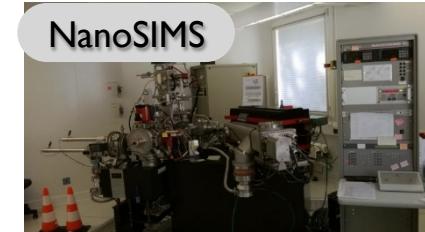
STXM + XANES



Raman



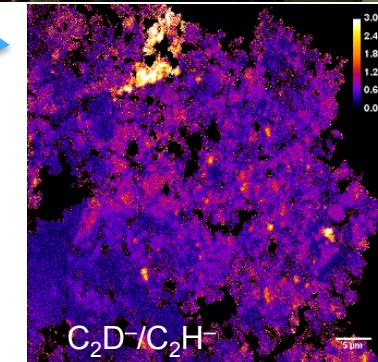
NanoSIMS



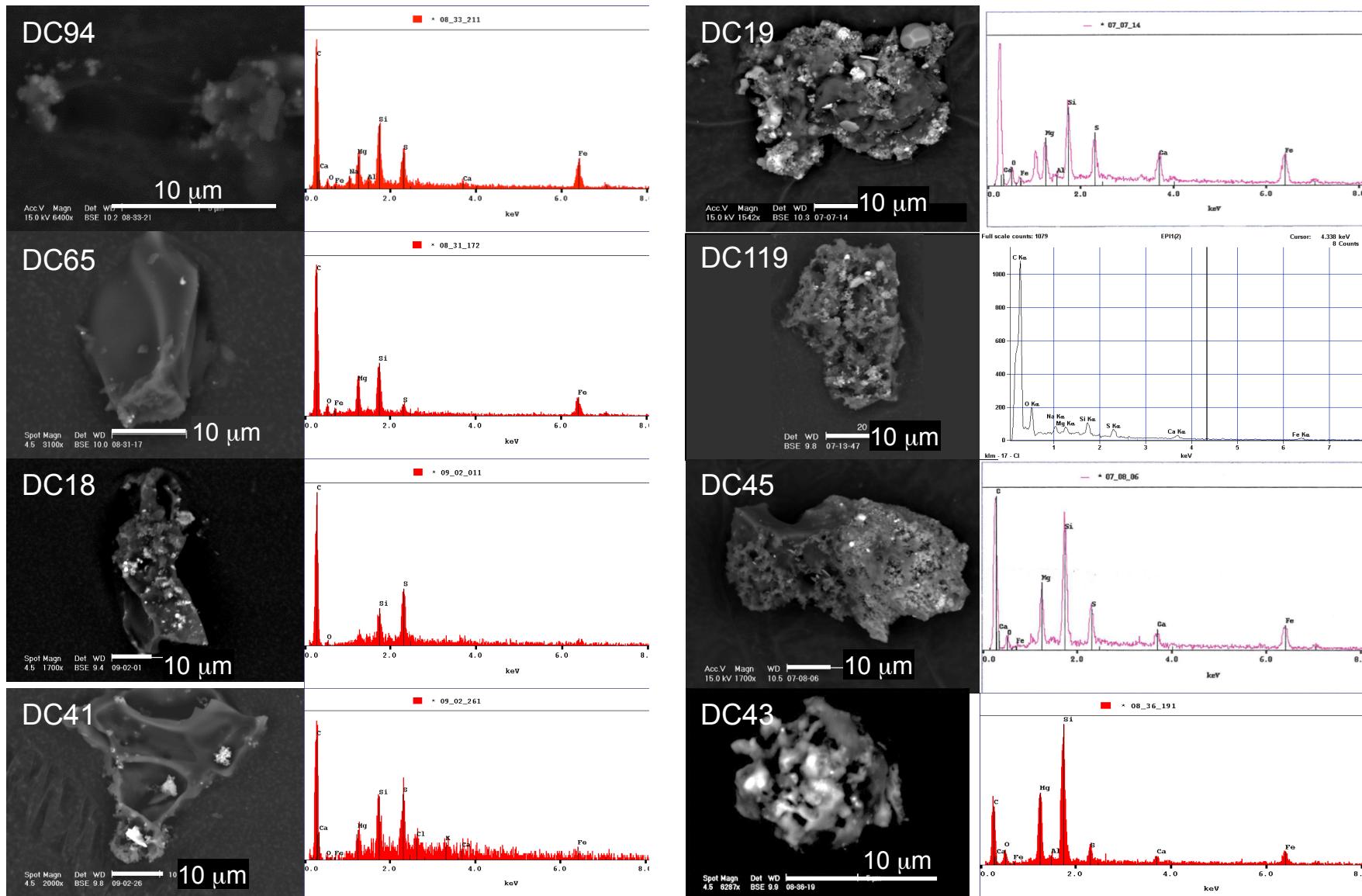
EMPA



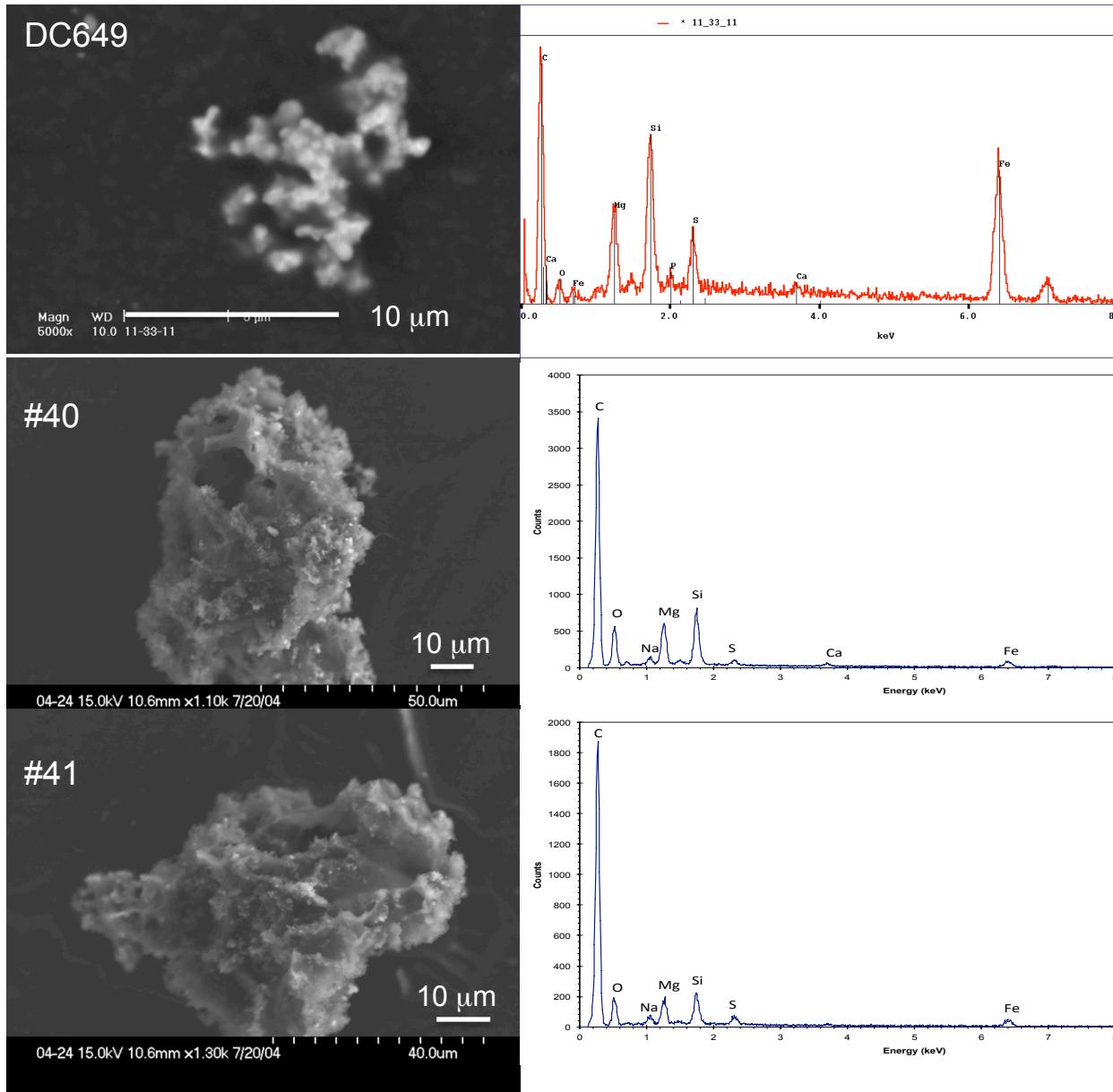
Noble Gas (Ne)



11 CONCORDIA UCAMMs



11 CONCORDIA UCAMMs (cont'd)

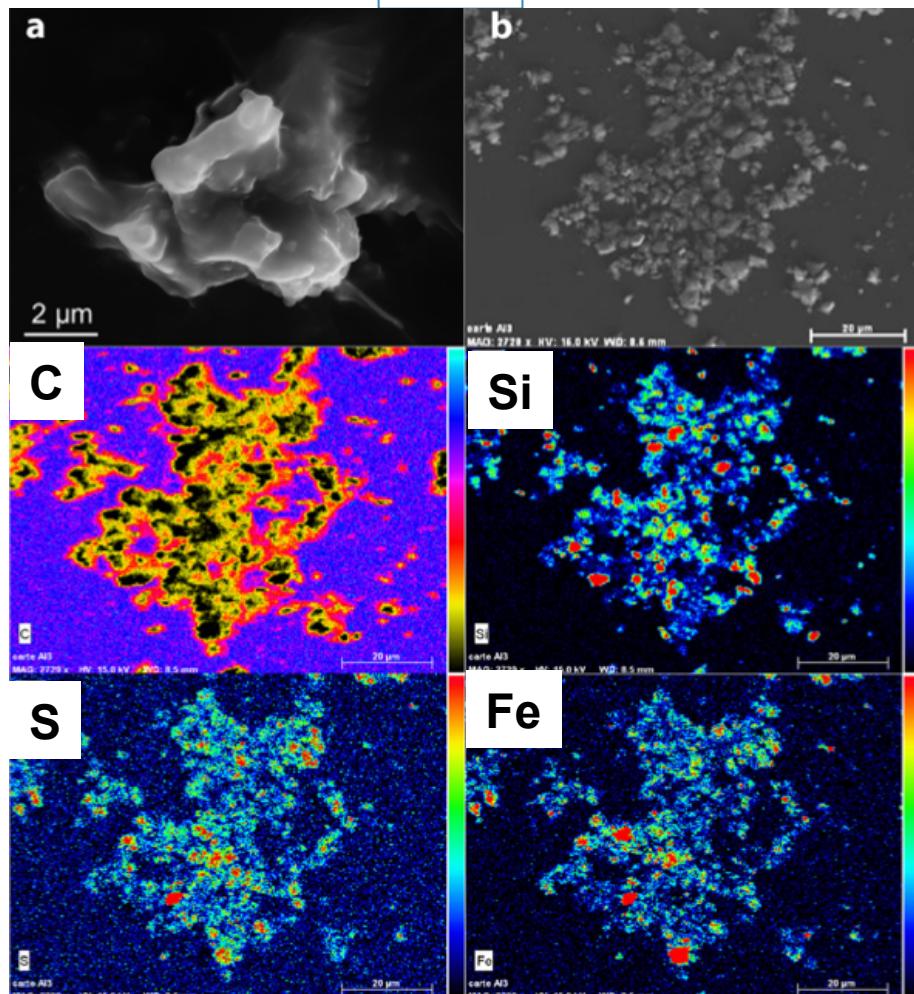


And still searching...

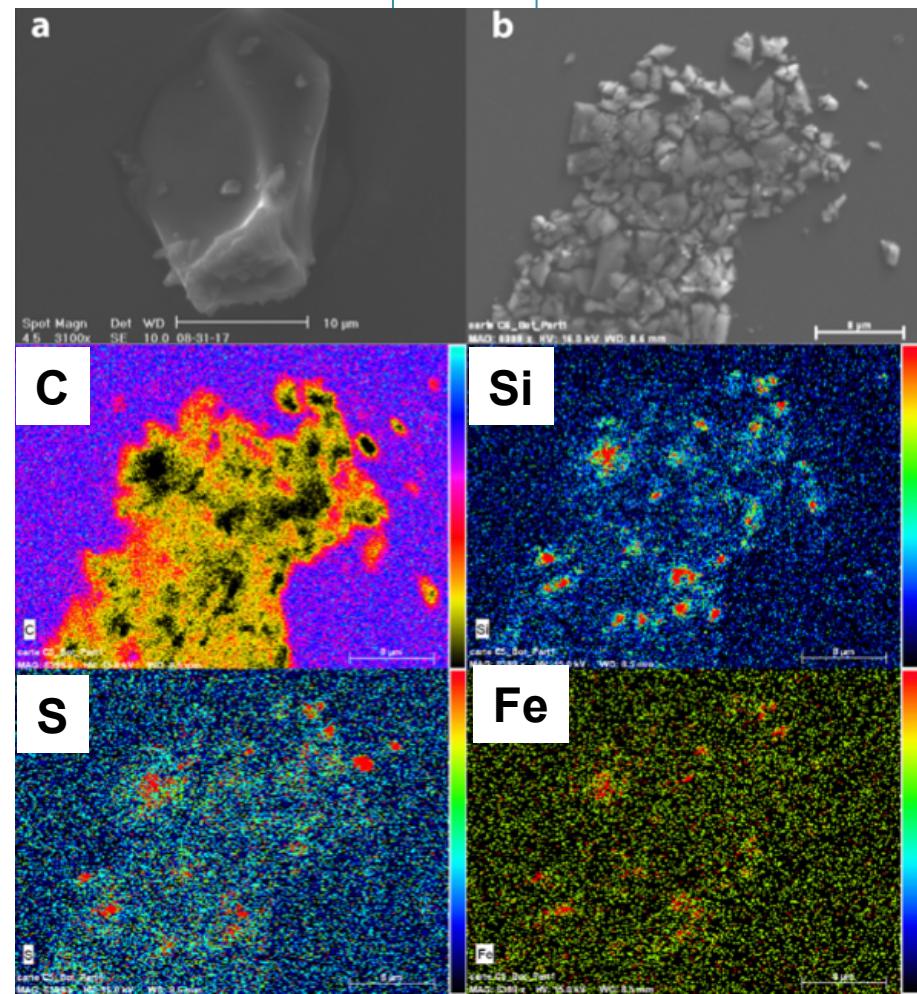
UCAMMS for IR and SIMS



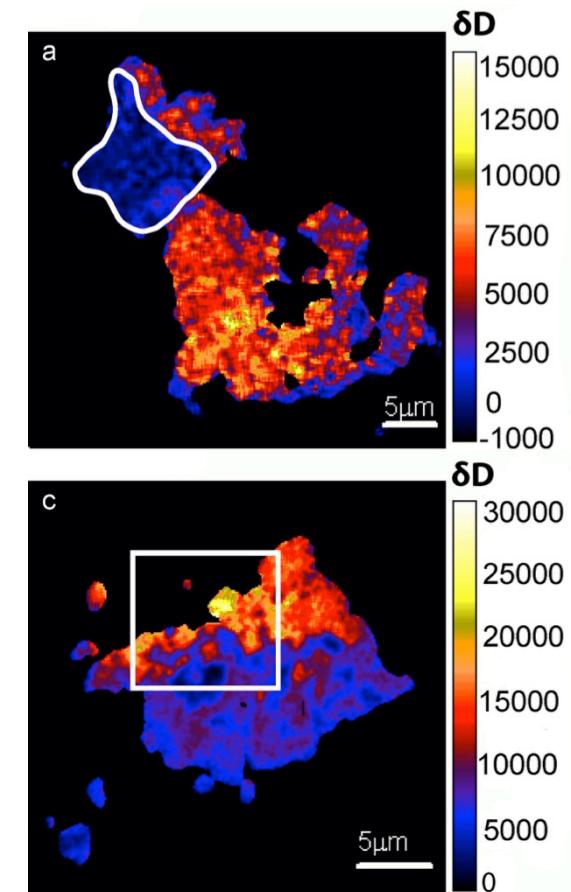
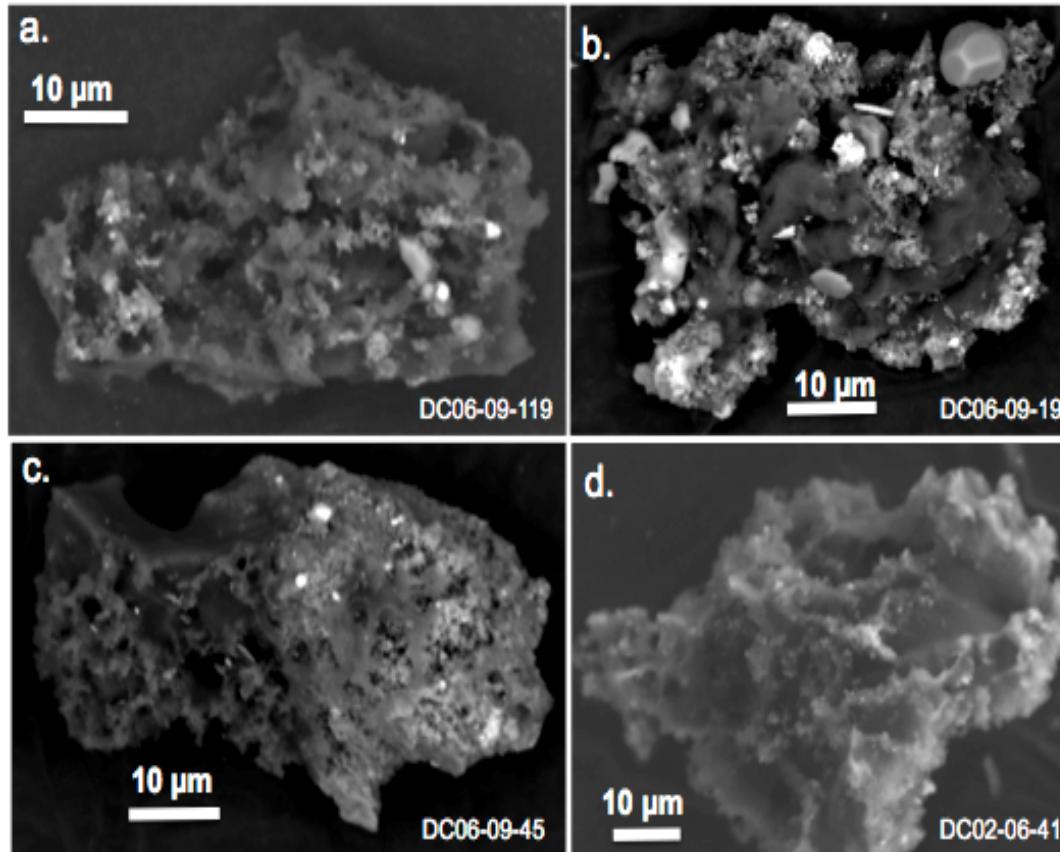
DC94



DC65



UCAMMs – Hydrogen isotopes (NanoSIMS)



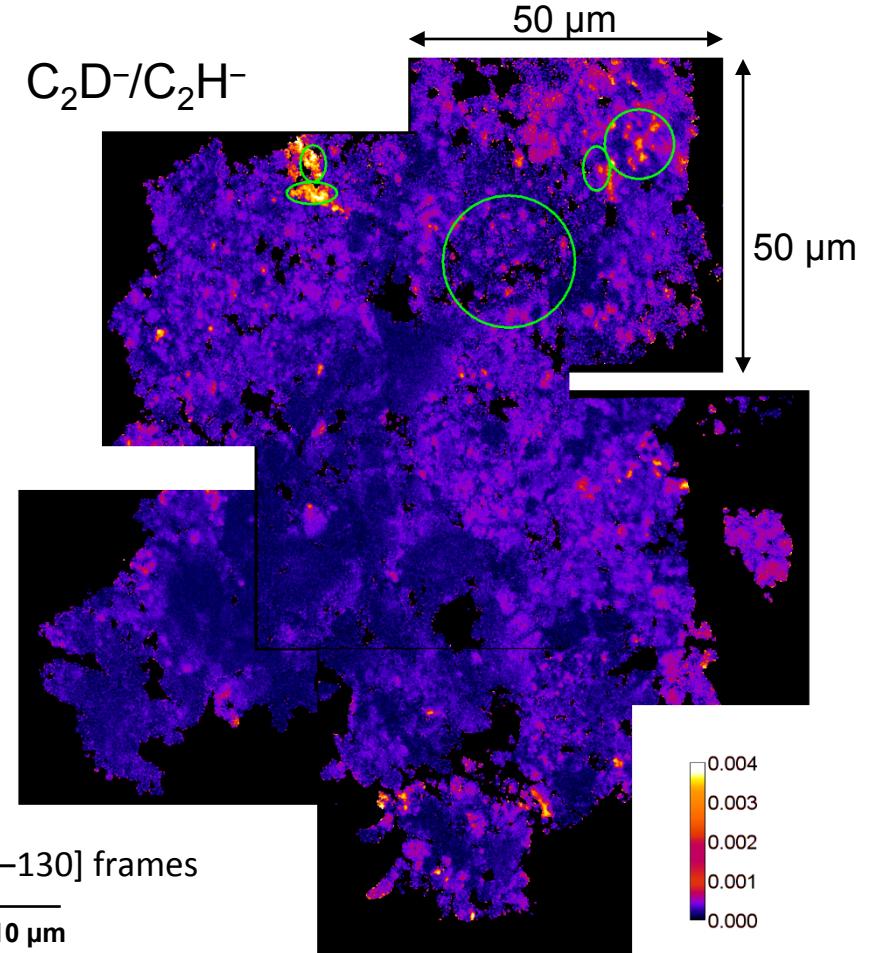
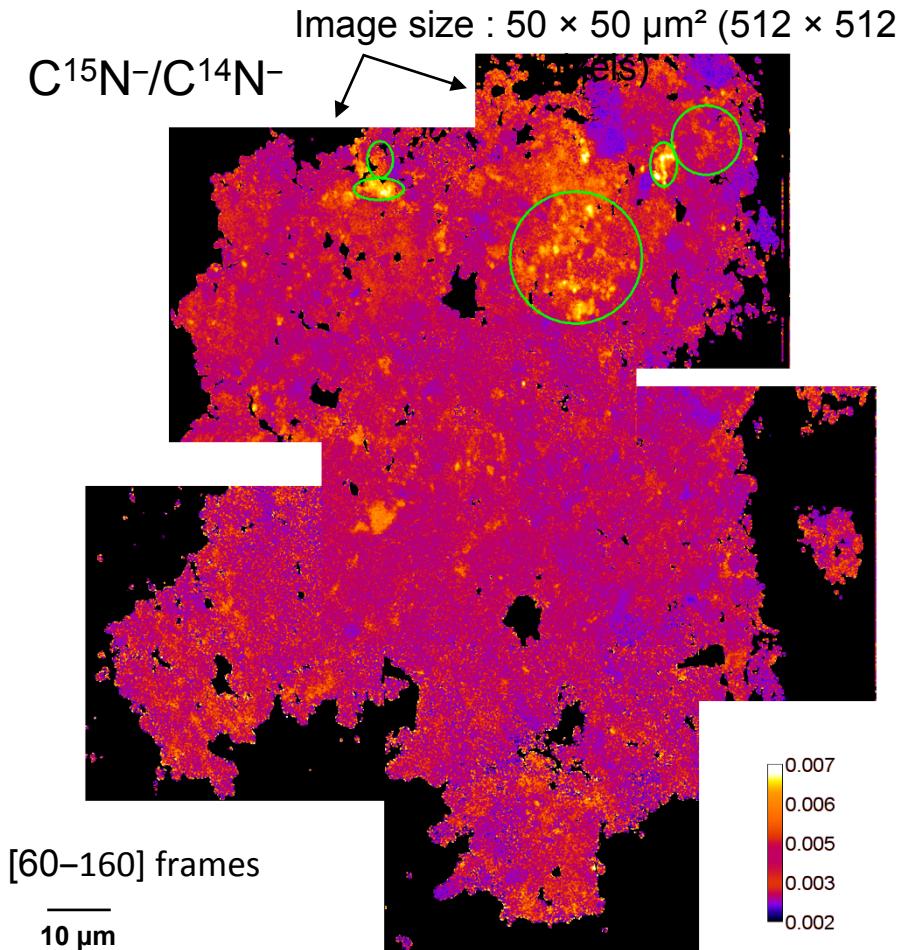
Duprat+2010 Science

- A globally D-rich carbonaceous phase ($D/H_{bulk} \sim 5$ to 10 times SMOW)
- Maximum D/H ratios up to 30x terrestrial value (i.e. $4.6 \pm 0.5 \times 10^{-3}$)

$\text{C}^{15}\text{N}^-/\text{C}^{14}\text{N}^-$ and $\text{C}_2\text{D}^-/\text{C}_2\text{H}^-$ isotopic ratio images on DC94



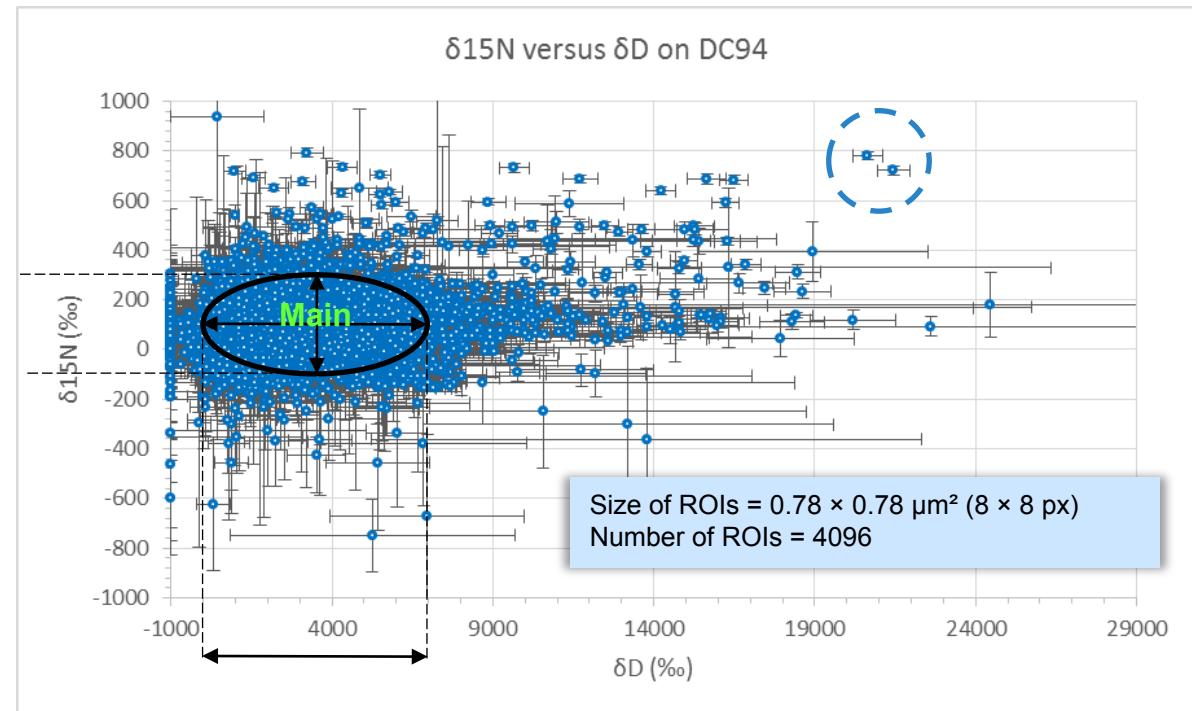
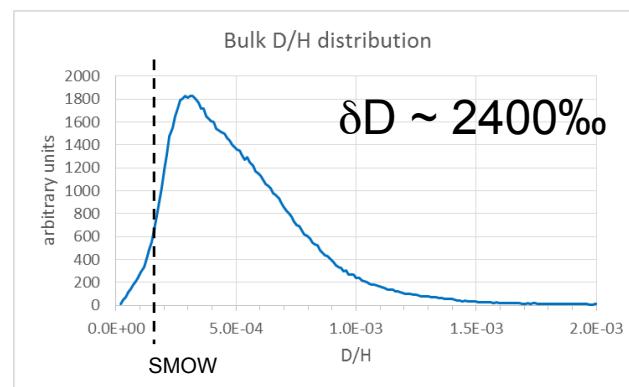
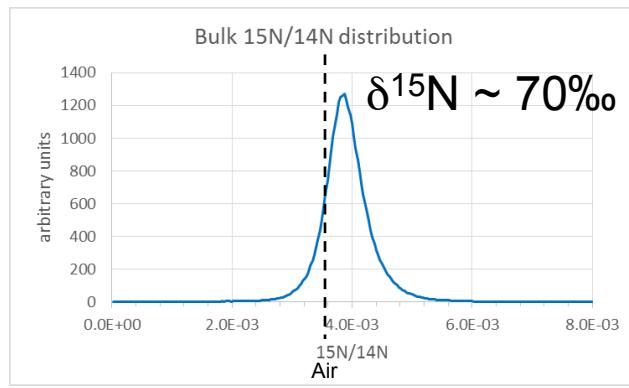
It is possible to compare the $\text{C}^{15}\text{N}^-/\text{C}^{14}\text{N}^-$ and $\text{C}_2\text{D}^-/\text{C}_2\text{H}^-$ isotopic ratio images



(Bardin+ MetSoc 2015)

- DC94 fragment has heterogeneous N and H isotopic compositions

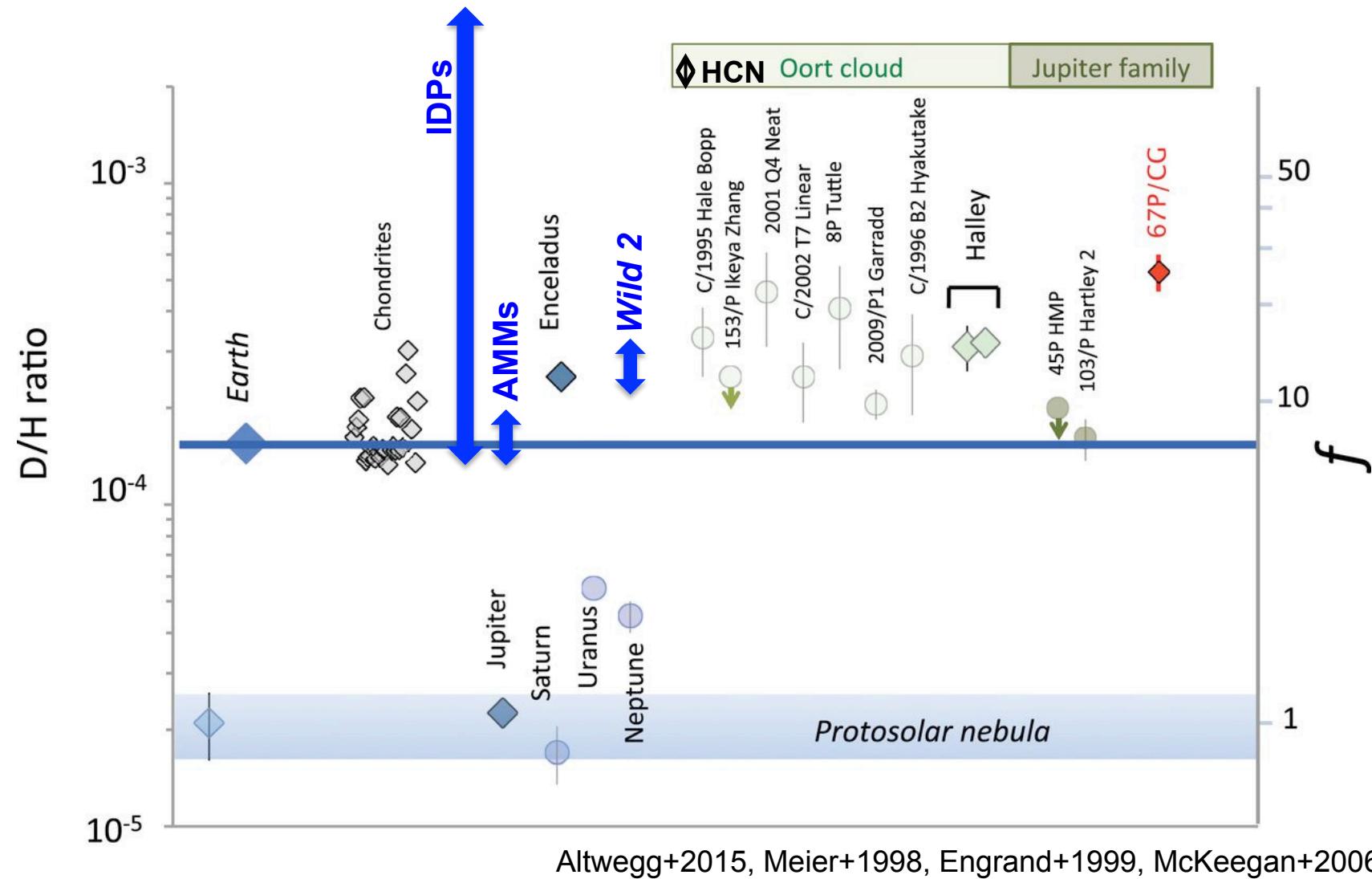
Search for spatial correlations between $\delta^{15}\text{N}$ and δD



(Bardin+ MetSoc 2015)

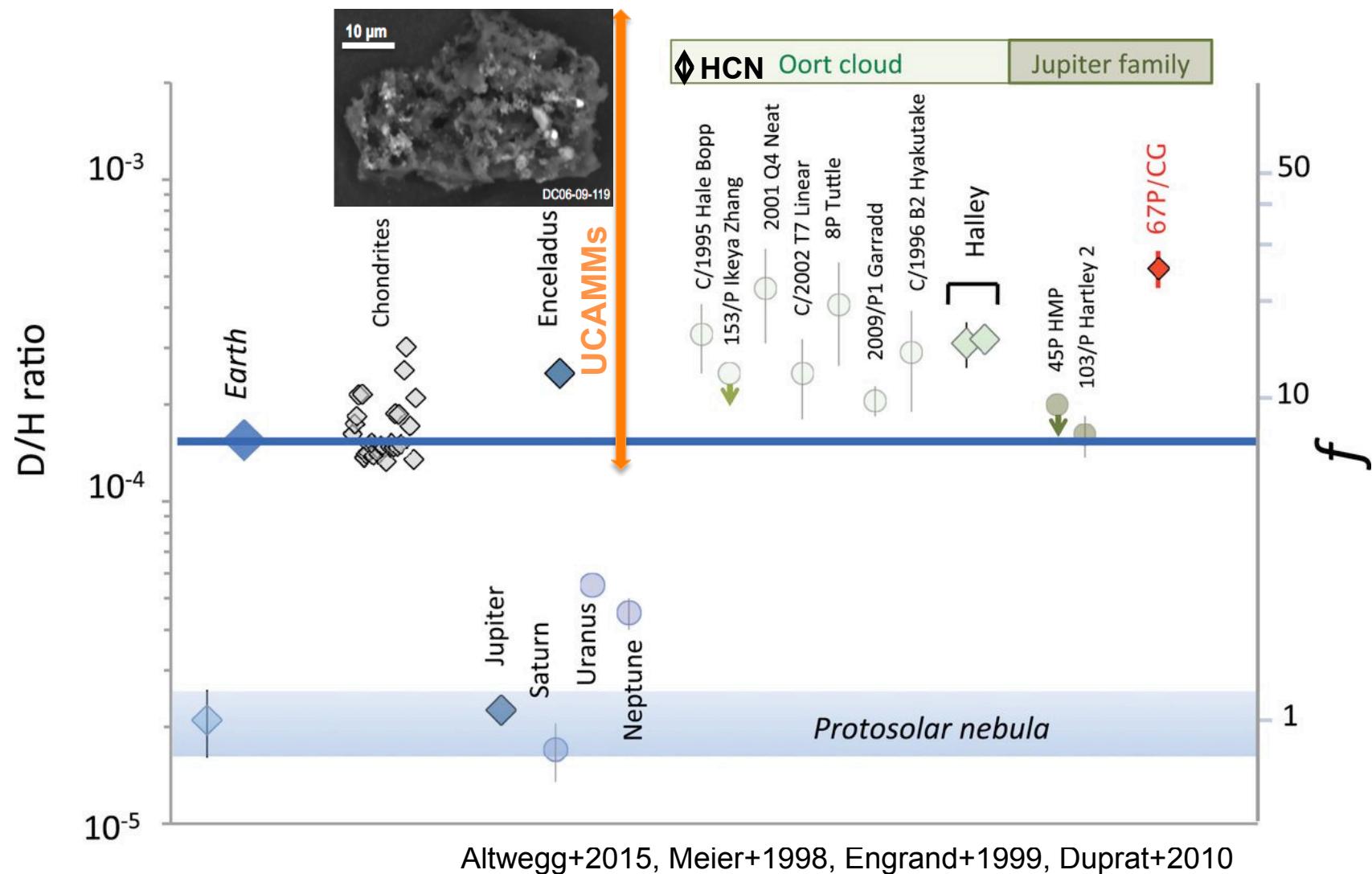
- main OM : large variations in δD and moderate variations in $\delta^{15}\text{N}$
- No clear correlation between δD and $\delta^{15}\text{N}$ for most of the grain

D/H ratios



- Caveat : comparing D/H in water and organics

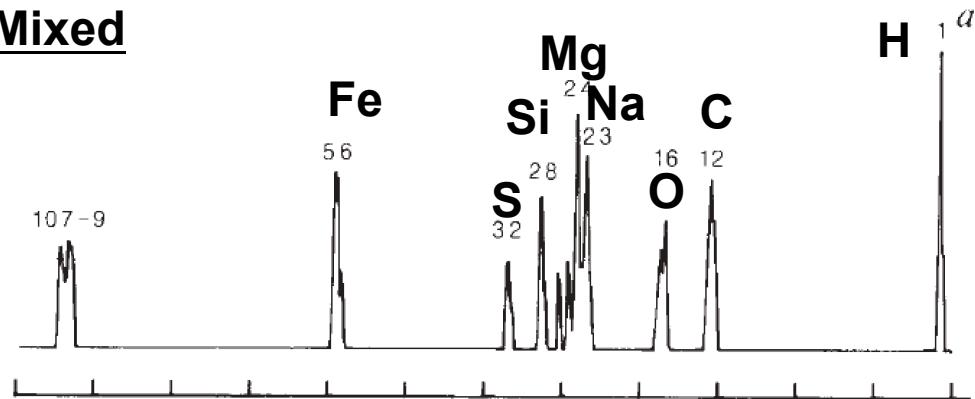
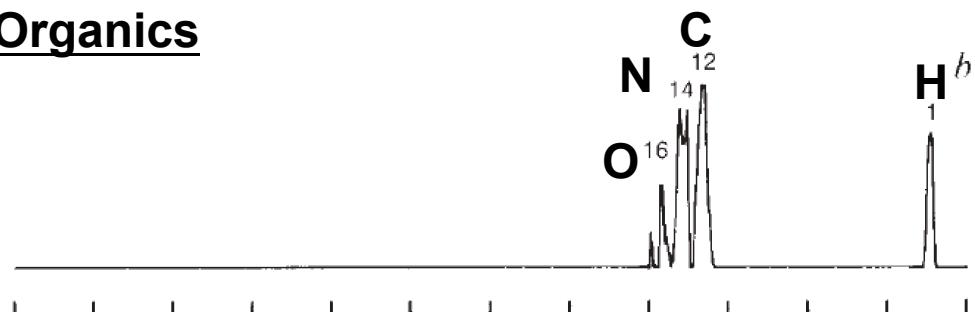
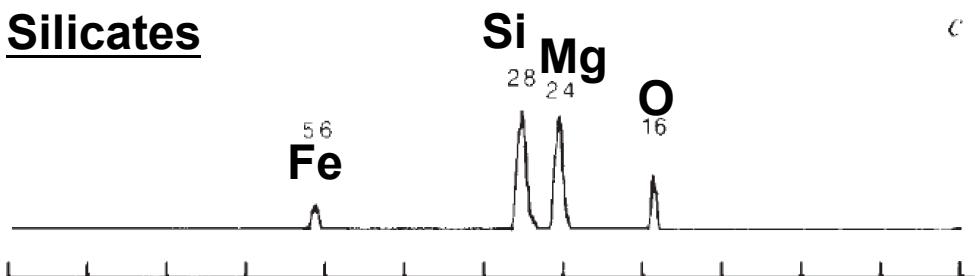
D/H ratios



- Caveat : comparing D/H in water and organics

1986

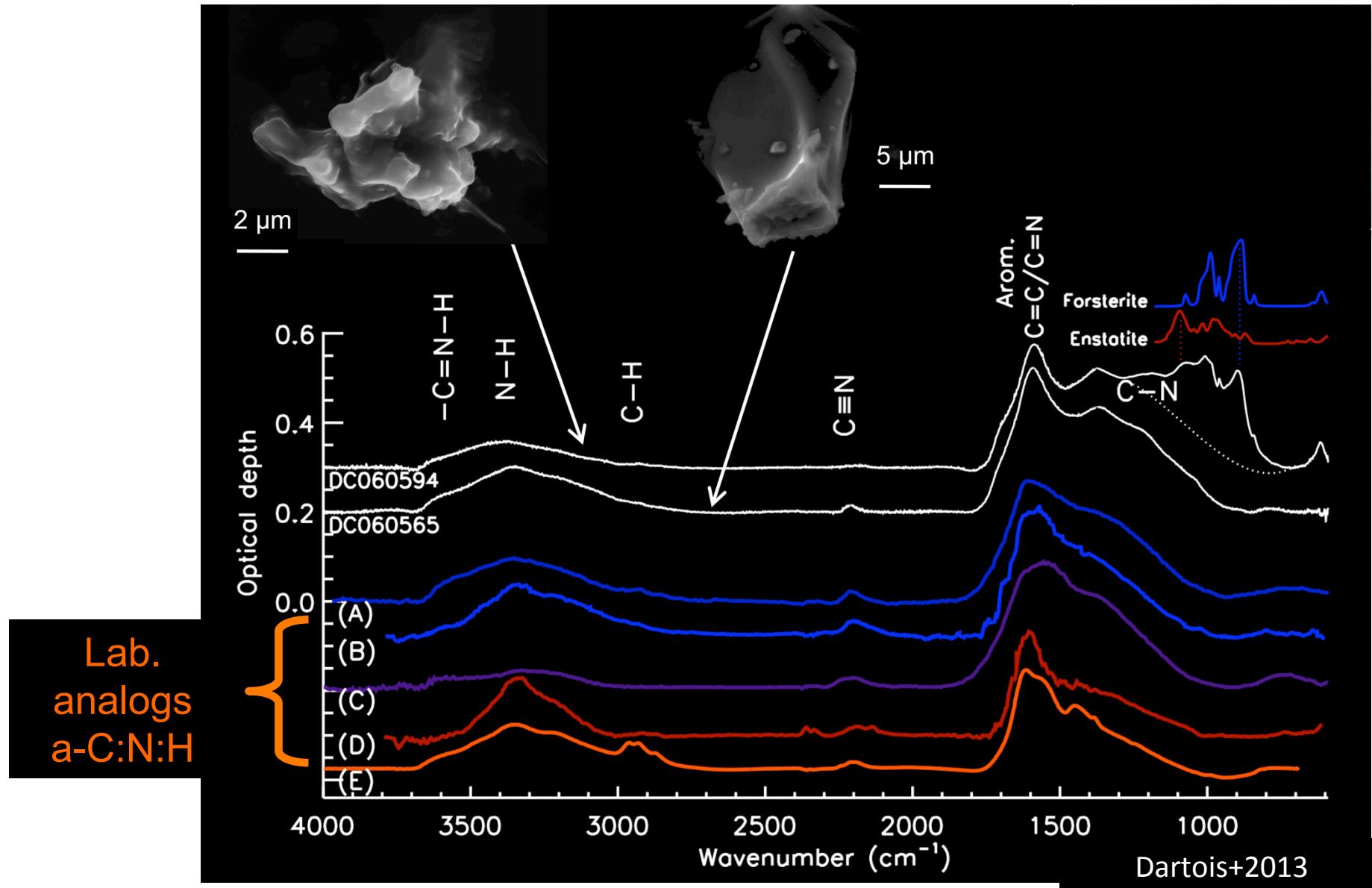
Comet Halley

80% of
the spectraMixedOrganicsSilicates

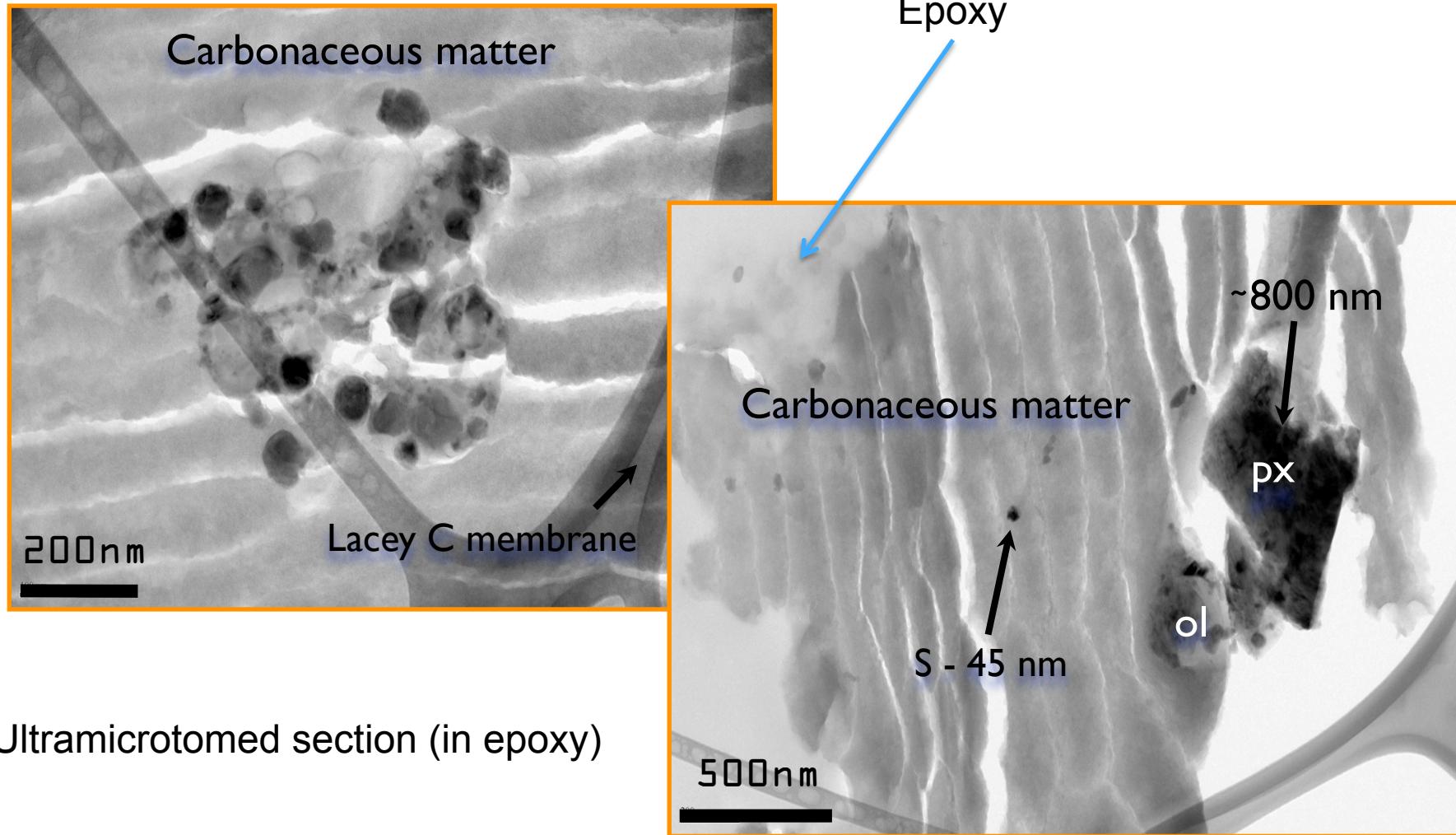
(PIA data)

(Kissel et al. 1986)

UCAMMs : a N-rich « CHON » ...

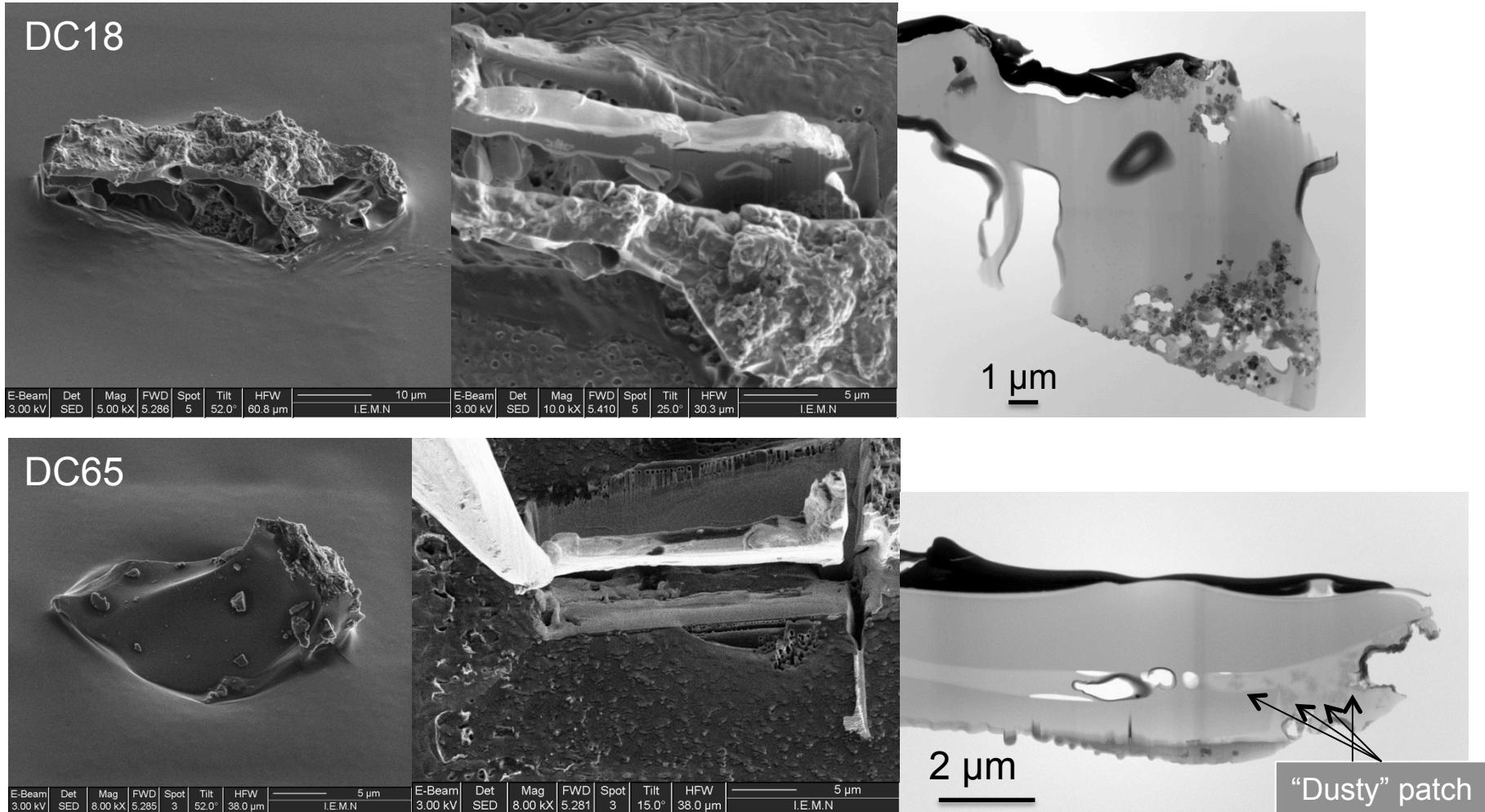


Intimate OM-minerals mixing in UCAMMs



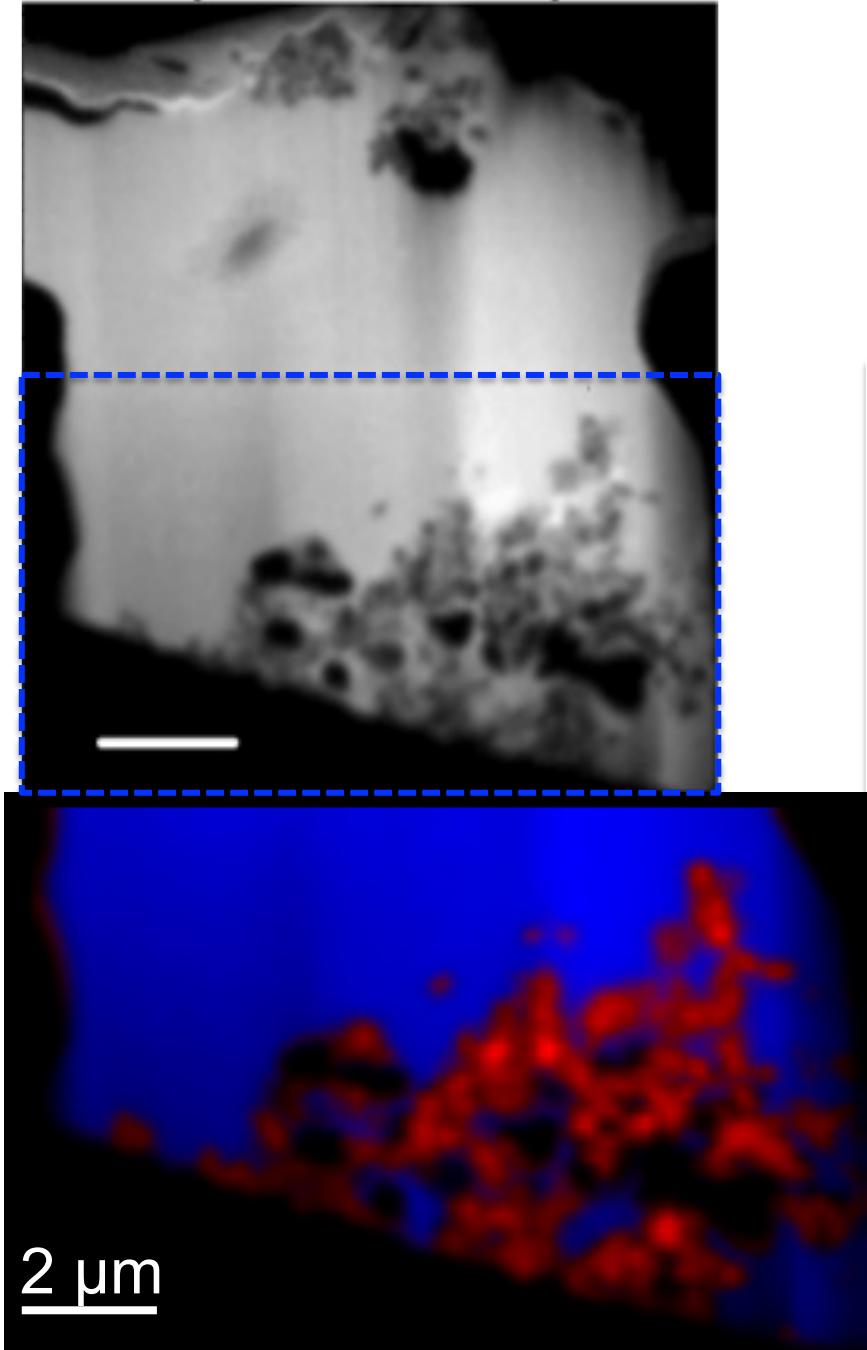
Dobrică+2012

100 nm FIB sections – no epoxy

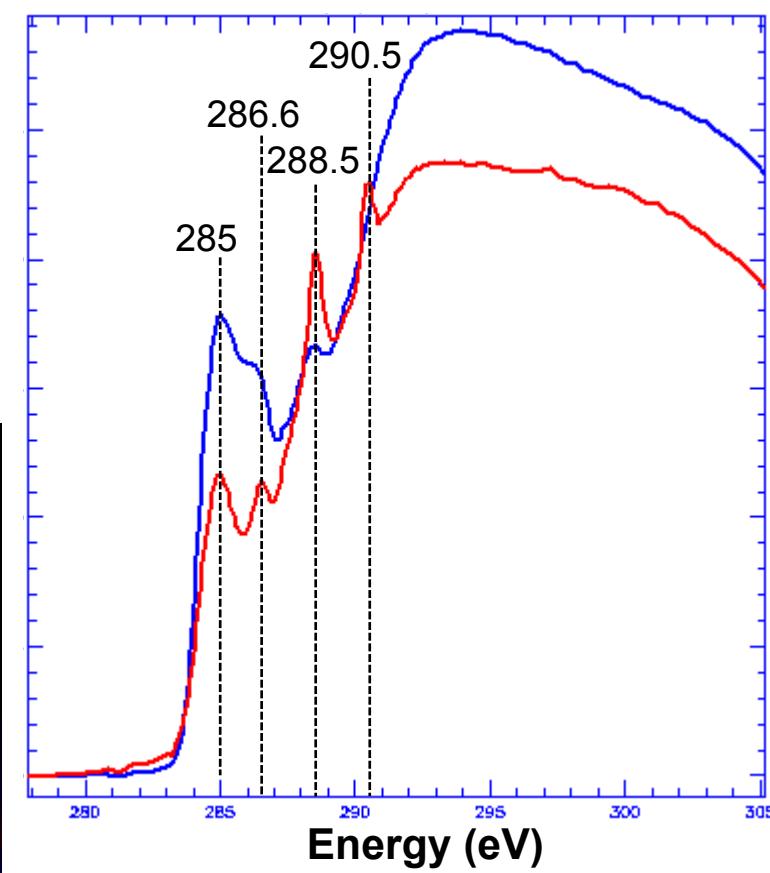
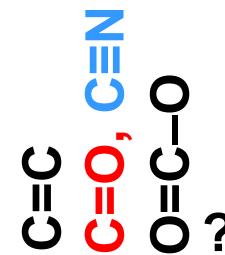


Engrand+2015 LPSC

(FIB Sections : D. Troadec IEMN Lille)

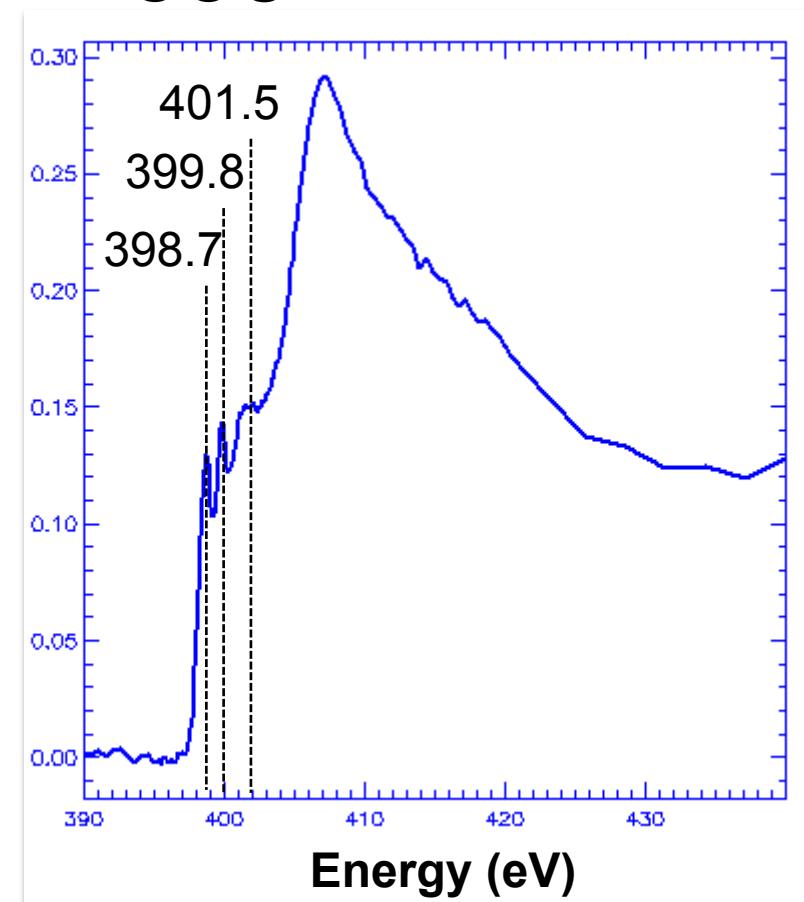
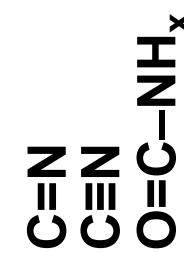
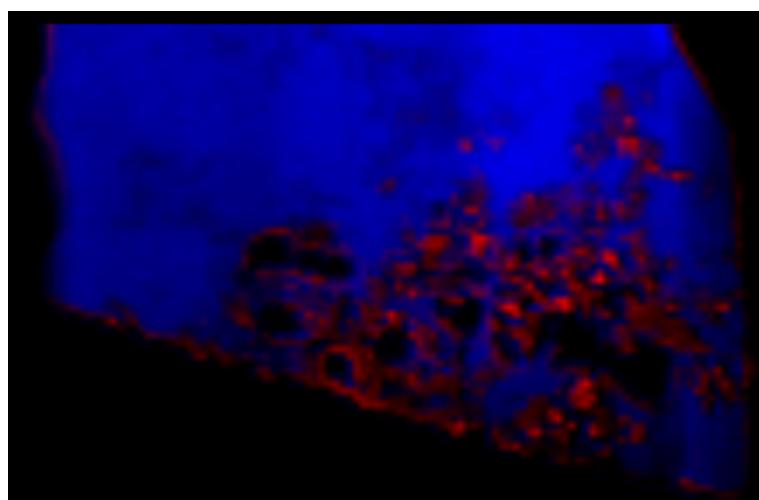
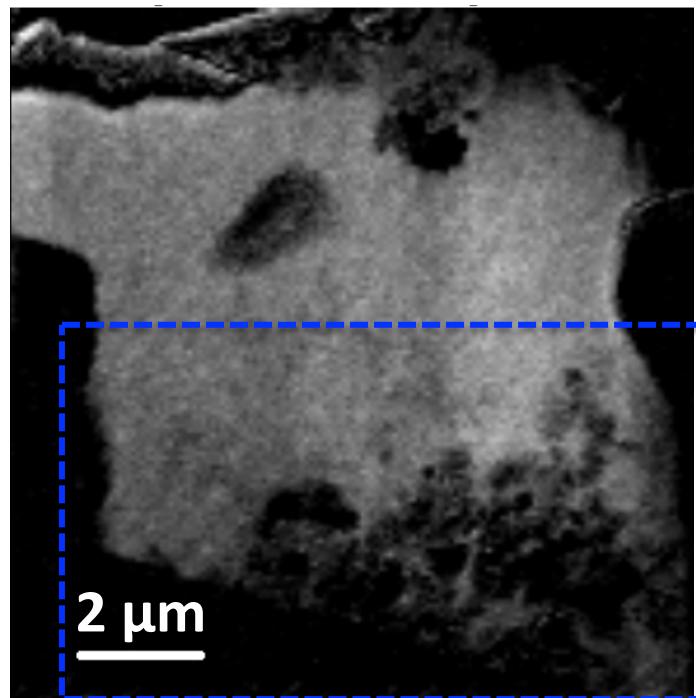


DC18
C-XANES

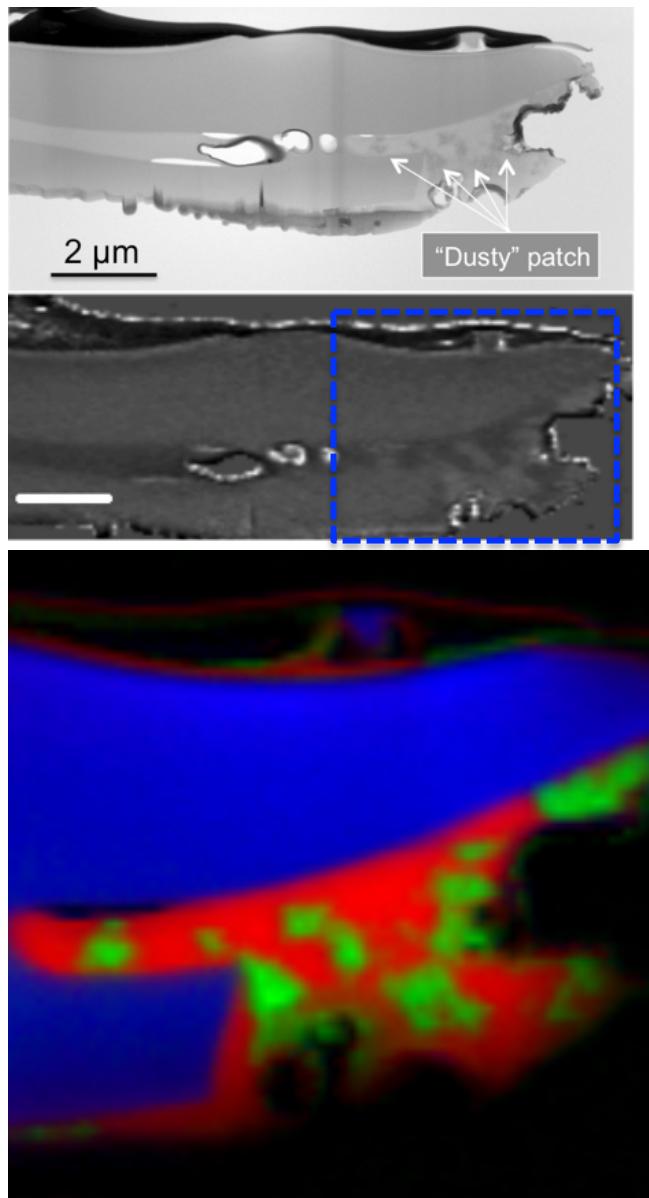


Engrand+2015 LPSC

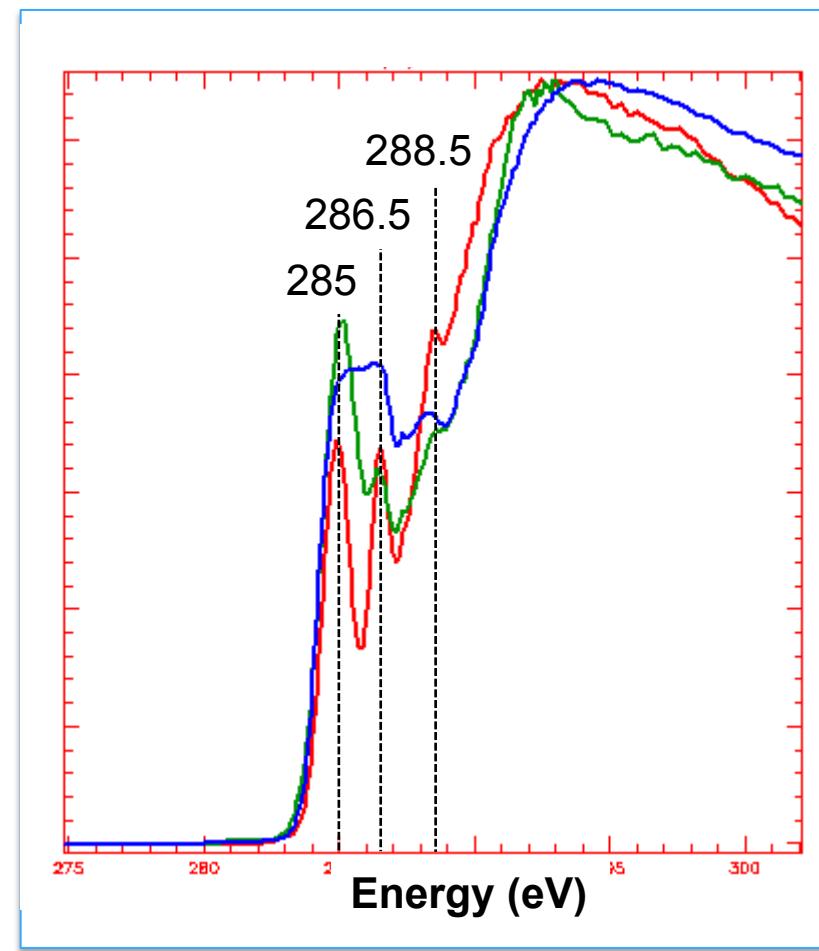
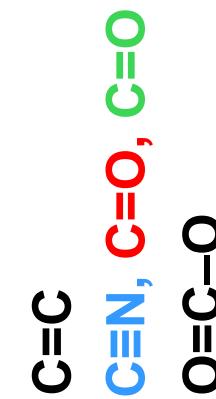
DC-18 N-XANES

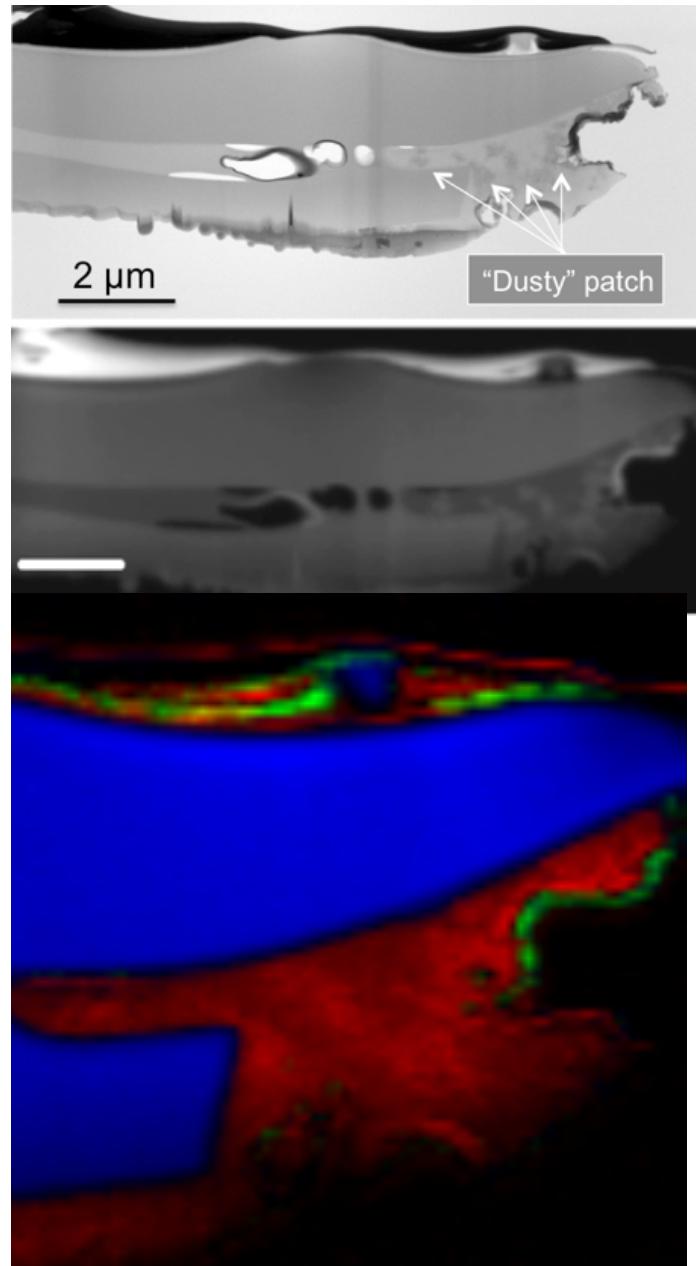


DC65 C-XANES



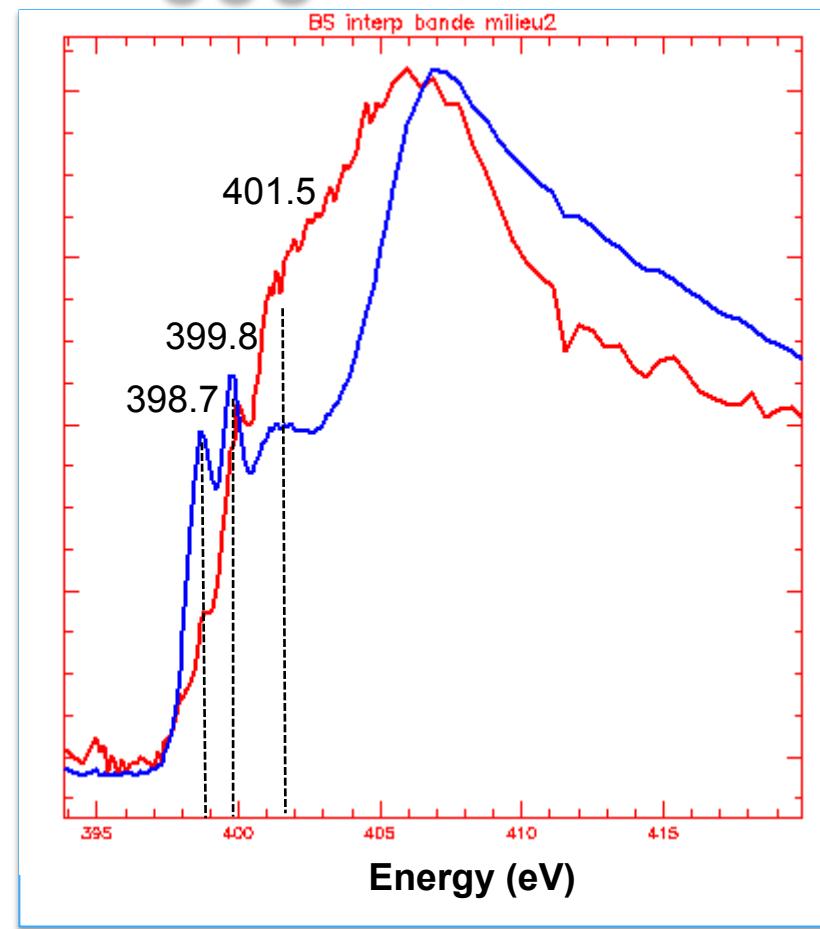
Engrand+2015 LPSC



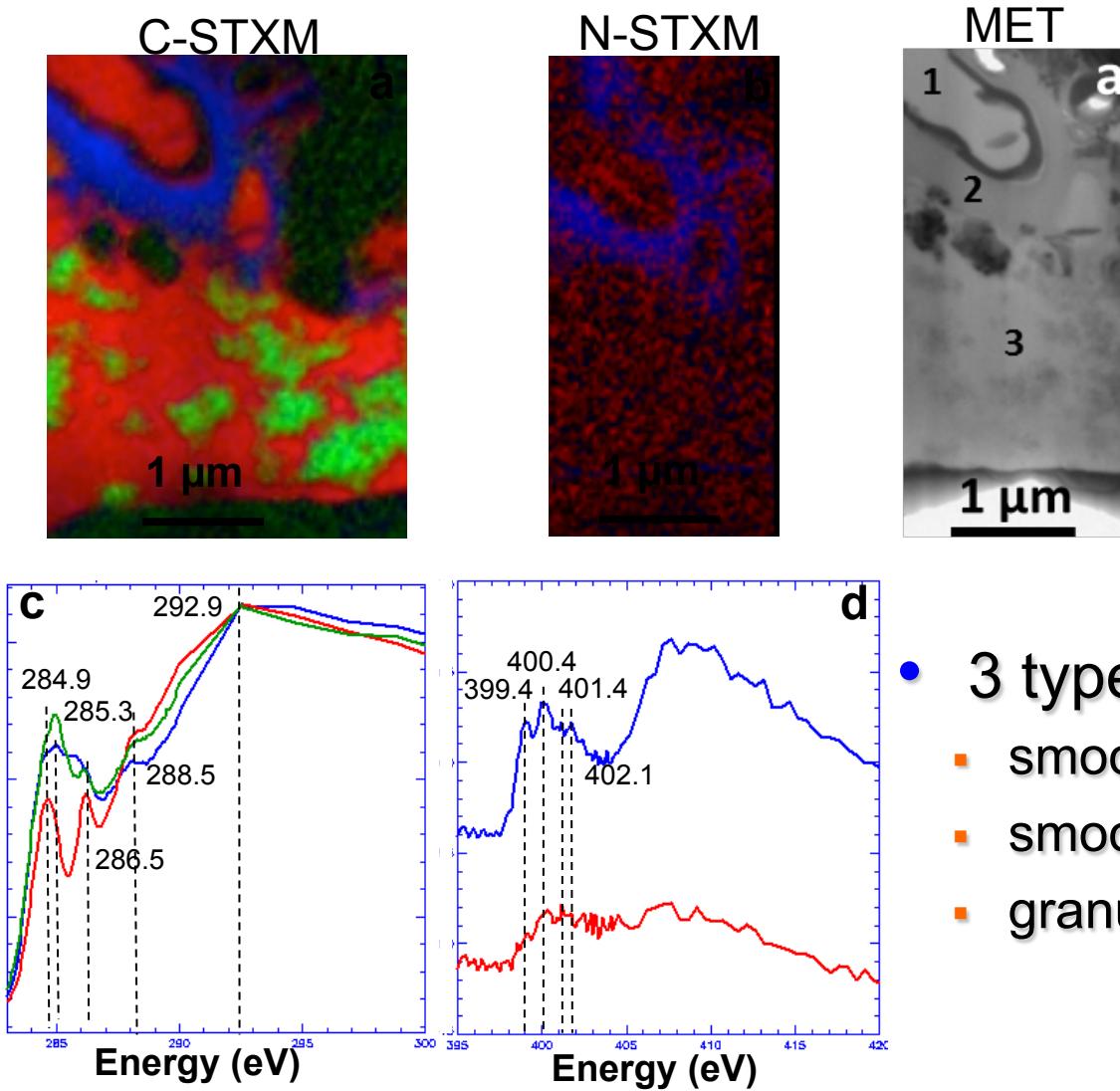


Engrand+2015 LPSC

DC65 N-XANES



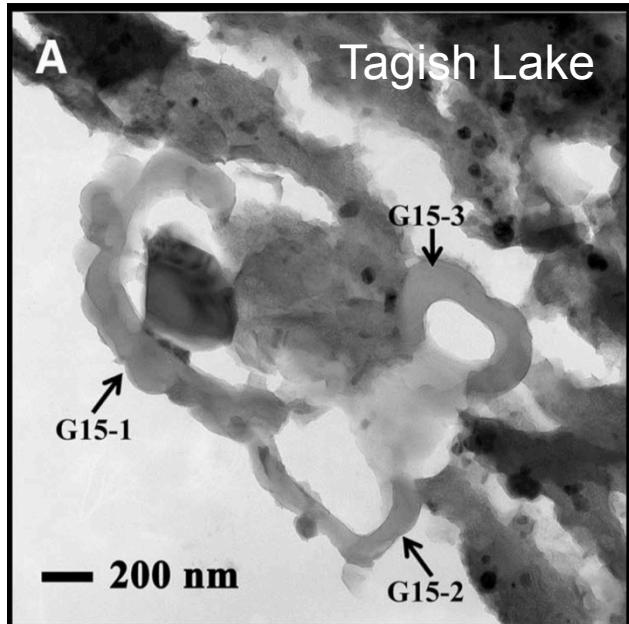
UCAMM DC43 : a unique sample



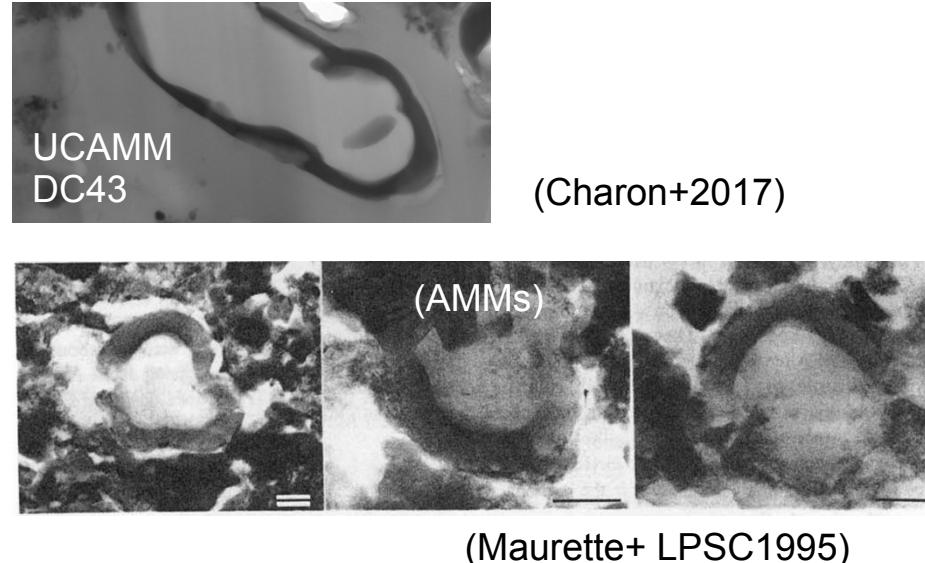
- 3 types of OM :
 - smooth : N-rich, mineral-poor (#2)
 - smooth: N-poor, mineral-poor (#1)
 - granular: N-poor, mineral-rich (#3)

(Charon+2017 LPSC)

Nanoglobules



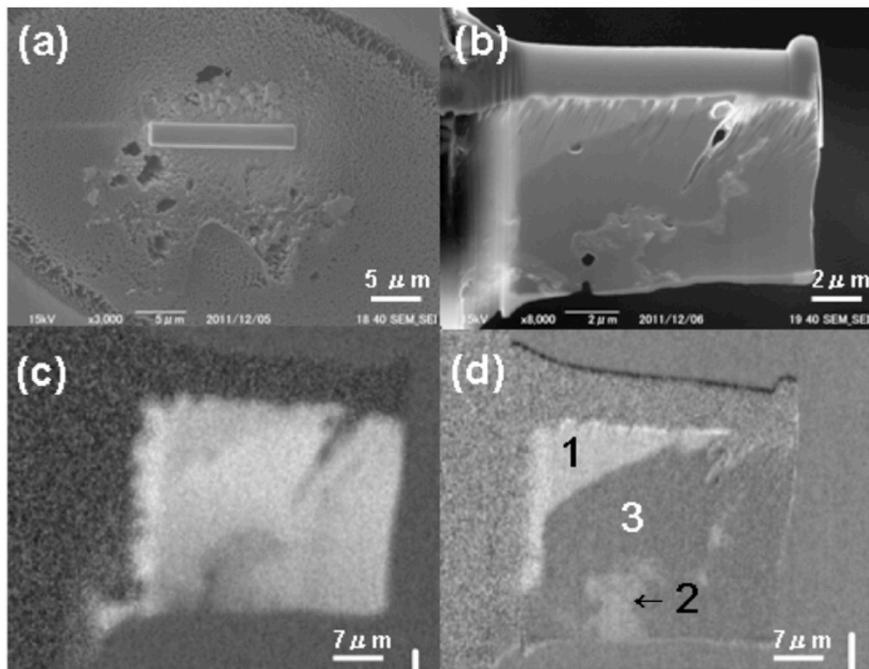
(Nakamura-Messenger+2003)



(Maurette+ LPSC1995)

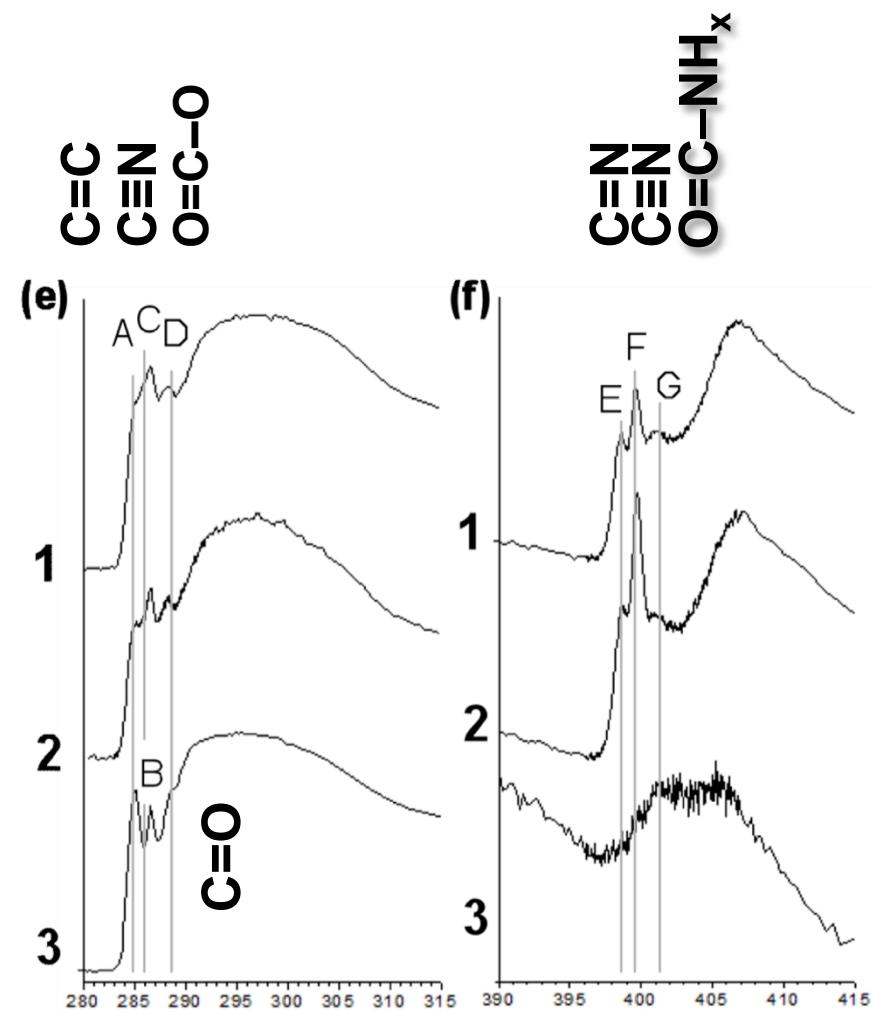
- D and N isotopic anomalies
 - δD from 1800 to 8100‰ in Tagish Lake
 - $\delta^{15}\text{N}$ from 200-1000‰ in Tagish Lake
- Present in many primitive materials
- Origin unclear : low T (outer regions or presolar cloud), processed ice coating?

UCMM (Japan)

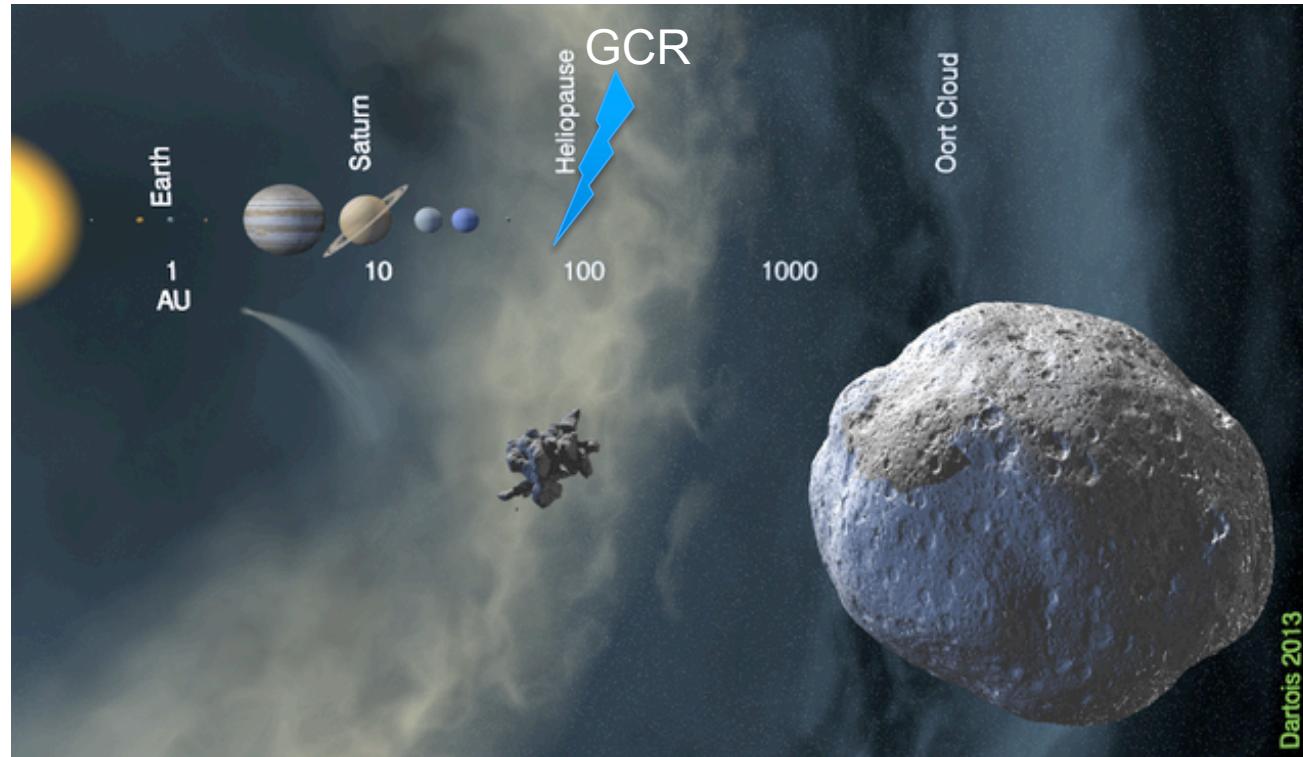


Yabuta+2012

Fig. 2. (a) Tungsten strap of 1.3 μm thickness deposited on the surface of UCMM for FIB extraction, (b) An SEM image of the FIB section, (c) Carbon- and (d) Nitrogen- distribution maps obtained by STXM, and (e) Carbon- and (f) Nitrogen-XANES spectra of the regions 1, 2, and 3 indicated in (d). Peak assignments are based on [4] ; A (285.1 eV): $\text{C}^*=\text{C}$, B (286.6 eV): $\text{C}=\text{C}-\text{C}^*=\text{O}$, C (286.6 eV): $\text{C}^*\equiv\text{N}$, D: $\text{NH}_x(\text{C}^*=\text{O})\text{C}$ (\sim 288.3 eV) or $\text{OR}(\text{C}^*=\text{O})\text{C}$ (\sim 288.6 eV), E (398.8 eV): $\text{C}=\text{N}^*$, F (399.7 eV): $\text{C}\equiv\text{N}^*$, G (\sim 401.5 eV): $\text{N}^*\text{H}_x(\text{C}=\text{O})\text{C}$.



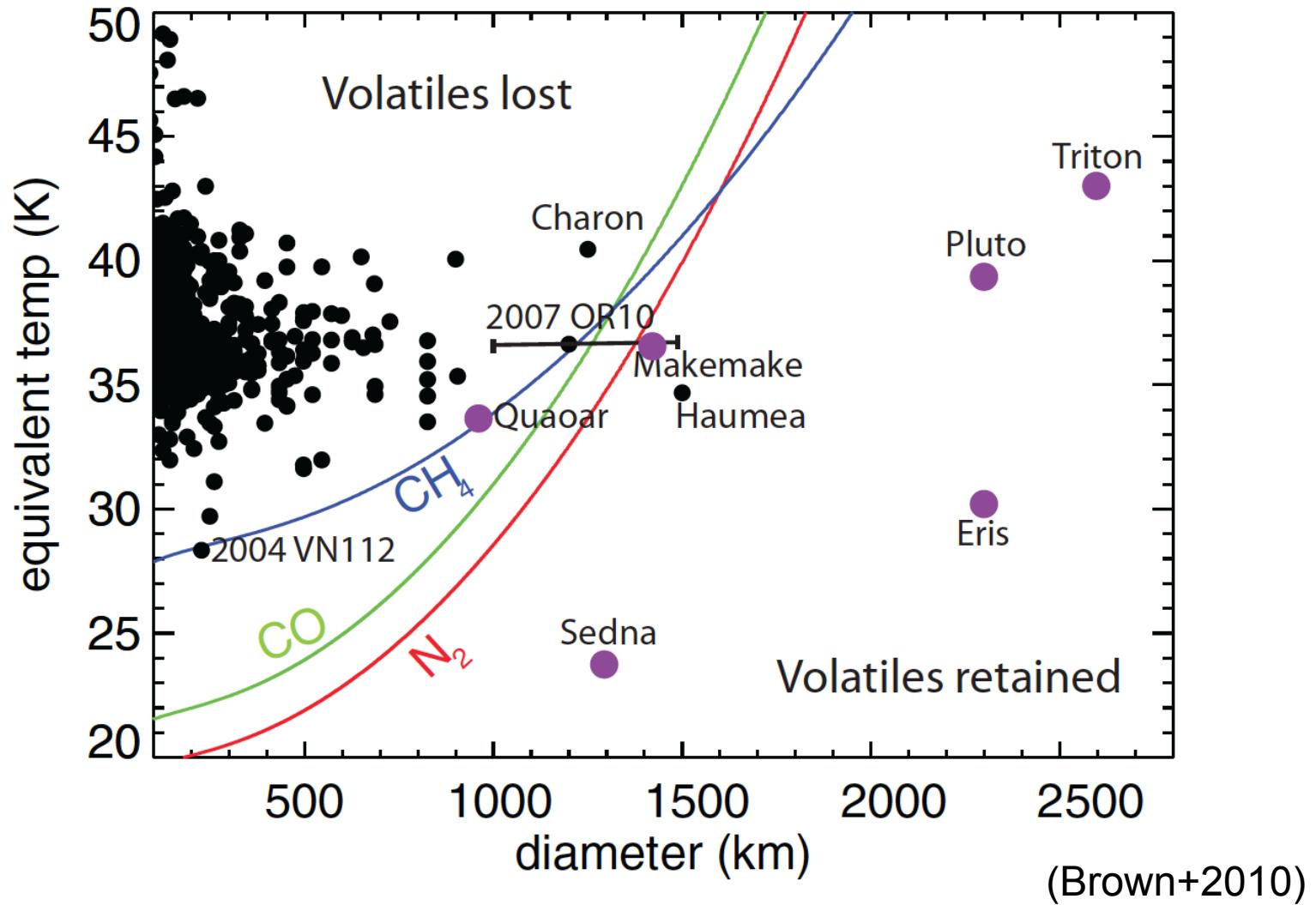
Formation of N-rich OM by irradiation?



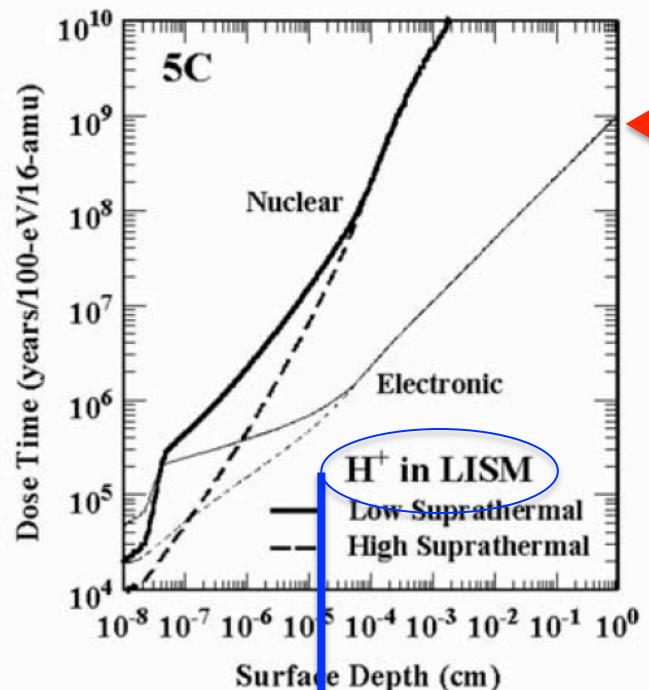
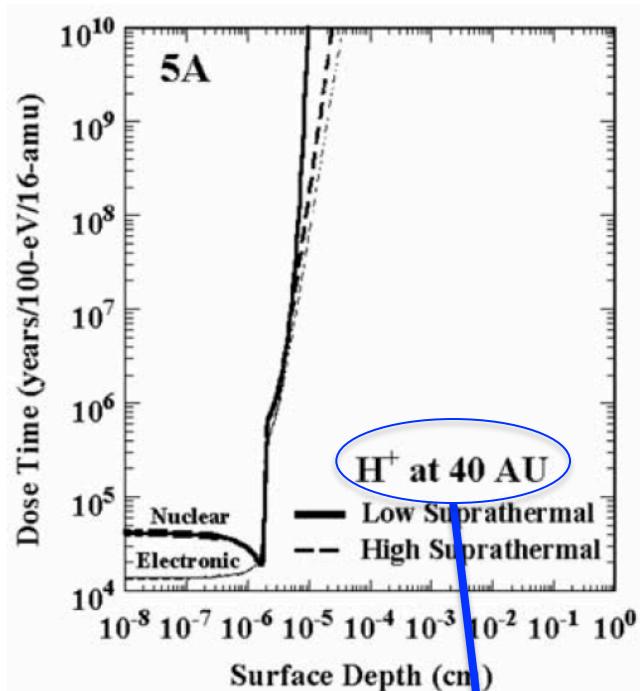
Dartois+2013

- UCAMMs : material from **beyond the nitrogen snow-line**
- **Formation of N-rich OM by GCR irradiation** at surface of a Kuiper belt or Oort cloud icy objet ?

Ices stability



Radiative environments

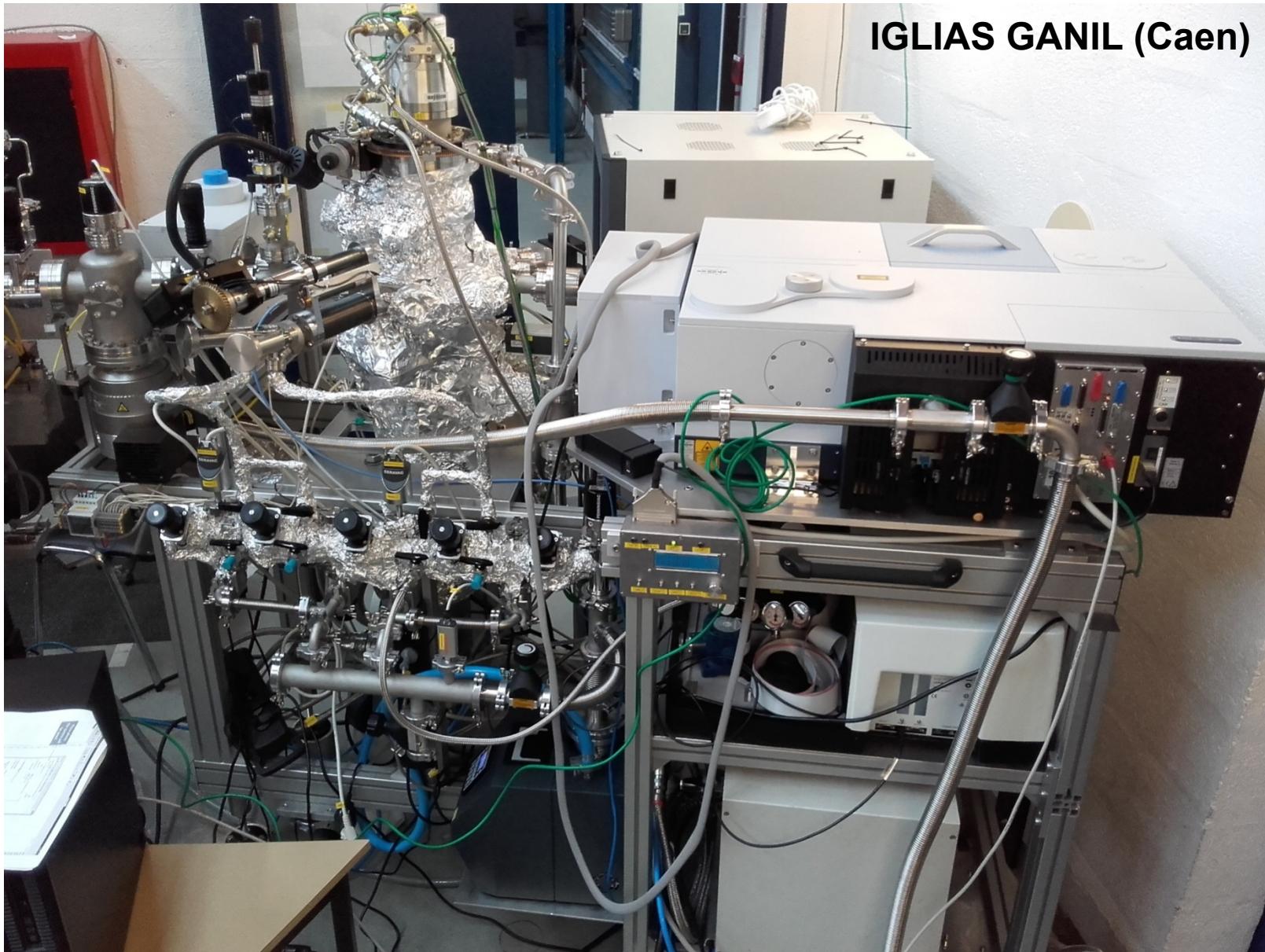


(Cooper+2003)





Centre de Recherche sur les Ions, les Matériaux et la Photonique



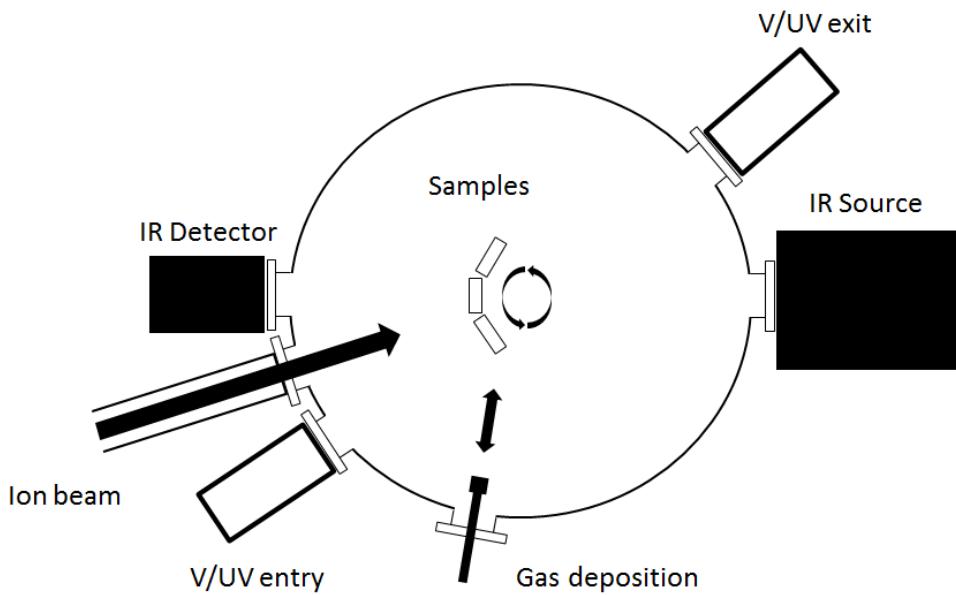
(Augé + 2016)





Irradiation de GLaces d'Intérêt Astrophysique (IGLIAS) – GANIL Caen

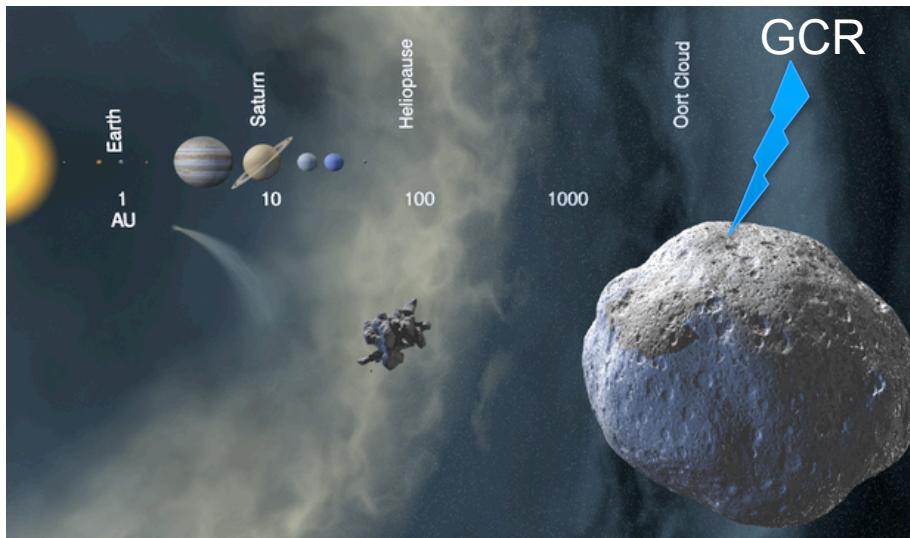
Centre de Recherche sur les Ions, les Matériaux et la Photonique



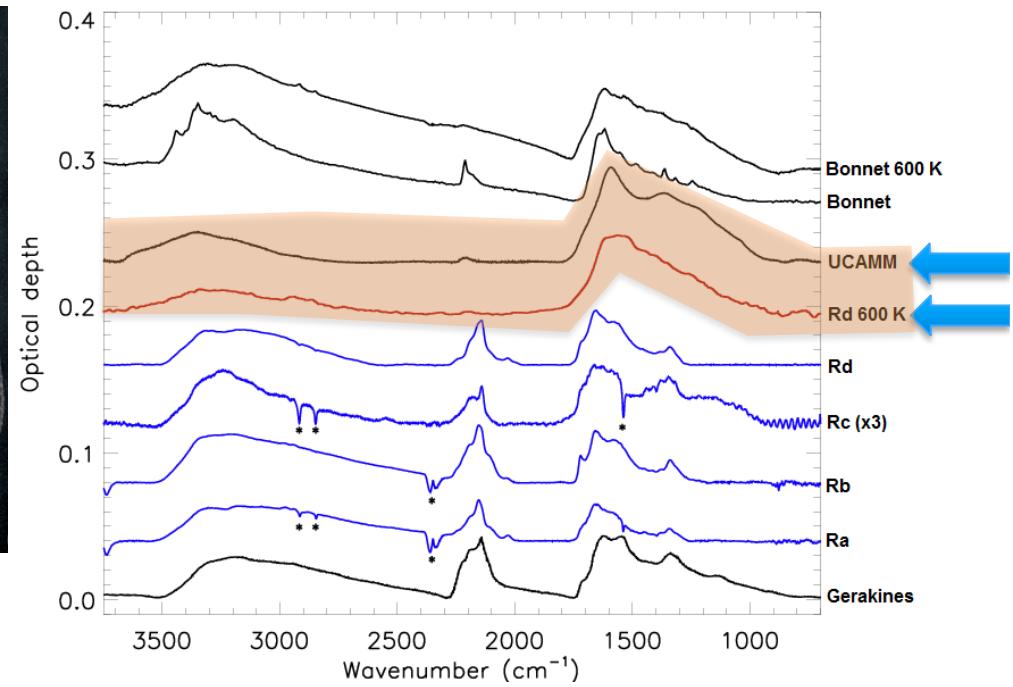
- P chamber : $1.5 \cdot 10^{-10}$ mbar @ 8 K
- Brucker Vertex 70v (6000-700 cm^{-1})
- Perkin Elmer λ 650 (800-200 nm)
- 3 substrates (CsI, SiO_2 , MgF_2 - 20*2 mm)

(Augé + 2016)

Formation of N-rich OM by irradiation?



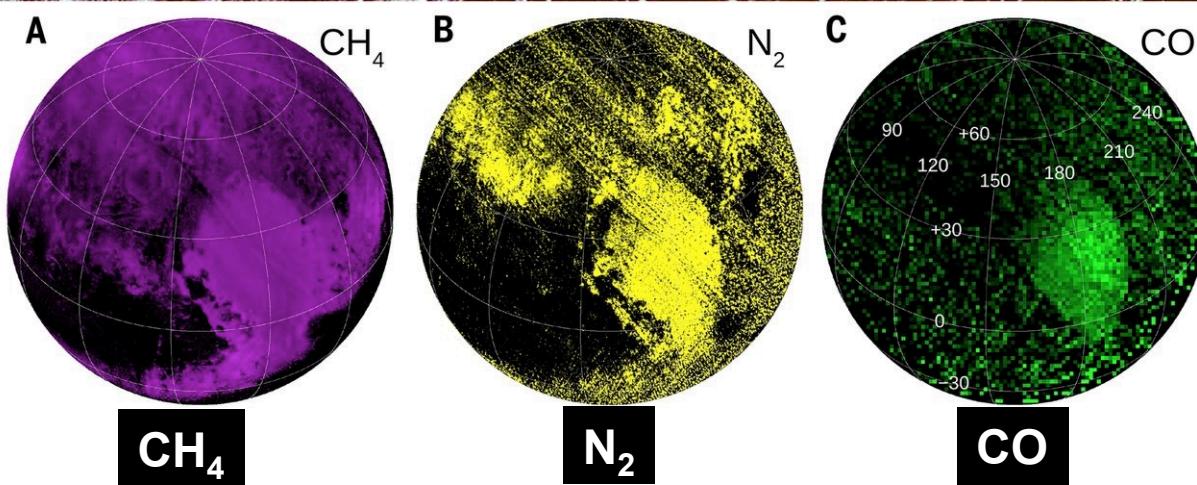
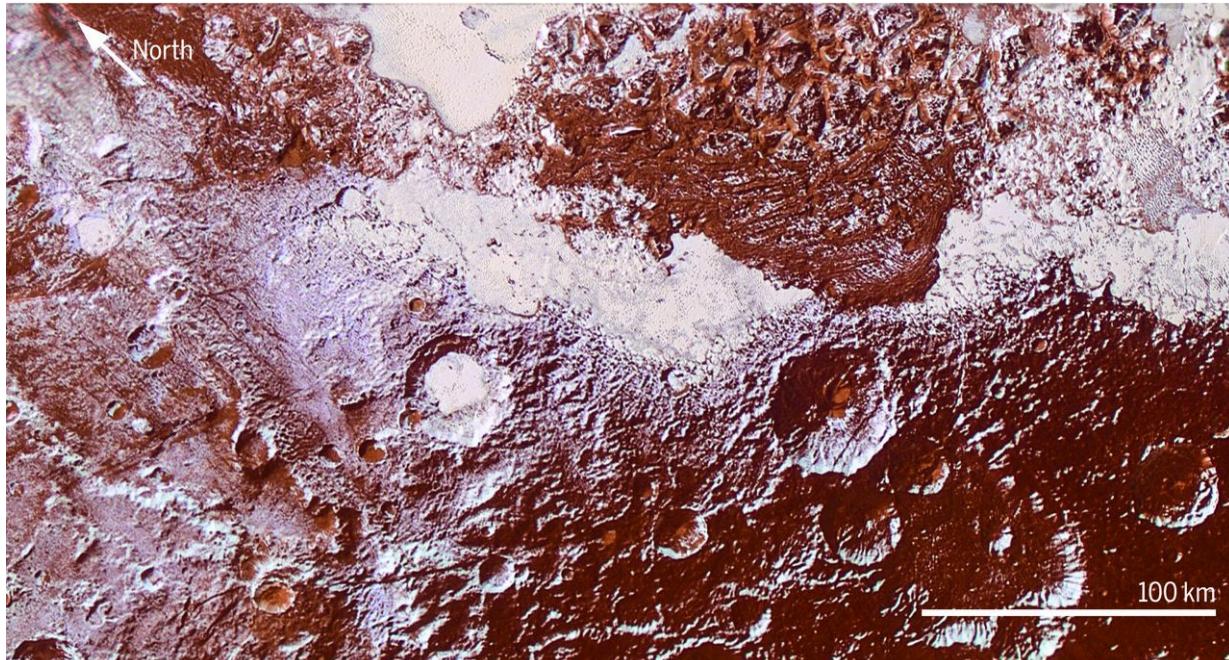
Dartois+2013



Augé+2016 A&A

- UCAMMs : material from **beyond the nitrogen snow-line**
- **Formation of N-rich OM by GCR irradiation at surface of a Kuiper belt or Oort cloud icy objet ?**

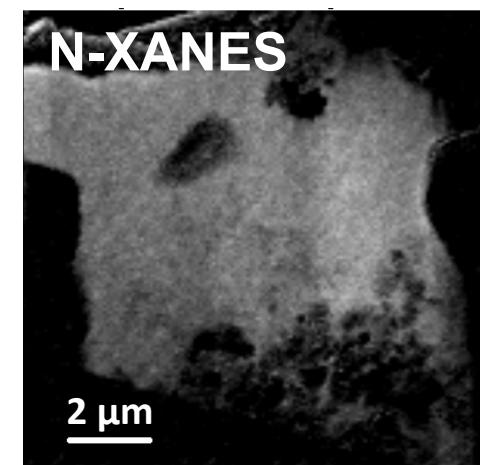
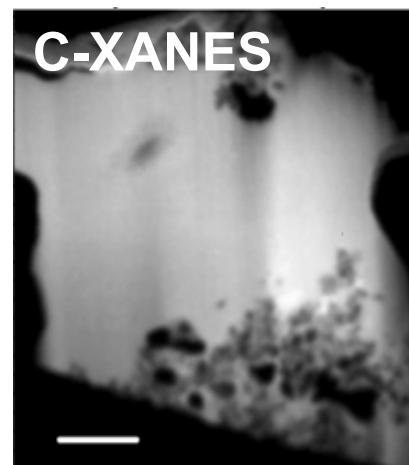
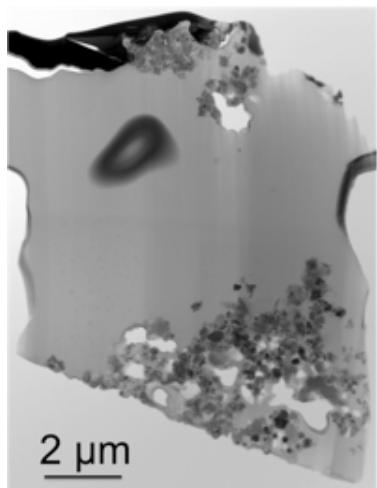
New Horizons – Pluto's Surface diversity



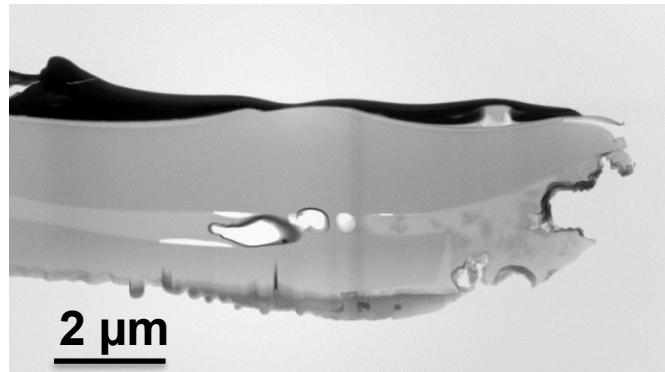
Grundy+2016

N-rich and N-poor OM in UCAMMs Intimate mixing with minerals

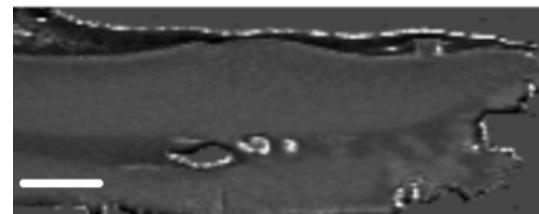
DC18 -
TEM



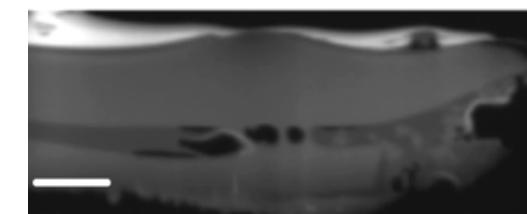
DC65 - TEM



C-XANES

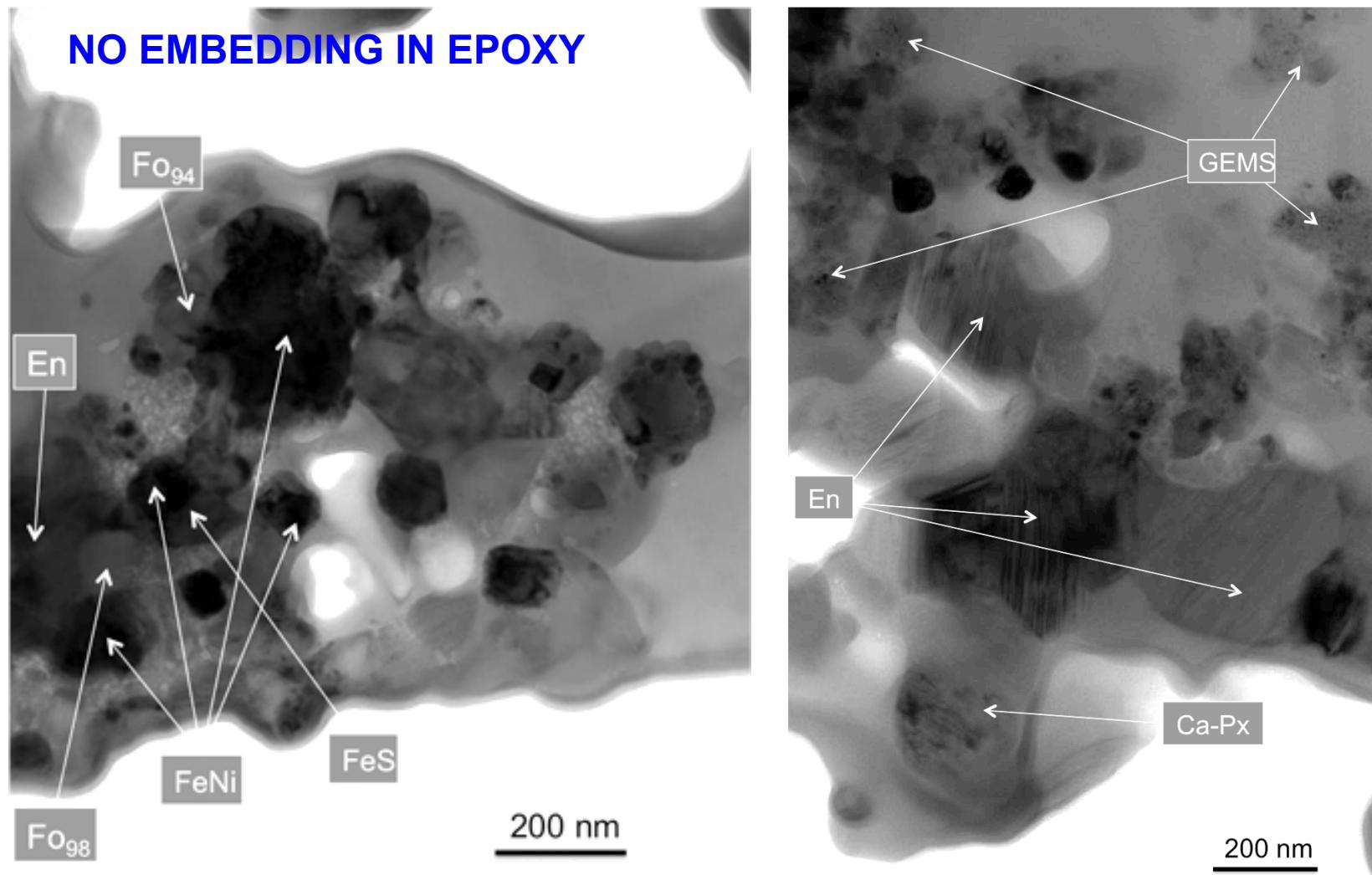


N-XANES



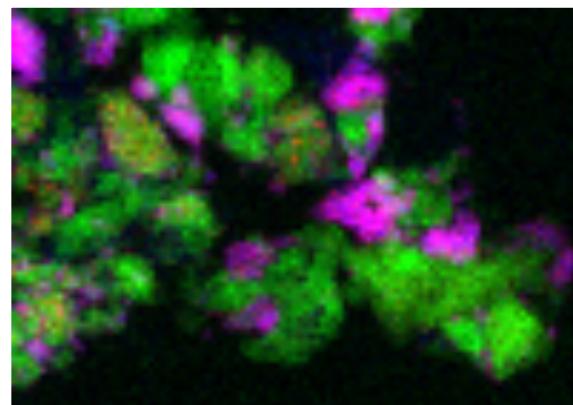
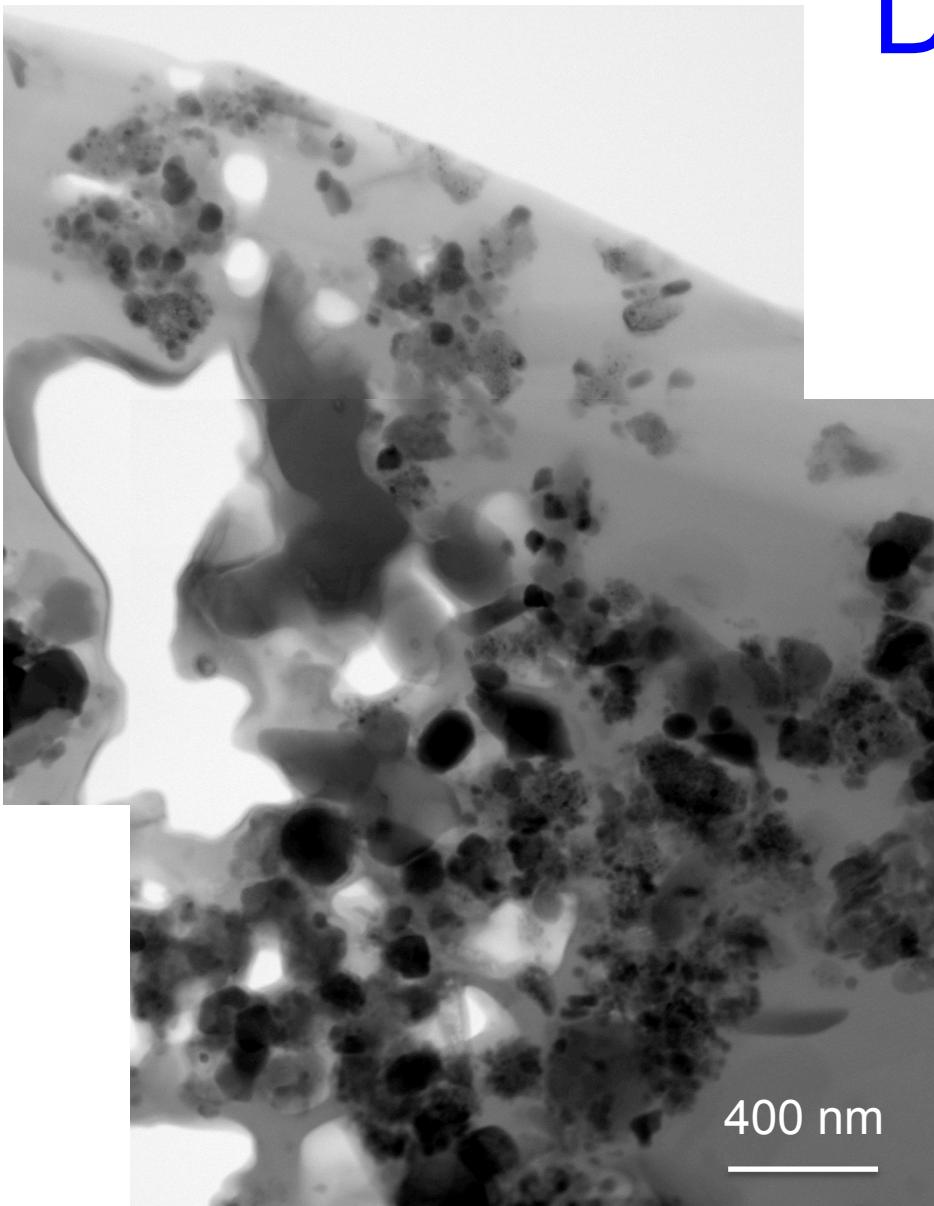
Engrand+ LPSC 2015
Yabuta+ LPSC 2012
Charon+ LPSC 2017

UCAMMs : Mineral complexity

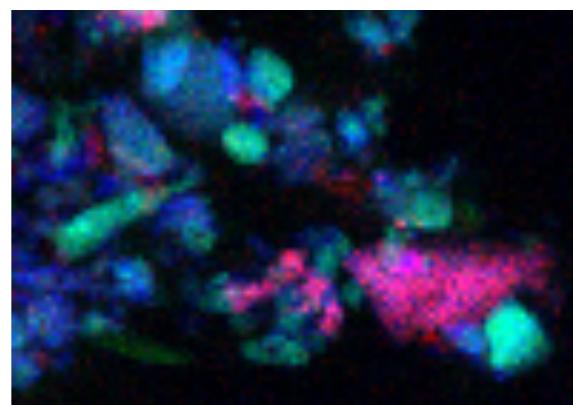


- Px/OI ratio (in numbers) ~ 1

DC18

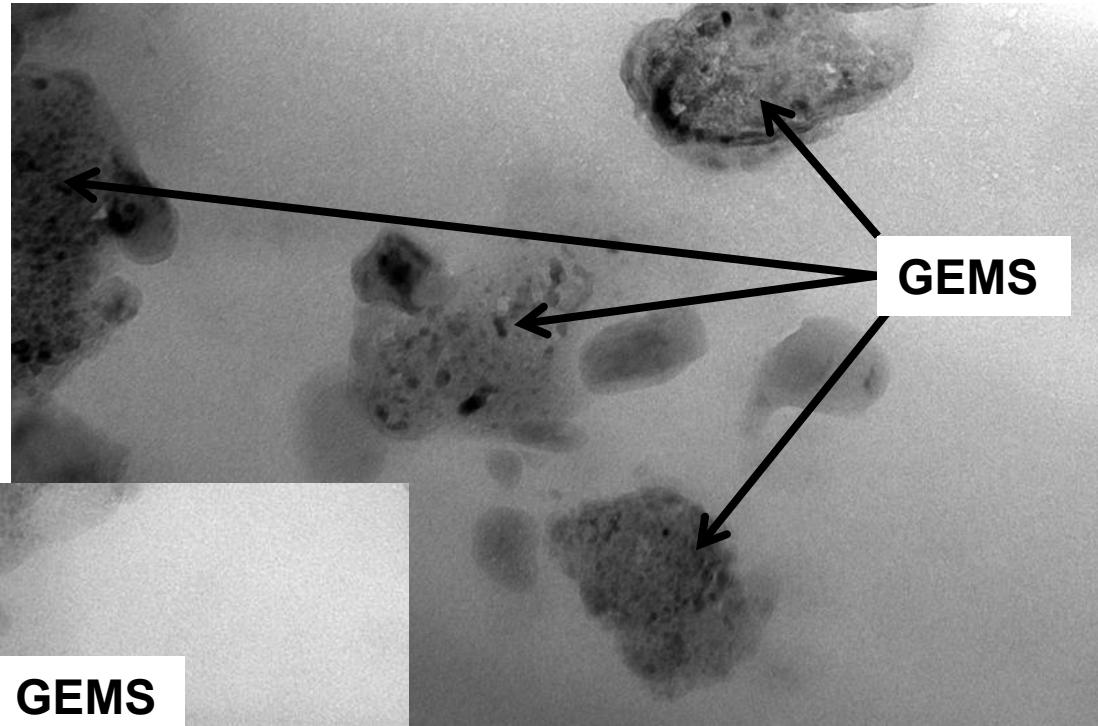
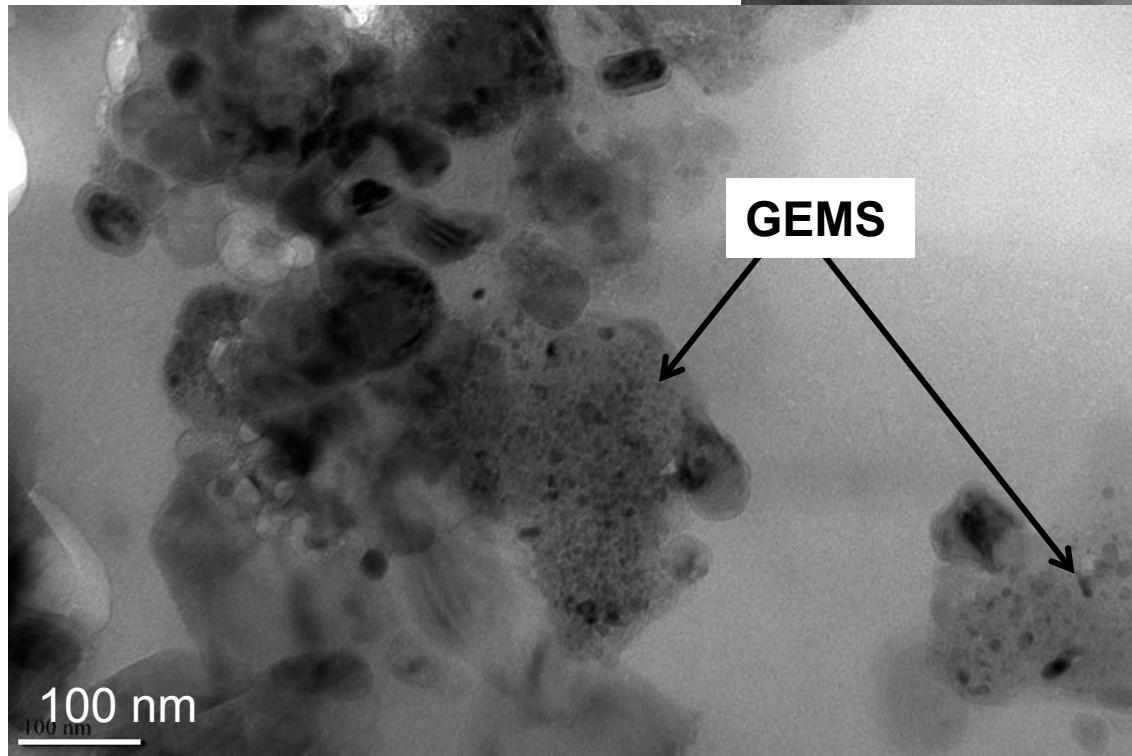


FeS
Silicate



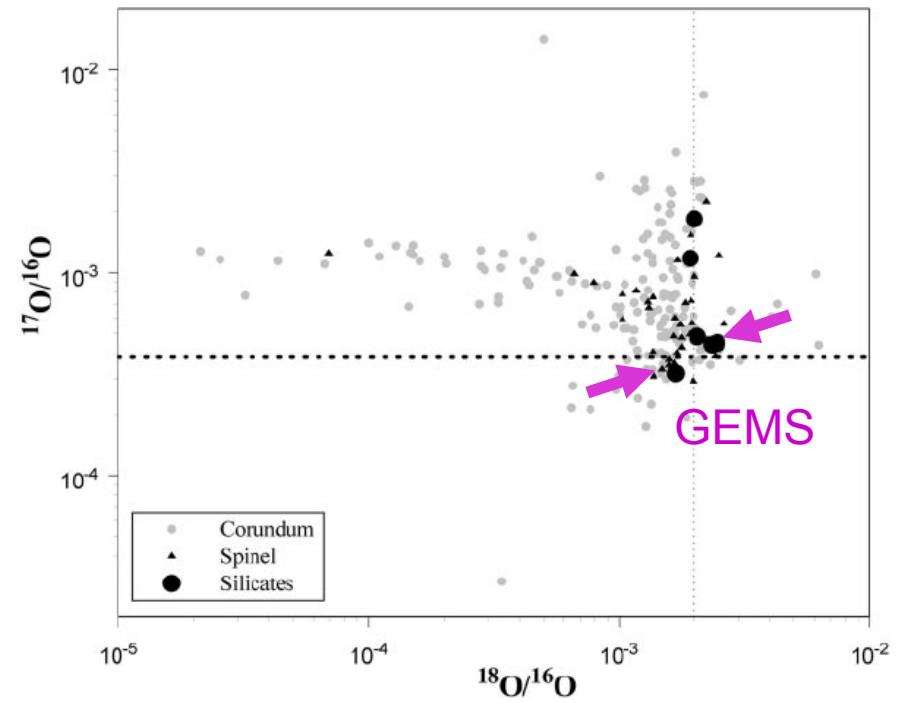
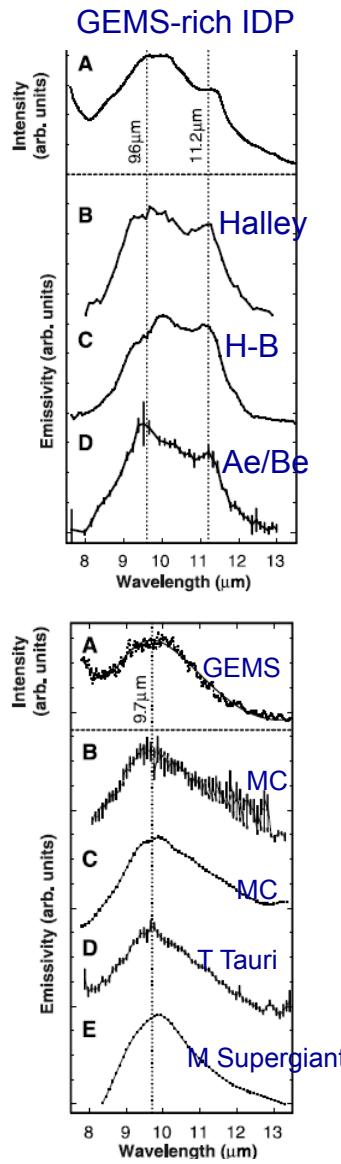
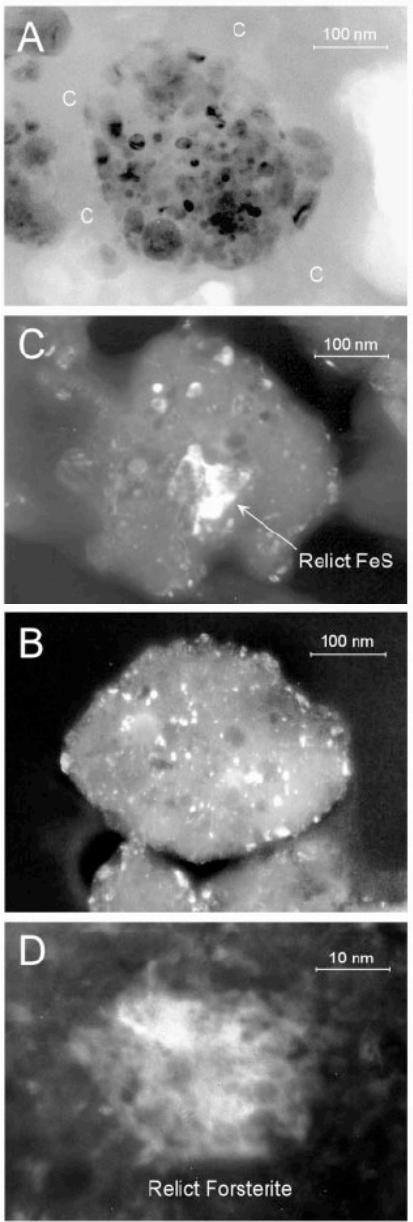
OI
Px
Ca-rich Px

Glassy phases : GEMS



Engrand+2015 LPSC

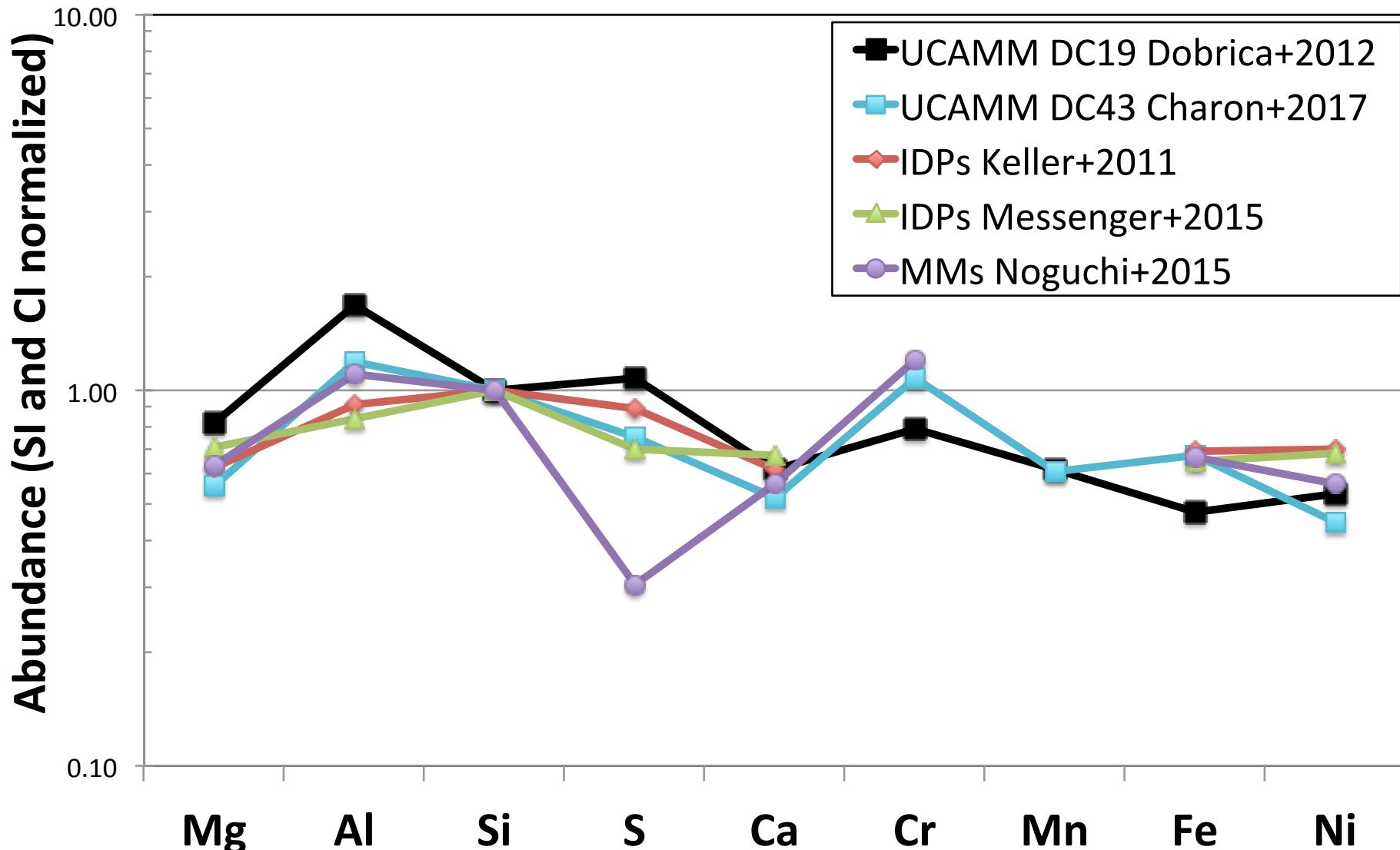
CP-IDPs contain abundant GEMS : an interstellar signature?



Messenger+2003

Bradley+1999

GEMS composition in IDPs/(UC)AMMs



Formation of GEMS?

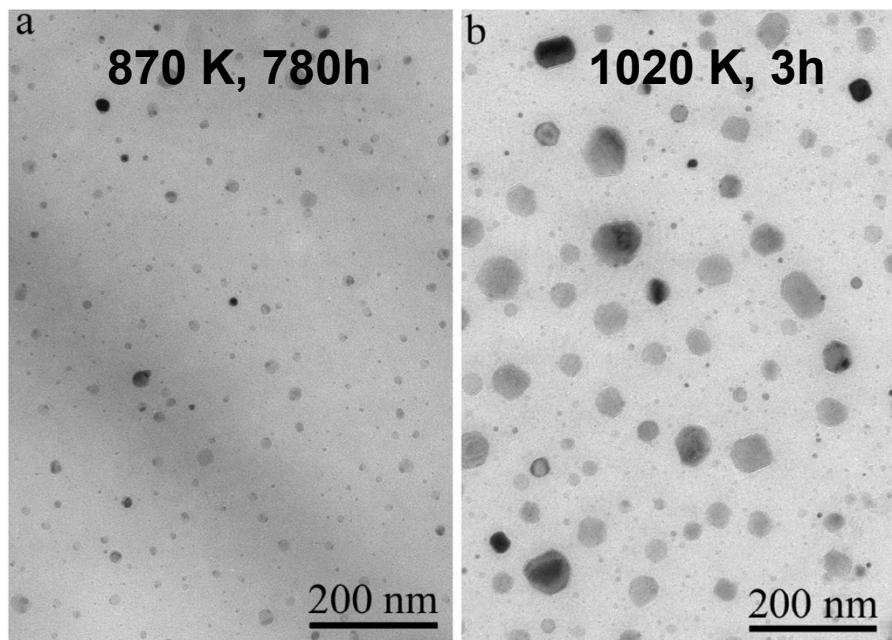


Fig. 1. TEM micrograph of annealed sample **a**) at 870 K for 780 h and **b**) at 1020 K for 3 h. Rounded metallic nano-particles enclosed in the amorphous silicate. They formed by a reduction reaction and further precipitation since metallic Fe is immiscible in silicates. The microstructure closely resembles to those to GEMS found in IDPs.

- annealing (<1000K) of amorphous olivine under vaccum?

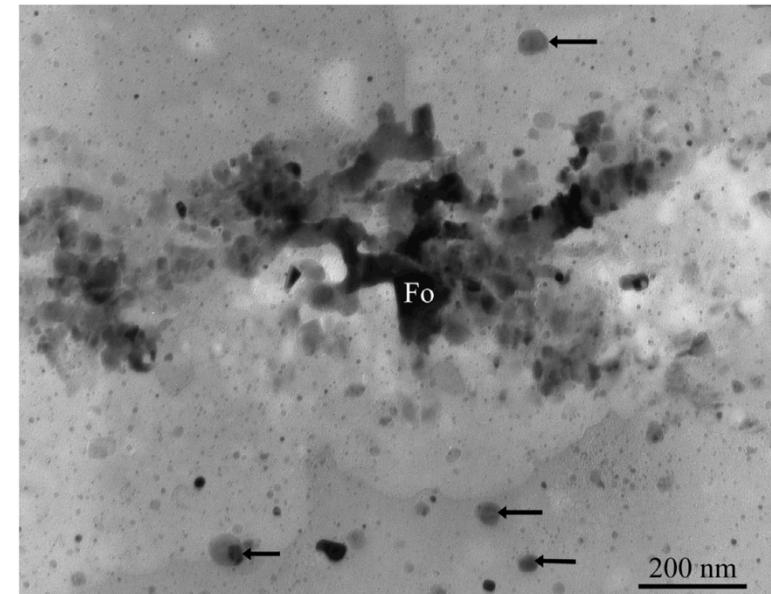
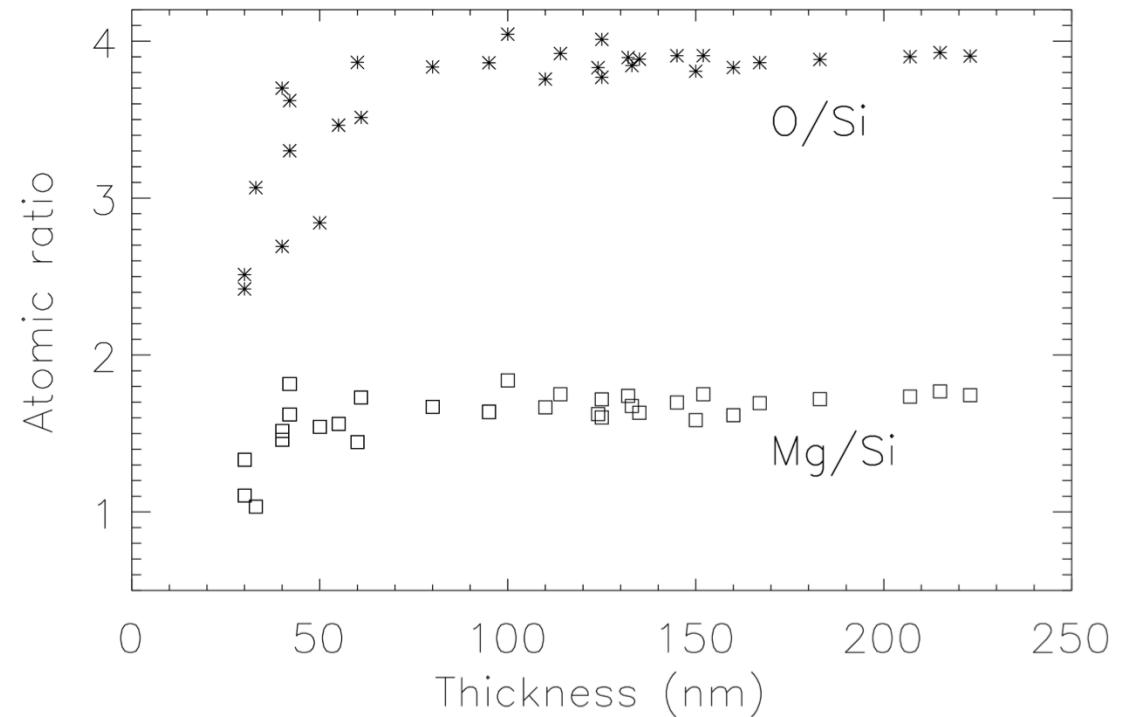
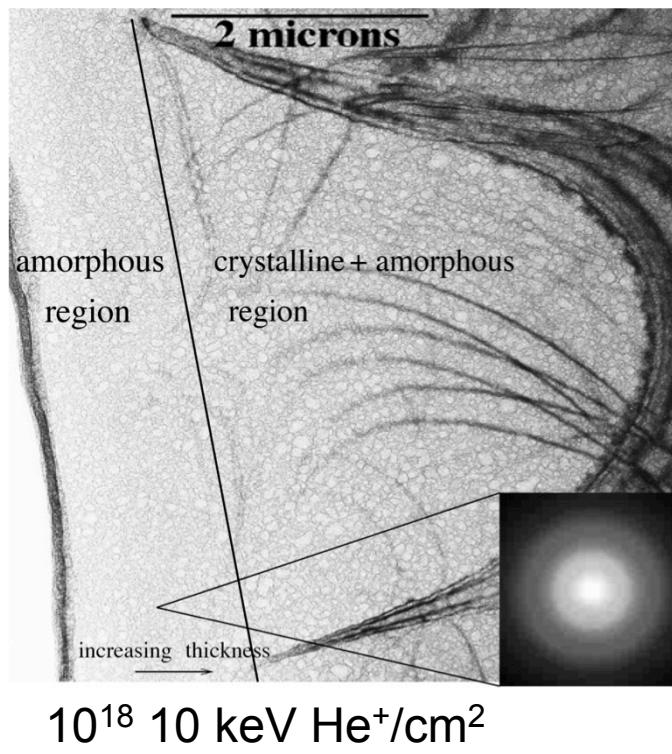


Fig. 2. TEM micrograph of sample annealed at 970 K (55 h) showing a forsterite crystal (Fo) embedded in a amorphous matrix. Note the dendritic structure at the edge of the grains. Some metal particles are also present in the amorphous phase (some of them are arrowed).

Davoisne+2006

Formation of GEMS? Low energy irradiation of olivine

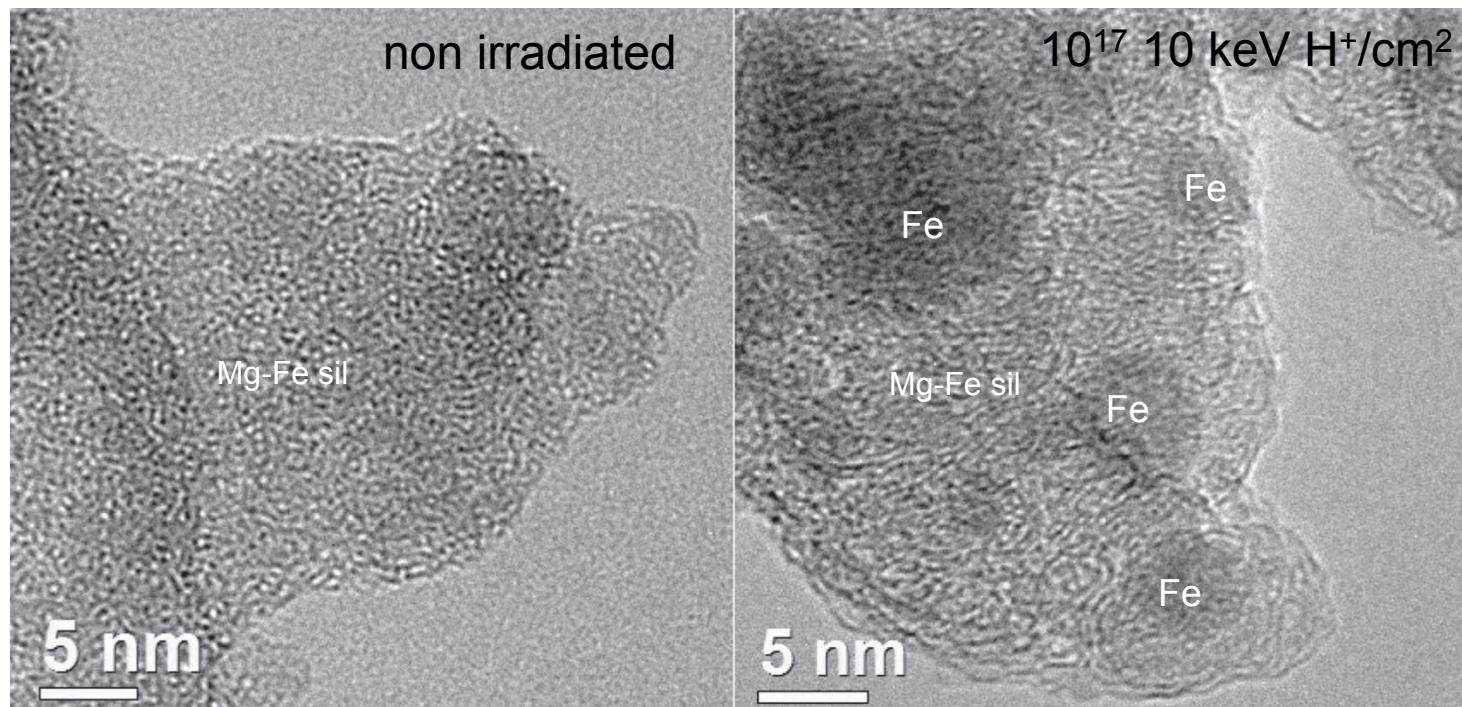


Demyk+2001
Davoisne+2008

- Disordered rim (40-90 nm)
- Change of composition (O, Mg)

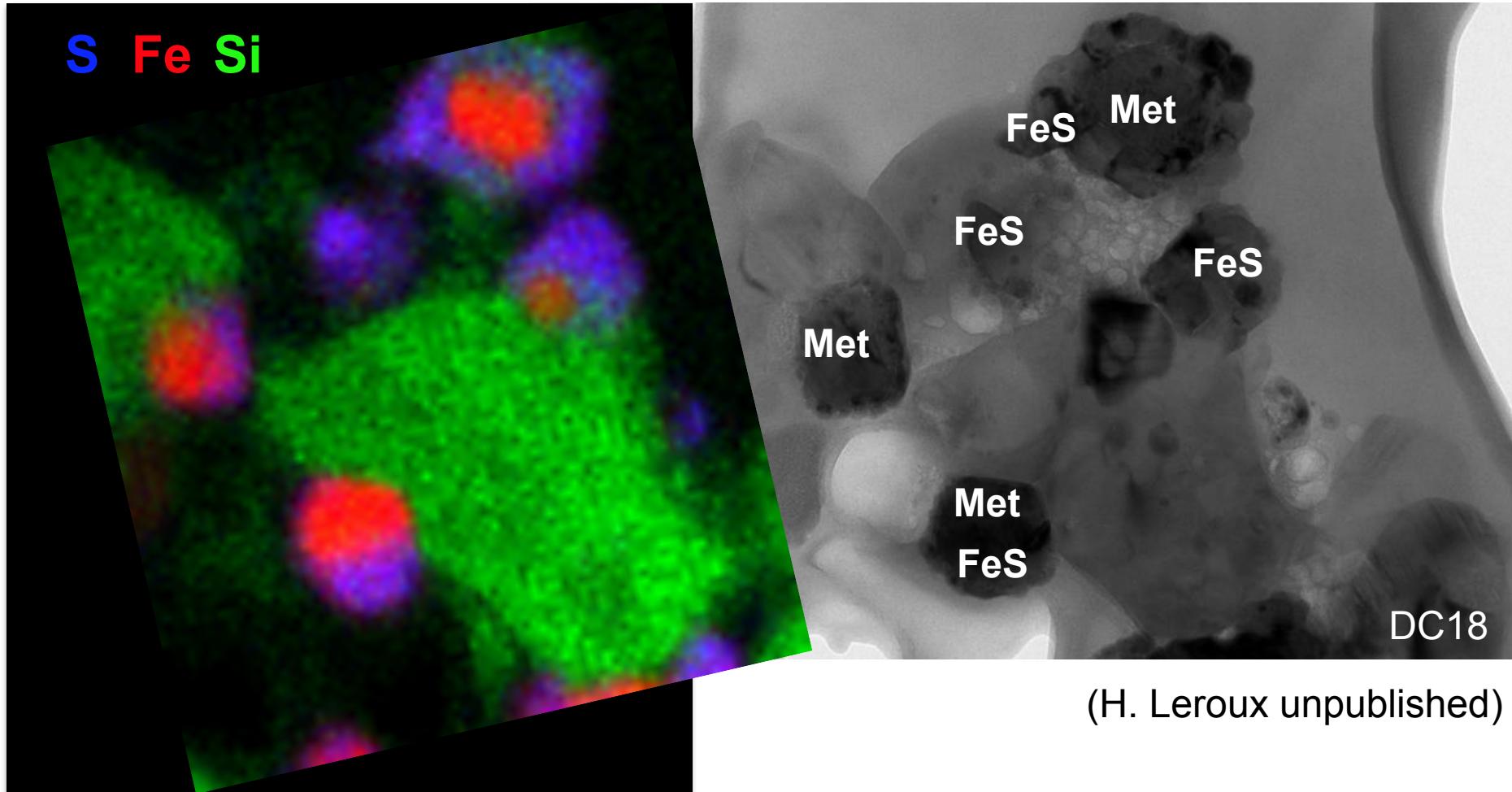
Formation of GEMS? Irradiation of amorphous silicates?

- Amorphous nanometer-sized Mg-Fe and Mg silicate of olivine-type stoichiometry (LA of metal targets (Mg-Fe-Si) in He/O₂ or of Mg-Fe silicates in 4 mbar He atmosphere)



Jäger+2016

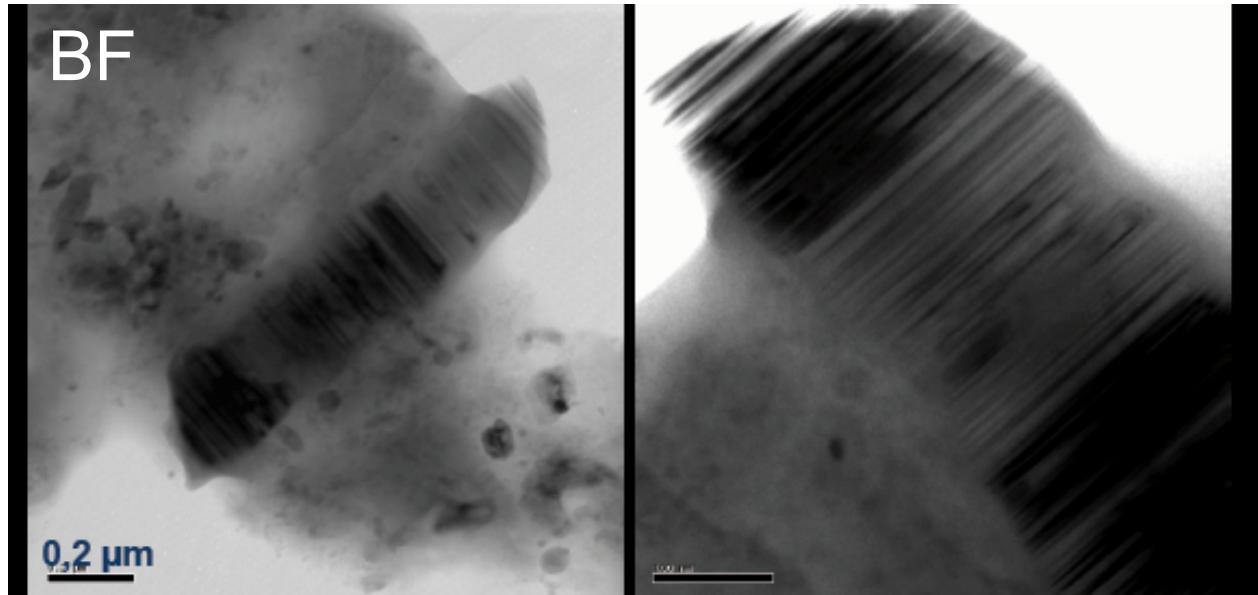
Metal – Fe sulfides



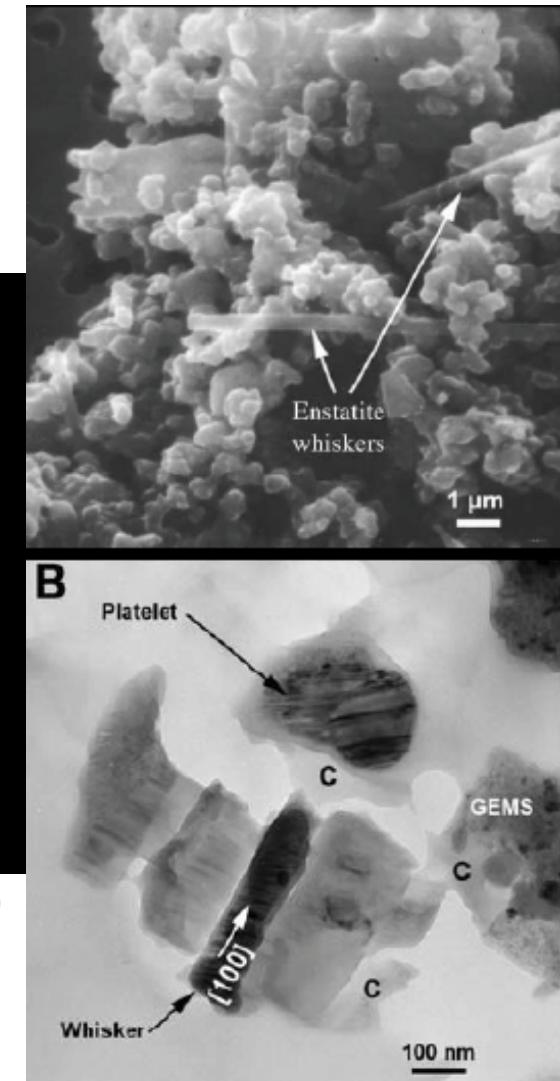
Mineral condensates : enstatite whiskers

IDP

AMM

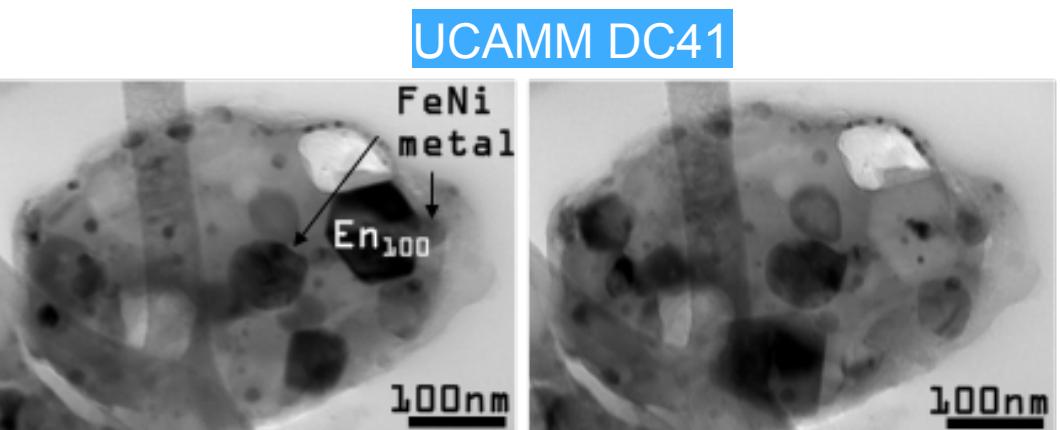
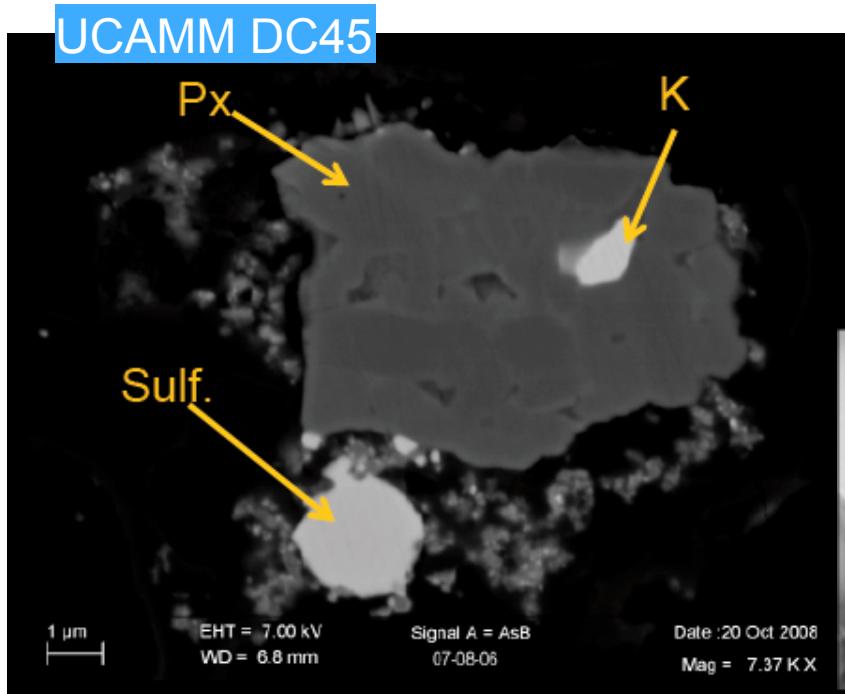


Dobrica+2009

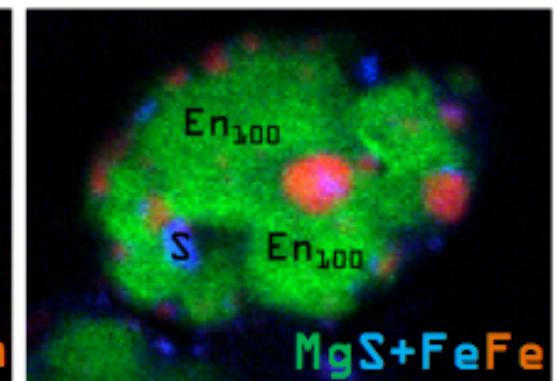
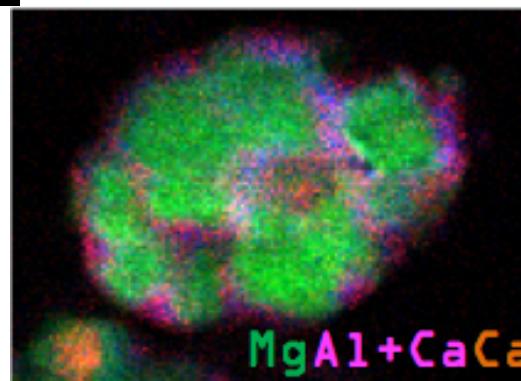


Bradley+1983

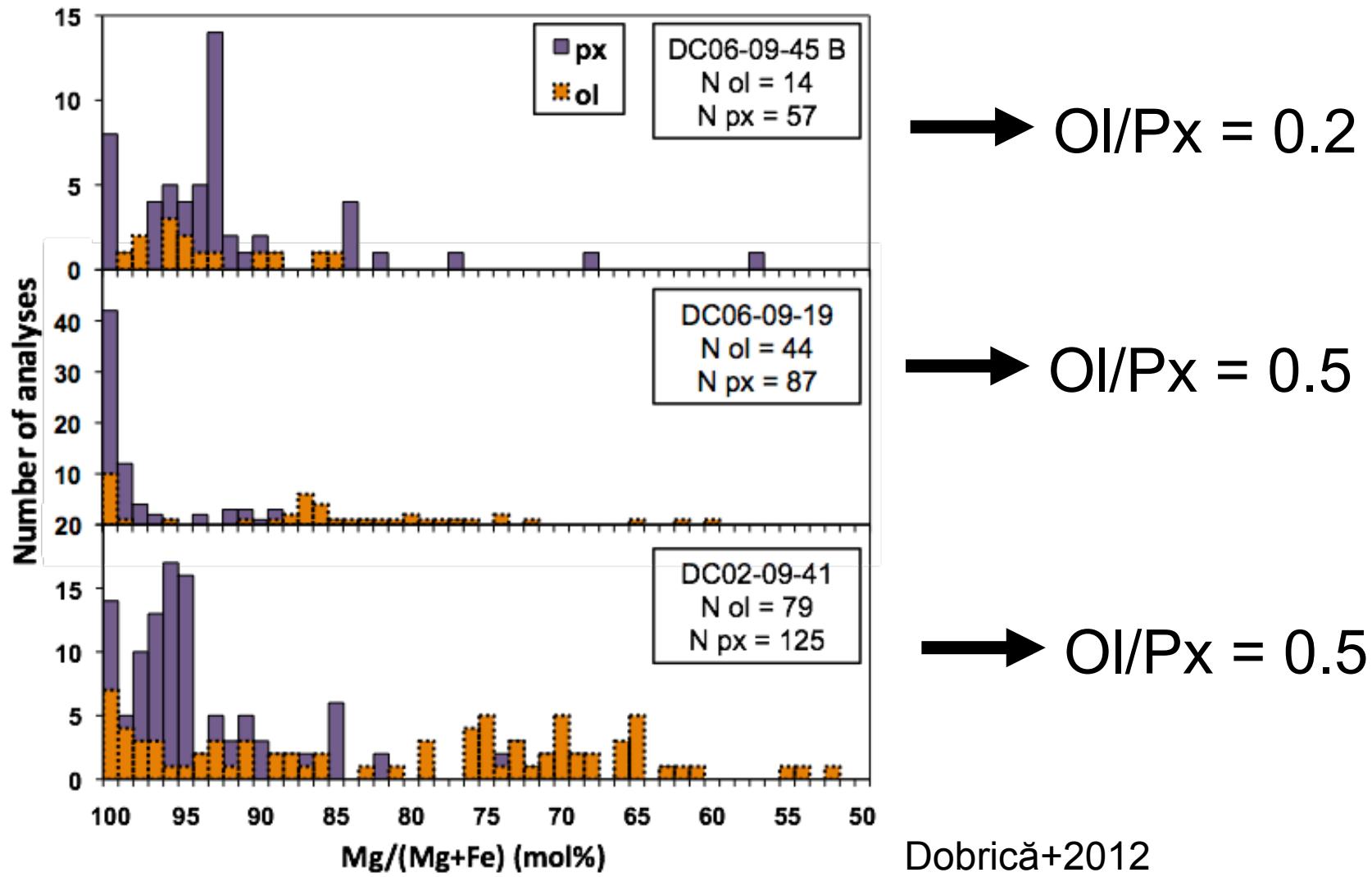
Chondrules – « Nanochondrules »



Dobrica+2012 GCA



UCAMM Olivines & Pyroxenes



Dobrică+2012

- Crystallinity > 25%

Minerals: ISM vs protoplanetary disks

- Circumstellar environments:
 - Minerals crystalline, very small
- Diffuse ISM :
 - > 98% minerals amorphous
 - Silicates ~ 50% OI - 50% Px
 - Size < 100 nm
- Disks :
 - 10-95% crystalline minerals
 - Sizes 0.1 – 2 μm
 - Silicates ~ 50% OI – 50% Px
 - Mineralogical zonation

A controversial mineralogical zoning in PPD

Herbig Ae/Ab

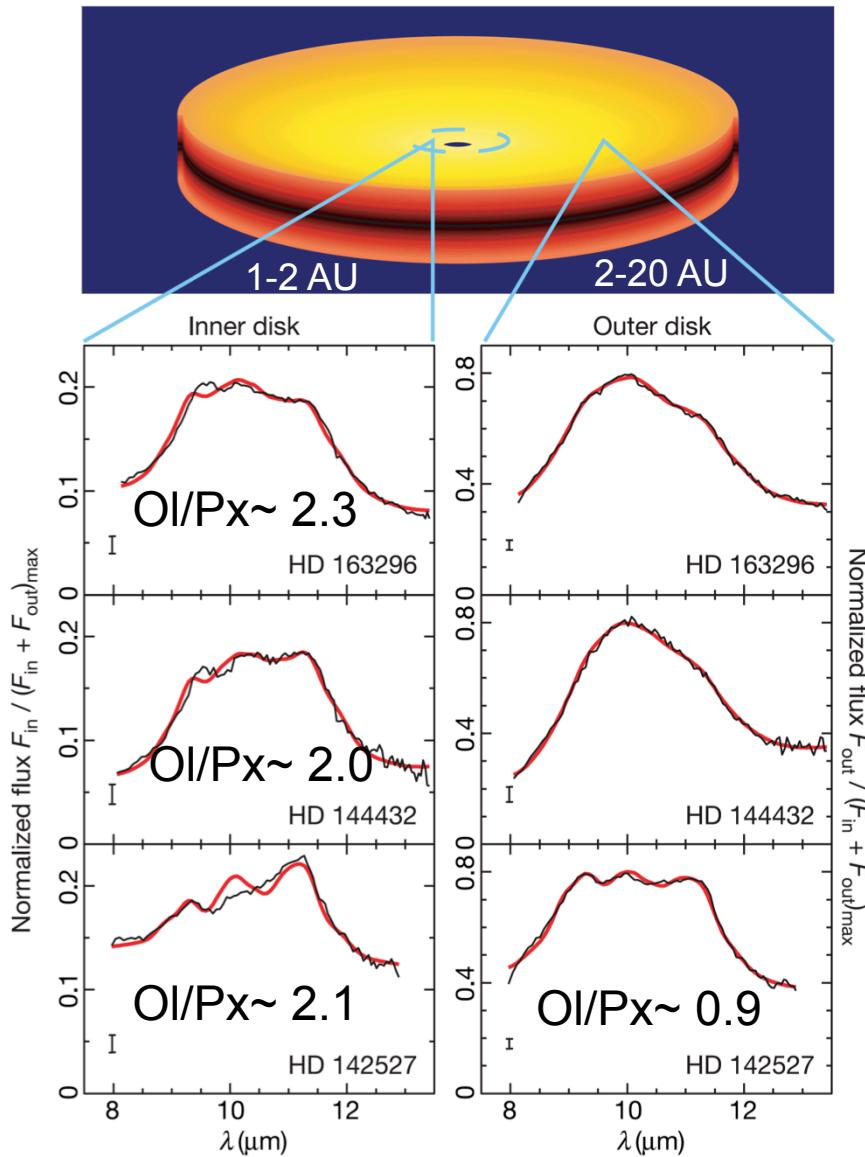


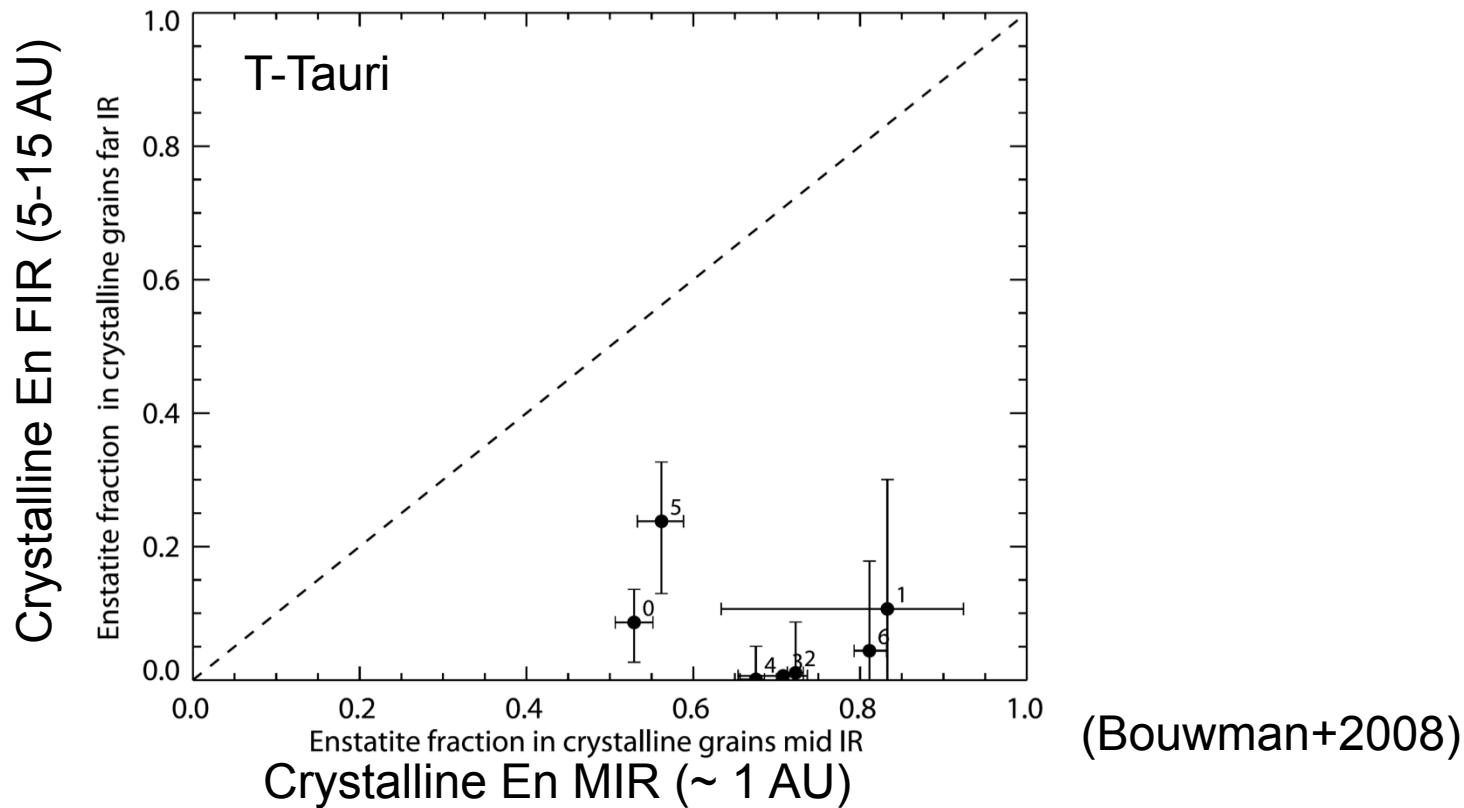
Table 1 Dust properties in the inner and outer disk

	Crystallinity (%)		Fraction of large grains (%)		Crystalline olivine to pyroxene ratio	
	Inner disk	Outer disk	Inner disk	Outer disk	Inner disk	Outer disk
HD 163296	40^{+20}_{-20}	15^{+10}_{-10}	95^{+5}_{-10}	65^{+20}_{-20}	$2.3^{+3.7}_{-0.5}$	-
HD 144432	55^{+30}_{-20}	10^{+10}_{-5}	90^{+10}_{-10}	35^{+20}_{-20}	$2.0^{+1.8}_{-0.6}$	-
HD 142527	95^{+5}_{-15}	40^{+20}_{-15}	65^{+15}_{-10}	80^{+10}_{-30}	$2.1^{+1.3}_{-0.7}$	$0.9^{+0.2}_{-0.1}$

(Van Boekel+2004)

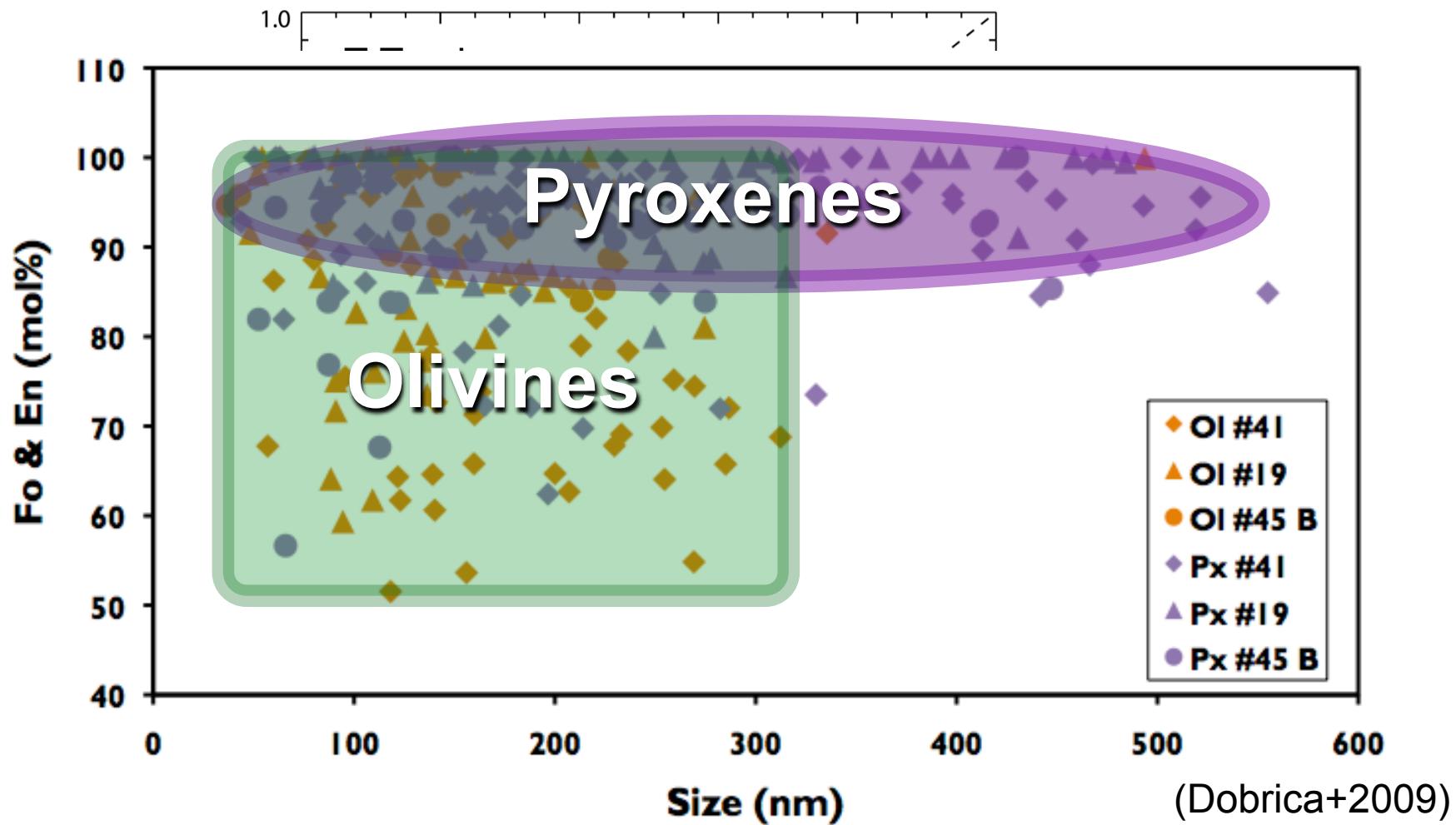
- Would agree with tendency observed in CP-IDPs, UCAMMs, Wild 2...
- But for T-Tauri...

Mineralogical zoning in PPD



- Crystalline OI/Px ~ 1.5 in inner regions (1 AU)
- Crystalline OI/Px > 5 in outer regions (5-15 AU)
- Enstatite grains ($\sim 1 \mu\text{m}$) larger than Fo grains ($\sim 0.1 \mu\text{m}$)
- (Amorphous silicates $\sim 6 \mu\text{m}$)

Mineralogical zoning in PPD



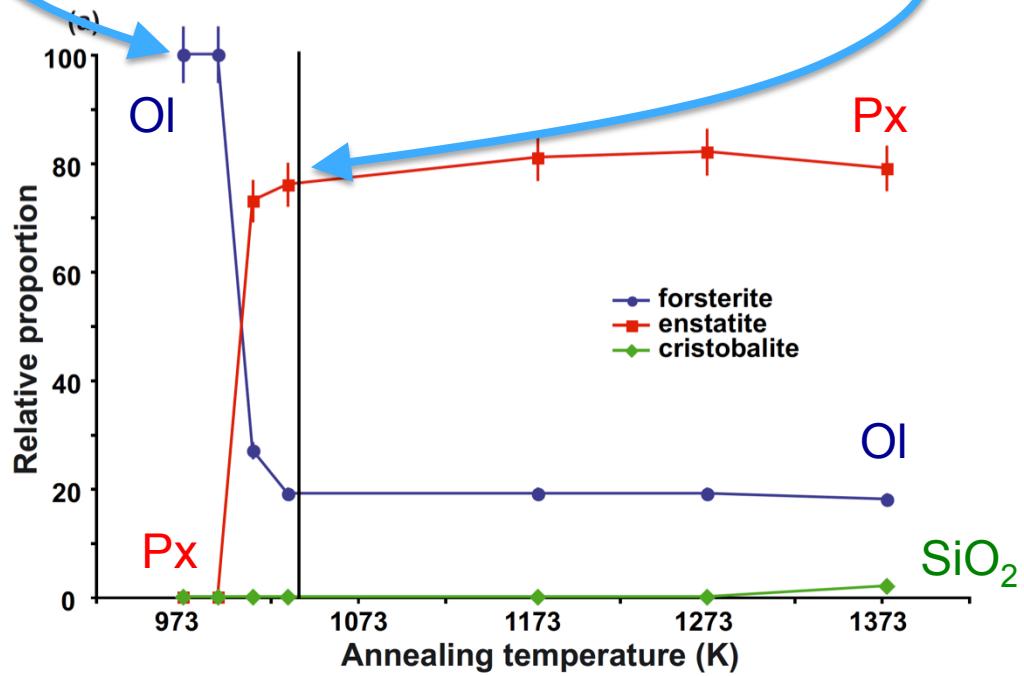
- Enstatite grains ($\sim 1 \mu\text{m}$) larger than Fo grains ($\sim 0.1 \mu\text{m}$)
- (Amorphous silicates $\sim 6 \mu\text{m}$)

Outward transport

- All minerals/CAIs in cosmic dust particles (IDPs/MMs) and Stardust samples (comet 81P/Wild2) are miniature compared to their meteorite counterparts
- A consequence of radial transport ?

Zonation of Ol and Px in PPDisk

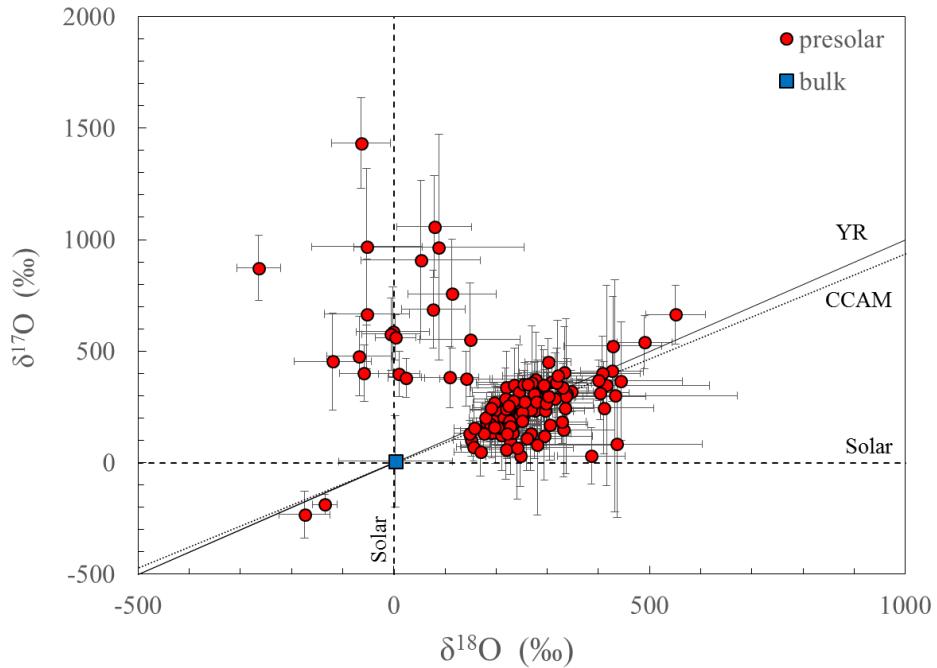
- Annealing of amorphous MgSiO_3 (Px) precursor ?
 - Formation of crystalline Fo
 - Then formation of crystalline Px
 - Silica formation as by-product



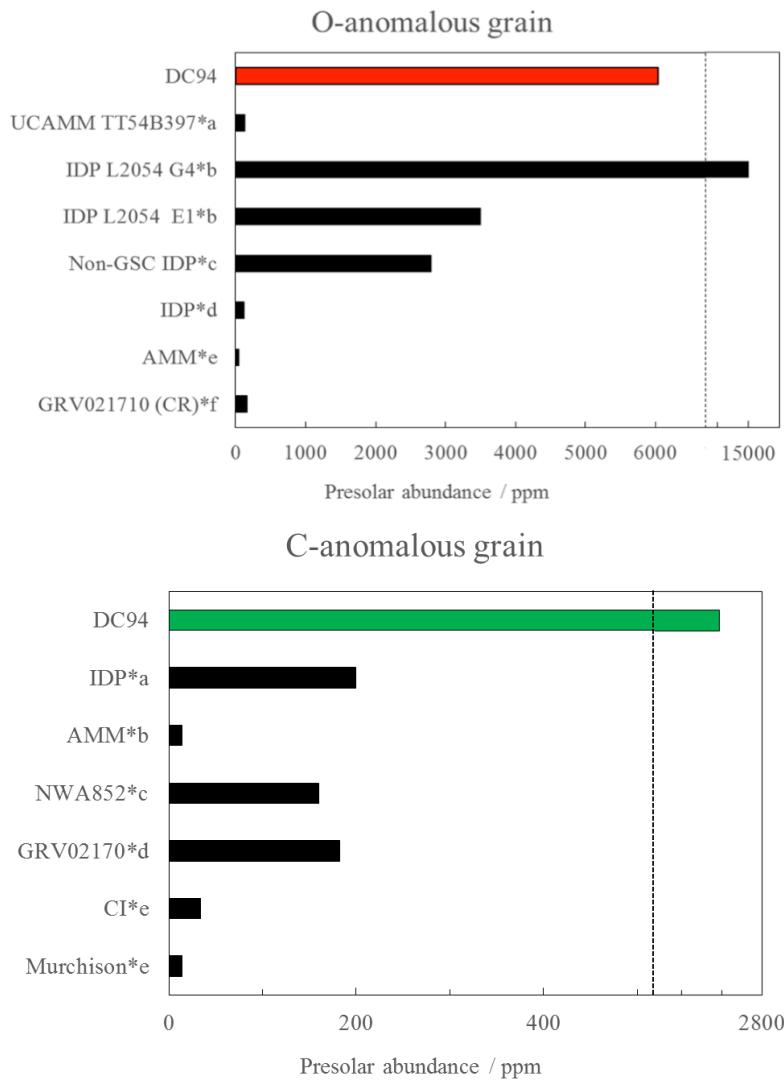
Roskoz+2011

- Crystalline phases in CP-IDPs/Wild2/UCAMMs preserved inner S.S. mineralogy during transport?
- But not meteorites?

Presolar grains in UCAMMs



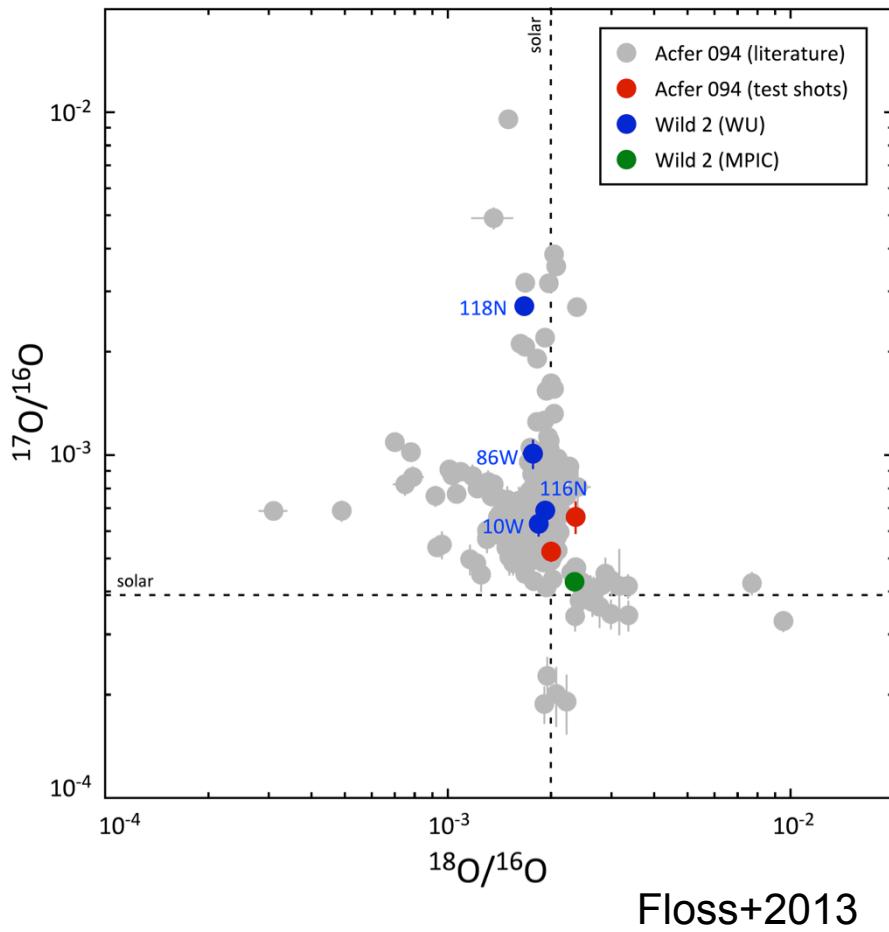
Kakazu et al. in prep



- Abundance of presolar grains $\sim 0.5\%$

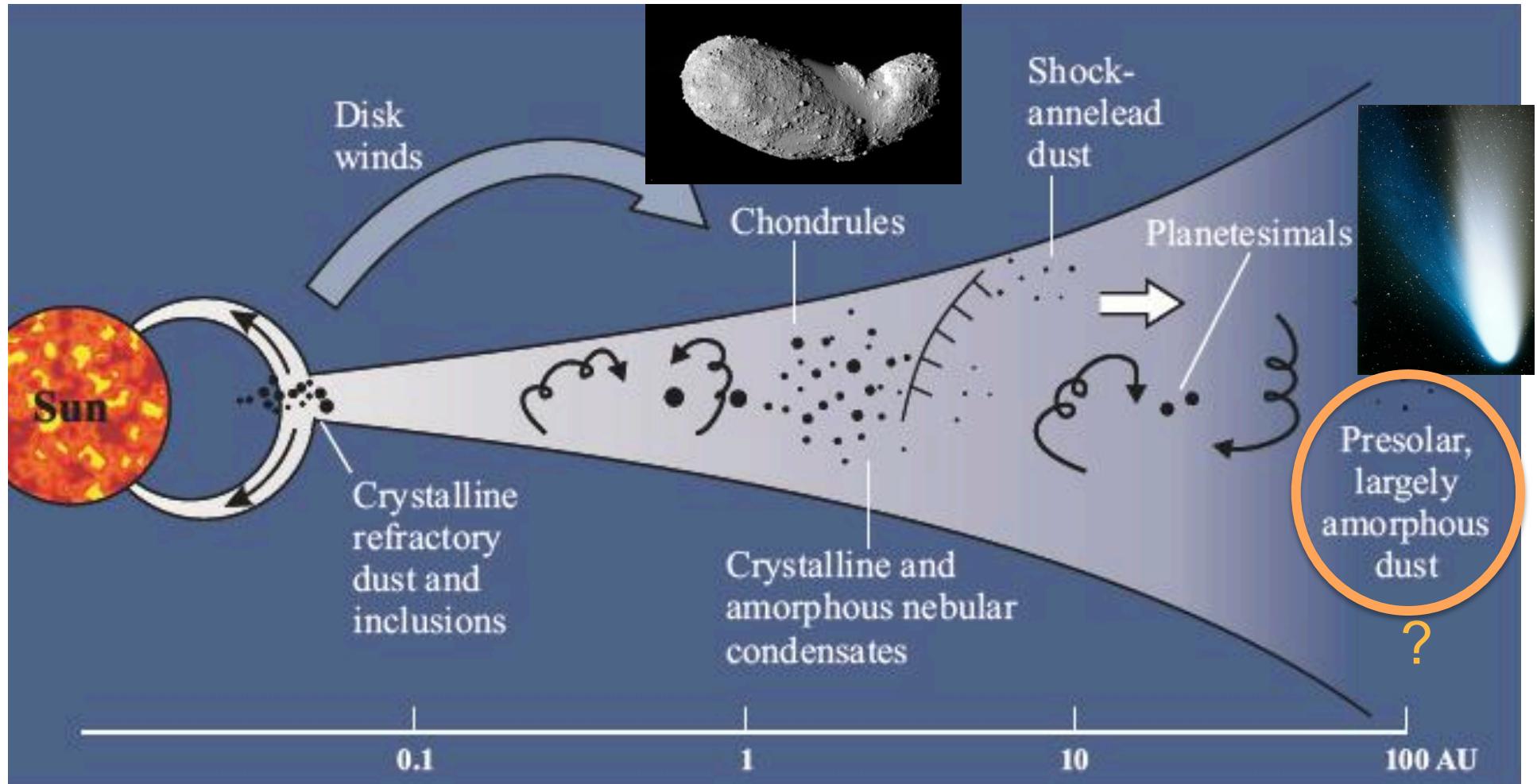
Stardust Presolar grains

- Stardust mission (81P/Wild2)
- anomalous silicates or oxides
 - corrected abundance (destruction impact) = 600-830 ppm
- SiC : 45 ppm



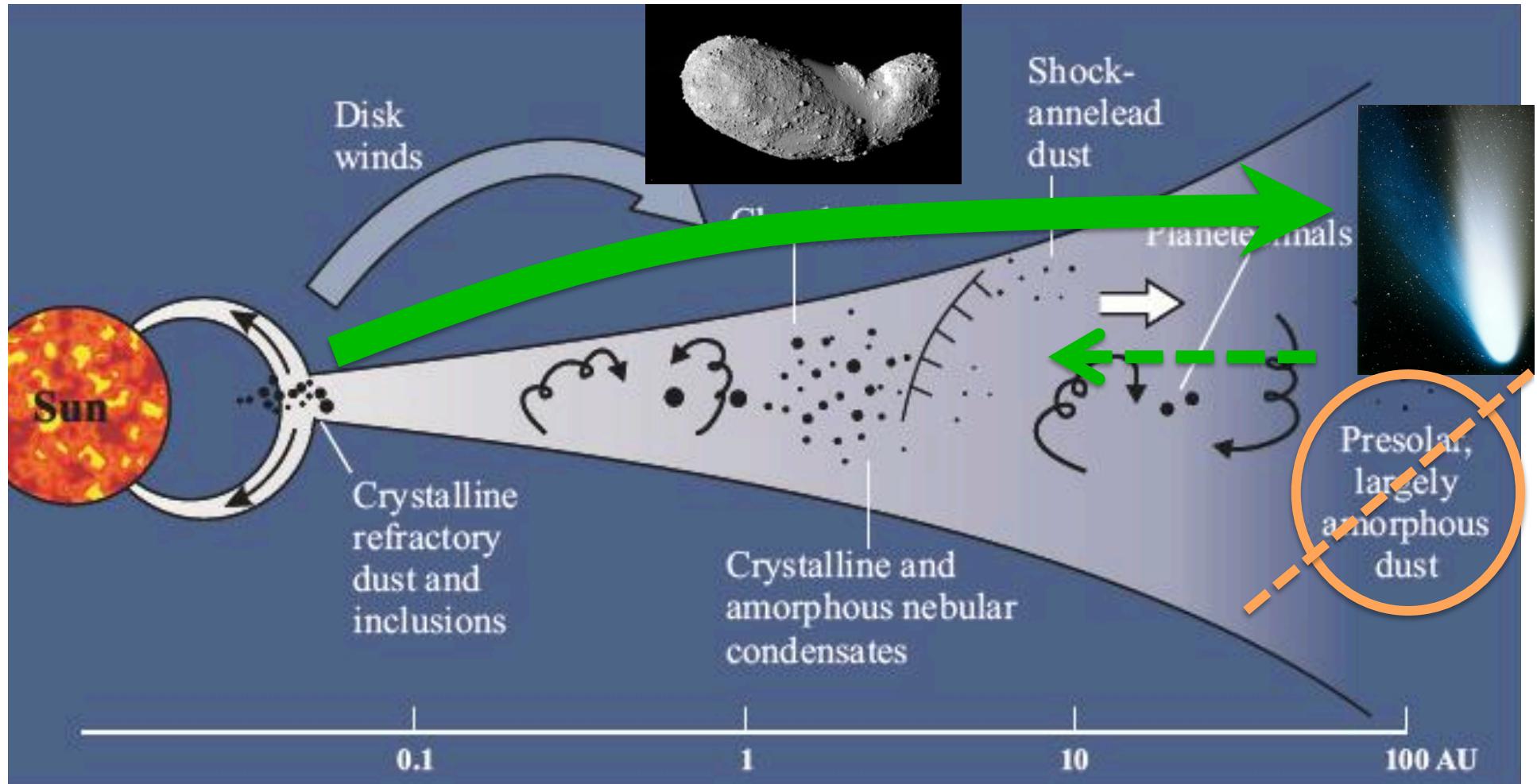
Floss+2013

Structure of the protoplanetary disk ?



(Scott, 2007)

Structure of the protoplanetary disk ?



(Scott, 2007)

Summary (1/2)

- UCAMMs provide informations on the outer regions of the protoplanetary disk
- 2 C-rich phases (at least) in UCAMMs: N-rich and N-poor (see also Yabuta et al., 2012)
 - N-rich carbonaceous phase is smooth, no minerals
 - N-poor carbonaceous phase is associated with minerals (and GEMS)
- Different origins/formation mechanisms for N-rich and N-poor phases
 - Genetic link to meteorite IOM for N-poor phase? (but contains GEMS...)
 - Formation by irradiation of $\text{CH}_4\text{-N}_2$ ices for N-rich phase? (Dartois+2013; Augé+2016)

Summary (2/2)

- Large D enrichments (low T chemistry) – no direct correlation with N isotopes
- « Large » abundance of presolar silicates (~0.5%) and C-rich anomalous grains
- Inner solar system material transported to the outer regions
 - Incomplete mixing from the inside out. OK with models (Shu et al. 1997, Bockelée-Morvan et al. 2002, Ciesla 2007, Vinković 2009...).
 - Mineral sizes usually smaller – size sorting effect?
 - Mineralogical gradient in the PPdisk ?

Open Questions

- Wild2 – CP-IDPs – UCAMMs :
 - Differences...=> ?
- Formation of GEMS – ISM or not ISM?
- Zonation of the PPD : CP-IDPs& UCAMMs richer in px than meteorites... signature? alteration effect?
- Atmospheric entry models: UCAMMs (outer S.S.) do not know evidence for heating above ~300°C?